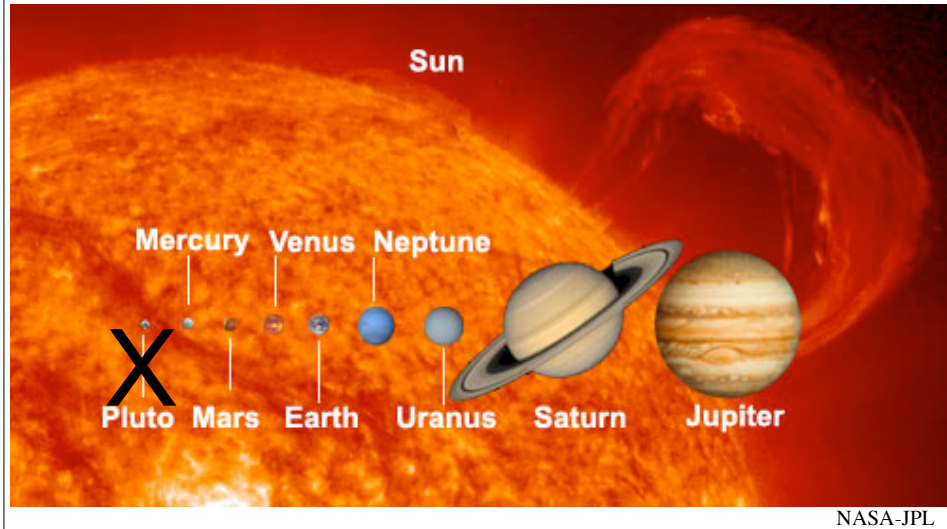


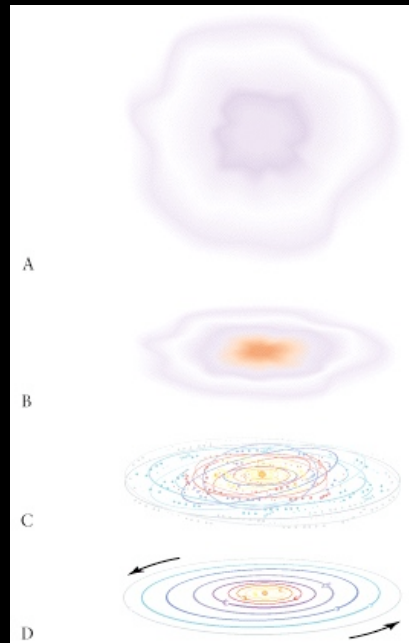
# Formation of the Solar System & the Structure of Earth

*OCEAN 355 Lecture Notes #2*



## Origin of Solar System from solar nebula

- Slowly rotating cloud of gas & dust
- Gravitational contraction
- High P=High T ( $PV=nRT$ )
- Rotation rate increases (conserve angular momentum--e.g., skaters)
- Rings of material condense to form planets (Accretion)

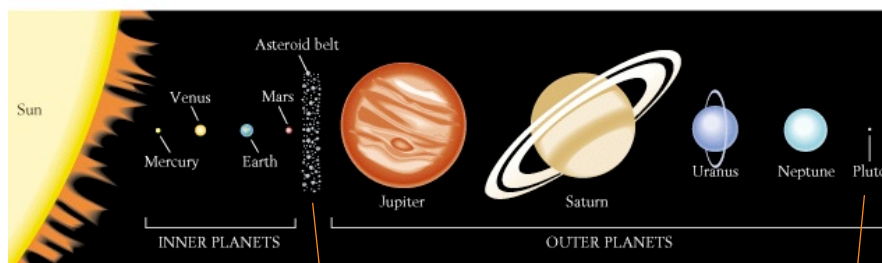
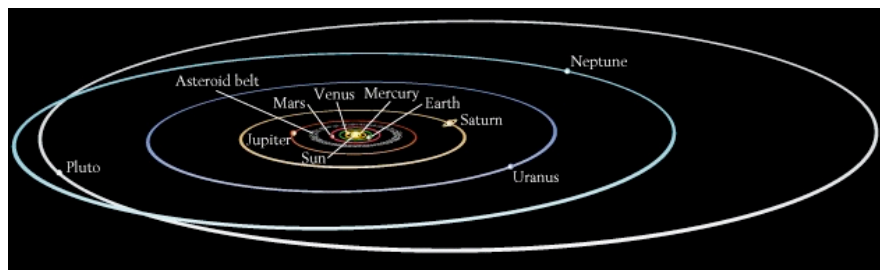


*Stanley (1999)*

## Observational Clues to the Origin of the Planets

- Inner planets are small and dense
- Outer planets are large and have low density
- Satellites of the outer planets are made mostly of ices
- Cratered surfaces are everywhere in the Solar System
- Saturn has such a low density that it can't be solid anywhere

## The Solar System



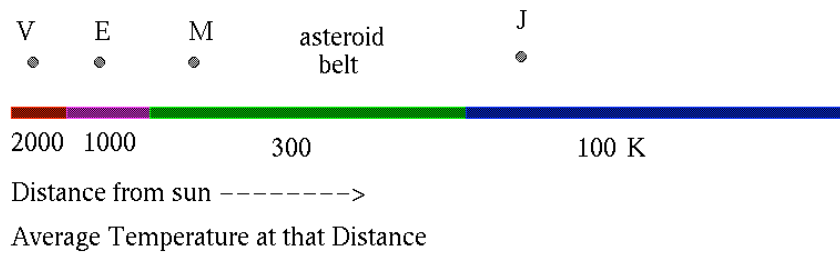
“Failed Planet”

Anomalous Planet

Stanley (1999)

## Formation of the Earth by Accretion: 1

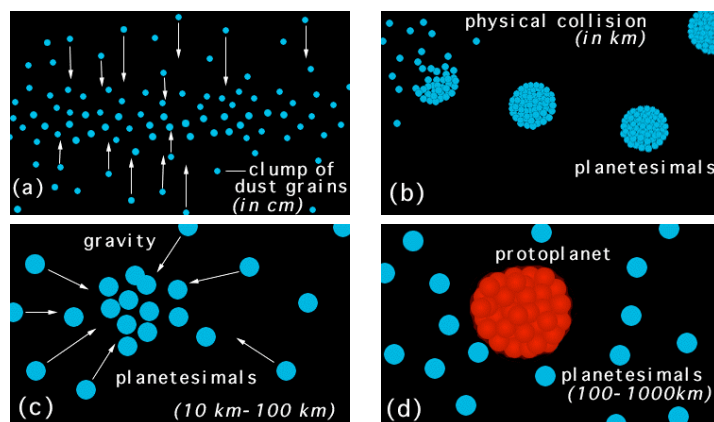
• **Initial solar nebula consisted of cosmic dust & ice with least volatile material condensing closest to the Sun and most volatile material condensing in outer solar system.**



<http://zebu.uoregon.edu/ph121/l7.html>

## Formation of the Earth by Accretion: 2

- [Step 1: accretion of cm sized particles](#)
- [Step 2: Physical Collision on km scale](#)
- [Step 3: Gravitational accretion on 10-100 km scale](#)
- [Step 4: Molten protoplanet from the heat of accretion](#)



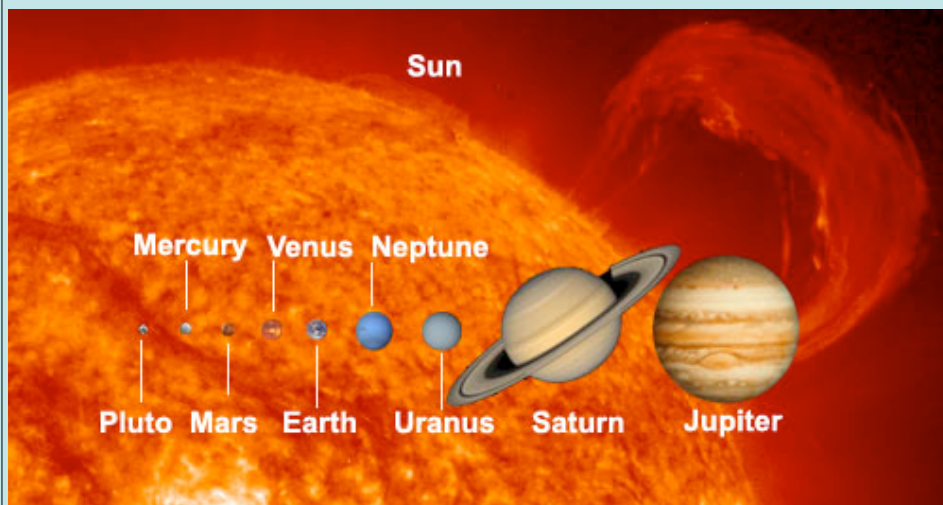
<http://zebu.uoregon.edu/ph121/l7.html>

## Formation of the Earth by Accretion: 3

- Tremendous heat generated in the final accretion process resulted in initially molten objects.
- Any molten object of size greater than about 500 km has sufficient gravity to cause gravitational separation of light and heavy elements thus producing a *differentiated* body.
- The accretion process is inefficient, there is lots of left over debris.
- In the inner part of the solar system, leftover rocky debris cratered the surfaces of the newly formed planets (*Heavy Bombardment*, 4.6-3.8 Ga).
- In the outer part of the solar system, the same 4 step process of accretion occurred but it was accretion of ices (cometisemals) instead of grains.

<http://zebu.uoregon.edu/ph121/17.html>

## The Sun & Planets to Scale

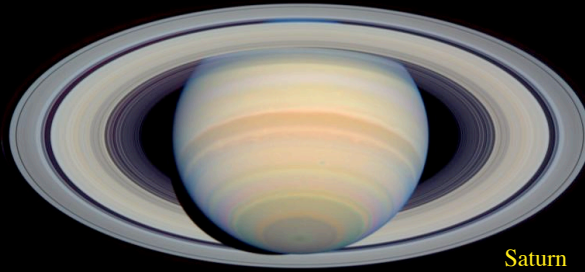


NASA-JPL



## More Planetary Eye Candy from Hubble

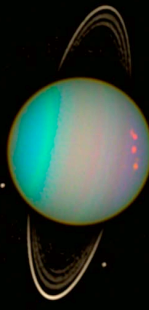
[http://hubblesite.org/gallery/album/solar\\_system\\_collection/4+1](http://hubblesite.org/gallery/album/solar_system_collection/4+1)



Saturn



Jupiter



Uranus



Mars

## After much outcry, NASA plans a final space shuttle servicing mission to HST in Aug. '08 to extend capabilities through 2013

In 2004, NASA administrators canceled a final servicing mission to the telescope that had been scheduled for 2006. The officials based the decision on concern for the safety of astronauts after the loss of the space shuttle Columbia in 2003. Scientists, politicians, and astronomy enthusiasts protested that the decision would bring an early end to the telescope's observations. NASA officials then agreed to study the possibility of sending a robotic craft to perform needed repairs.

STS-125: Final Shuttle Mission to Hubble Space Telescope



Image above: The astronauts selected for the final shuttle mission to perform work on the Hubble Space Telescope pose for a group photo. From left to right are astronauts Megan McArthur, Michael Good, Gregory C. Johnson, Scott Altman, John Grunsfeld, Michael Massimino and Andrew Feustel. Image credit: NASA |

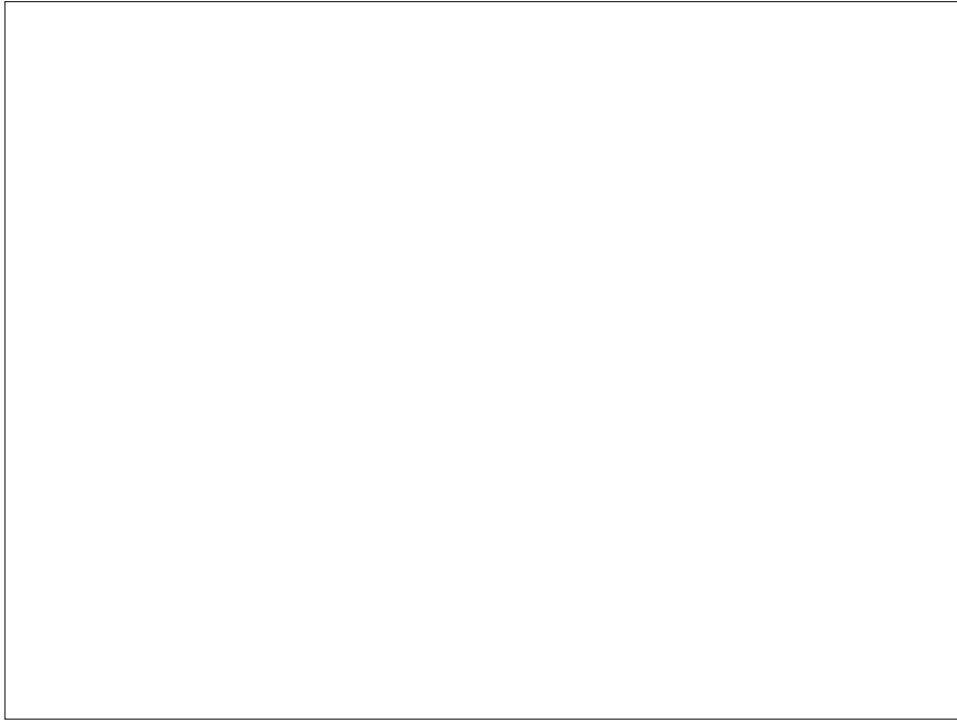
HOUSTON - NASA managers officially are targeting August 7, 2008, for the launch of the fifth and final space shuttle servicing mission to the Hubble Space Telescope. During the 11-day flight, Atlantis' seven astronauts will repair and improve the observatory's capabilities through 2013.

Mission planners have been working since last fall, when the flight was announced, to determine the best time in the shuttle manifest to support the needs of Hubble while minimizing the impact to International Space Station assembly. NASA also will support a "launch on need" flight during the Hubble mission. In the unlikely event a rescue flight becomes necessary, shuttle Endeavour currently is planned to lift off from Launch Pad 39-B at NASA's Kennedy Space Center, Fla. However, managers constantly are evaluating the manifest to determine the best mission options.

Shuttle missions beyond the Hubble flight, designated STS-125, still are being assessed. Shuttle and station program officials will continue to consider options for the remainder of the shuttle flights to complete construction of the space station by 2010, when the fleet will be retired. Those target launch dates are subject to change.

[http://www.nasa.gov/mission\\_pages/hubble/servicing/index.html](http://www.nasa.gov/mission_pages/hubble/servicing/index.html)

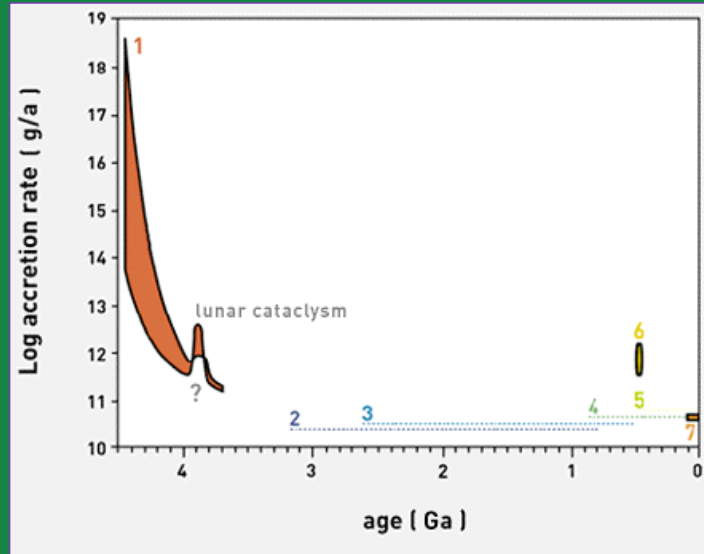
<http://spaceflight.nasa.gov/gallery/images/shuttle/sts-109/jucs/sts109-331-010.jpg>



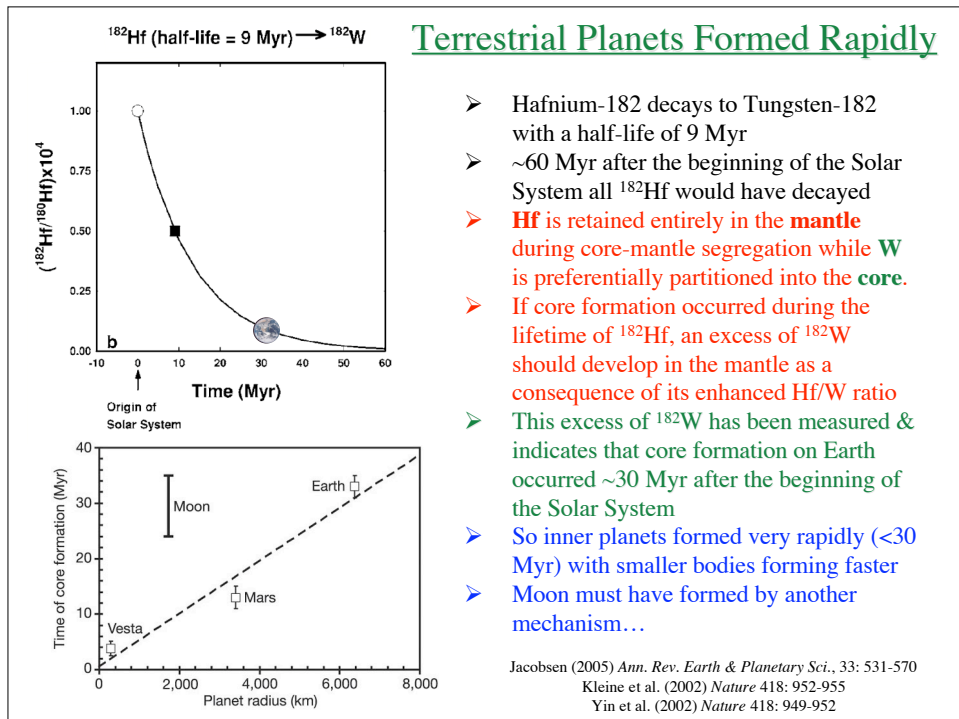
**\*\*\* Ended Here - 10/3/08 \*\*\***

Update Hubble repairs mission  
info

# Earth Accretion Rate Through Time



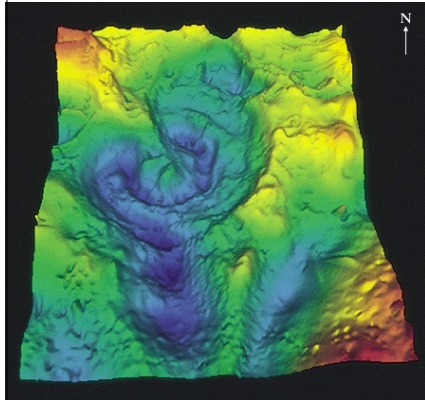
Schmitz et al., (1997) *Science*, Vol. 278: 88-90, and references therein.  
<http://www.whoi.edu/science/MCG/pge/project4.html>



# Accretion continues...

## Chicxulub Crater, Gulf of Mexico

- 200 km crater
- 10-km impactor
- 65 Myr BP
- Extinction of 75% of all species!

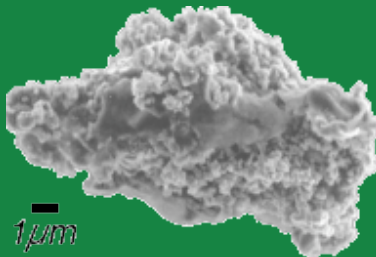


## Meteor (Barringer) Crater, Arizona

- 1 km diam. Crater
- 40-m diam Fe-meteorite
- 50 kyr BP
- 300,000 Mton
- 15 km/s

<http://www.gi.alaska.edu/rem sense/features/impactcrater/imageexplain.htm>

# Interplanetary Dust Accumulation



1  $\mu$ m

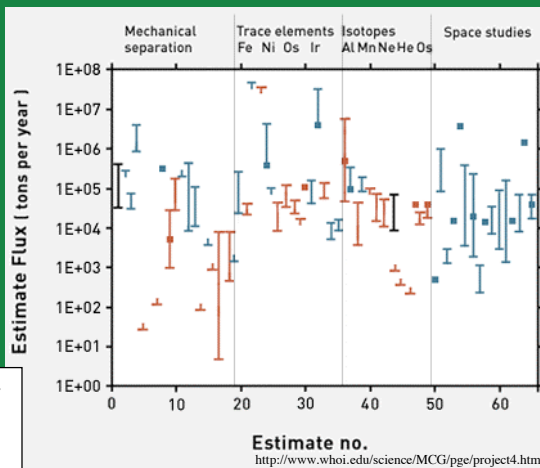
<http://presolar.wustl.edu/work/idp.html>

$40 \pm 20 \times 10^4$  metric tons/ yr  
( $40 \times 10^{10}$  g) interplanetary  
dust accretes every yr!

= 78.4 g/km<sup>2</sup>/yr

= 524.1

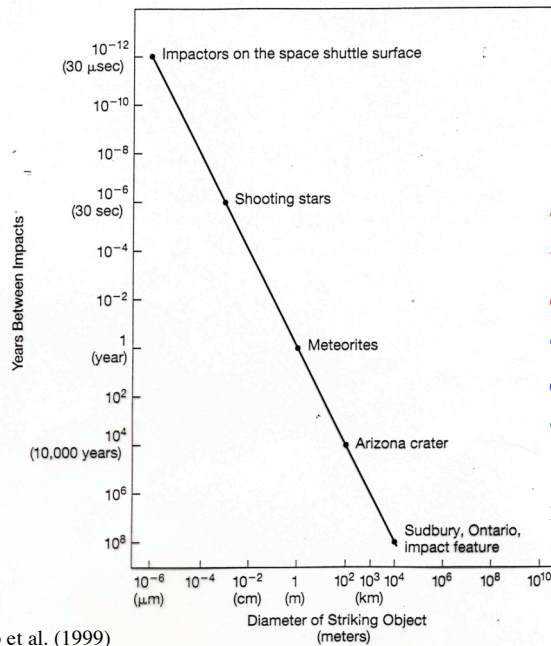
Ford Explorers/day



<http://www.whoi.edu/science/MCG/pge/project4.html>

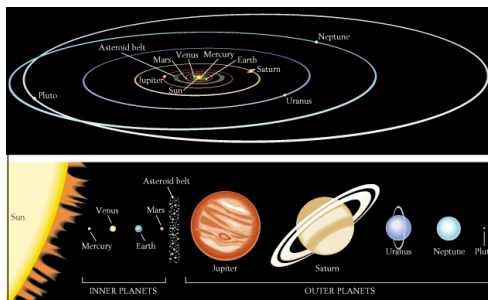


## Size & Frequency of Impacts



Kump et al. (1999)

- $40 \pm 20 \times 10^4$  metric tons/ yr interplanetary dust accretes every yr!
- 100 m object impacts every 10 kyr
- 10 km object every 100 Myr (e.g., K/T impact)



## The Asteroid Belt



- A relic of the accretion process. A failed planet.
- Gravitational influence of Jupiter accelerates material in that location to high velocity.
- High-velocity collisions between chunks of rock shatter them.
- The sizes of the largest asteroids are decreasing with time.

Total mass (Earth = 1)	0.001
Number of objects > 1 km	~100,000
Number of objects > 250 km	~12
Distance from Sun	2-4 AU
Width of asteroid belt (million km)	180

## Asteroid 243 IDA



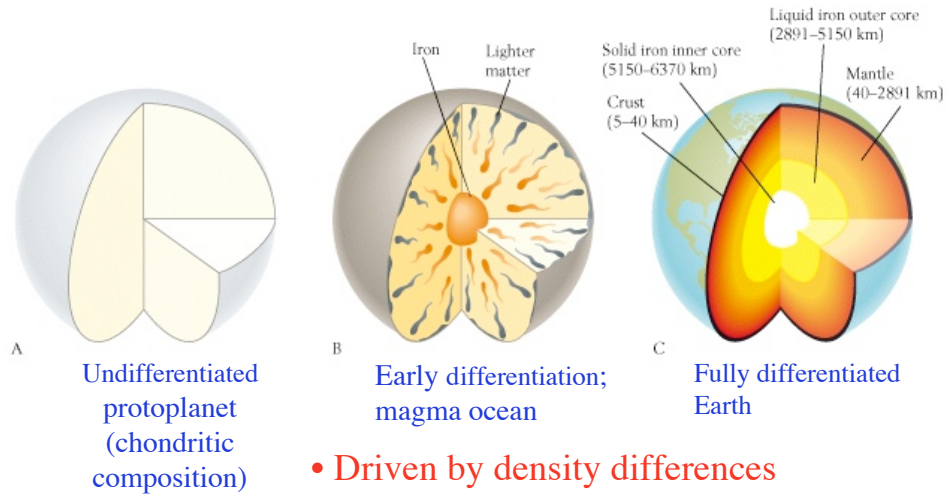
- Meteorite = asteroid that has landed on earth
- All chondrites (meteorites) date to ~4.5 B.y.
- Cratering indicates early origin

## Differentiation of the Earth: 1

- VM Goldschmidt (1922) published landmark paper “Differentiation of the Earth”:
  1. Earth has a chondritic (meteoritic) elemental composition.
  2. Surface rocks are not chemically representative of solar abundances, therefore must be differentiated.
- Proto-planet differentiated early into a dense iron-rich core surrounded by a metal sulfide-rich shell above which floated a low-density silicate-rich magma ocean.
- Cooling of the magma caused segregation of dense silicate minerals (pyroxenes & olivines) from less dense minerals (feldspars & quartz) which floated to surface to form crust.
- In molten phase, elements segregate according to affinities for:  
Fe = siderophile, sulfide = chalcophile & silicate = lithophile.

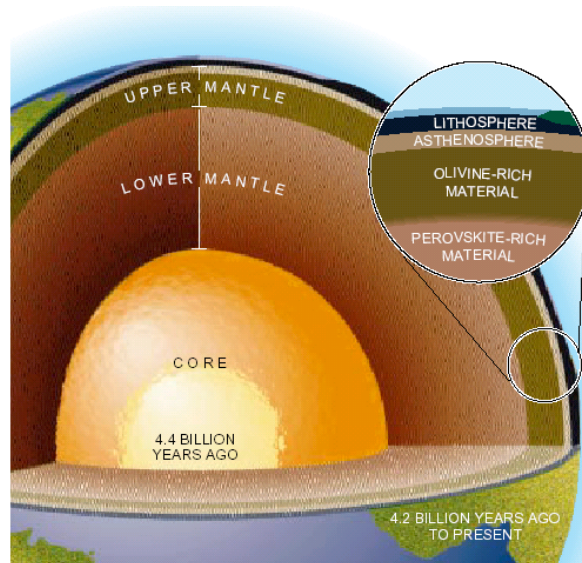


## Differentiation of Earth: 2



- Driven by density differences
- Occurred in ~30 Myr ( $^{182}\text{Hf}/^{180}\text{Hf}$ , c.f., Jacobsen (2005) *Ann. Rev. Earth & Planetary Sci.*, 33: 531-570)

## Differentiation of the Earth: 3



Allegre & Schneider, *Sci. Am.*(1994)

• **Differentiation of Earth**

Homogenous planetesimal

Earth heats up

Accretion and *compression* (T~1000°C)

*Radioactive decay* (T~2000°C)

Iron melts--migrates to center, forming *core*

*Frictional heating* as iron migrates

Light materials float--*crust*

Intermediate materials remain--*mantle*

Differentiation  
of Earth,  
Continents,  
Ocean &  
Atmosphere

• **Differentiation of Continents, Oceans, and Atmosphere**

Continental crust forms from differentiation of primal crust

Oceans and atmosphere

*Two hypotheses*

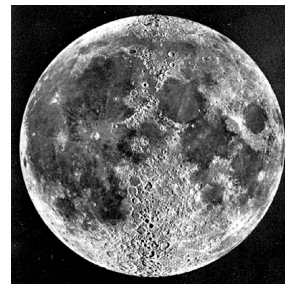
*internal*: degassing of Earth's interior (volcanic gases)

*external*: comet impacts add H<sub>2</sub>O CO<sub>2</sub>, and other gases

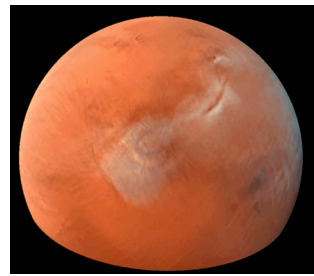
Early atmosphere rich in H<sub>2</sub>, H<sub>2</sub>O, N<sub>2</sub>, CO<sub>2</sub>; deficient in O<sub>2</sub>

**Earth's Crustal Evolution: 1.  
3 Types of Planetary Crust**

1° = original crystalline material to solidify from magma oceans of newly accreted bodies. None of this survives on Earth, but the white highlands of the moon are a good example. Impact that created moon produced 1° crust.



2° = slow heating by *radioactive decay* melts small quantities of rock in planetary interiors. Results in eruption of basaltic lavas. E.g., Earth's ocean floor, surfaces of Mars & Venus, lunar maria.



Taylor &  
McLennan  
(1996);  
NASA-JPL

## Origin of the Moon (Artist's Rendition!)



Stanley (1999)

## Moon-Forming Impact

Canup R & Aspaug E: Eos Trans. AGU, 82(47), Fall Meet. Suppl., Abstract U51A-02, 2001  
<http://www.swri.edu/9what/releases/canupmoon.htm>

Hypothesis for lunar origin - Moon forms from debris ejected as a result of the collision of a roughly Mars-sized impactor with early Earth

- Geophysical simulations use a method known as **smooth particle hydrodynamics**, or **SPH** and can achieve resolutions sufficient to study the production of orbit-bound debris necessary to yield the Moon.
- Off-center, low-velocity collisions yield material in bound orbit from which a satellite may then accumulate.
- Simulations must account for mass, angular momentum and compositions of the earth-Moon system.
- Must yield an Earth that retains an iron-rich core and a moon that is appropriately iron-depleted and the right density.

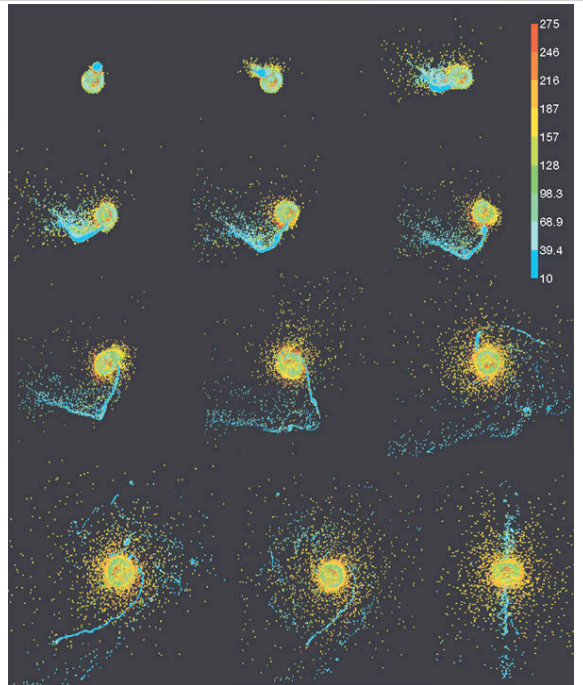
SPH results suggest:

- The object had 10-12% of Earth's mass (Mars-size!)
- Produces a satellite with <3% Fe by mass. Unable to be subsequently captured.
- Happened near end of Earth's accretional history.
- Resulted in melting of Earth crust.

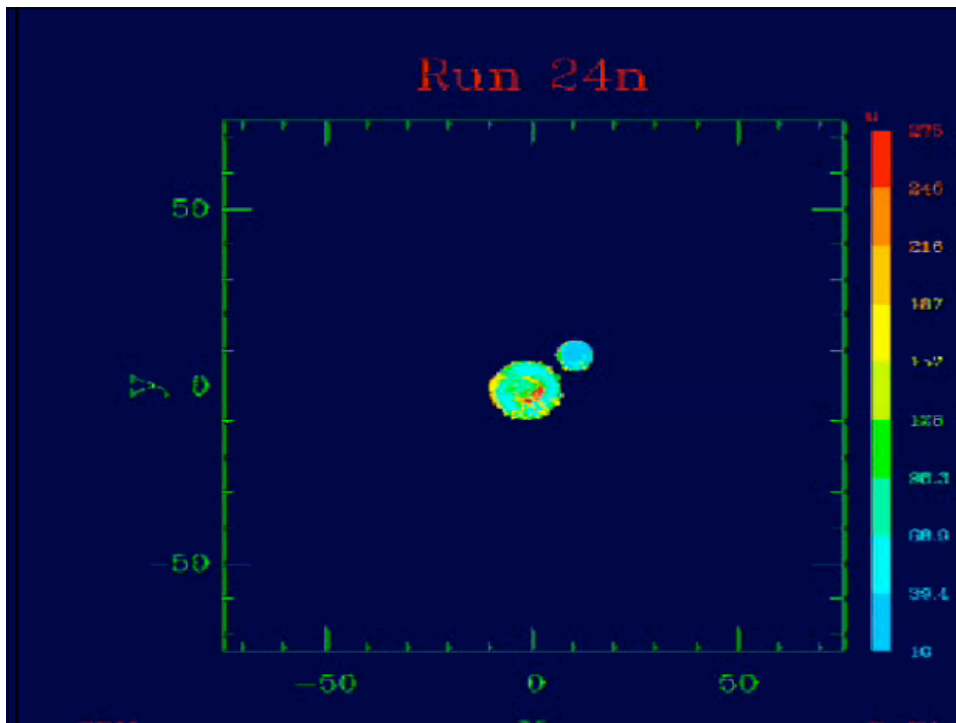
# Numerical Simulation of Moon-Formation Event

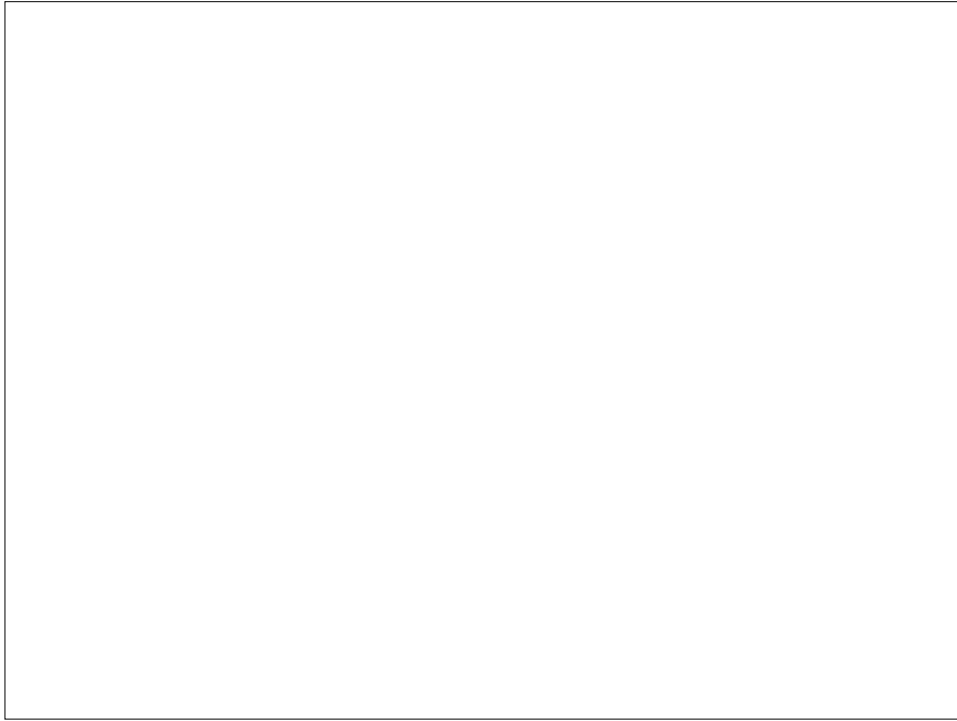
- Mars-size object ( $10\% M_E$ ) struck Earth
- core merged with Earth
- Moon coalesced from ejected debris of impactor's mantle

- Explains high rotation rate of Earth
- Heat of impact melted crust
- magma ocean #2

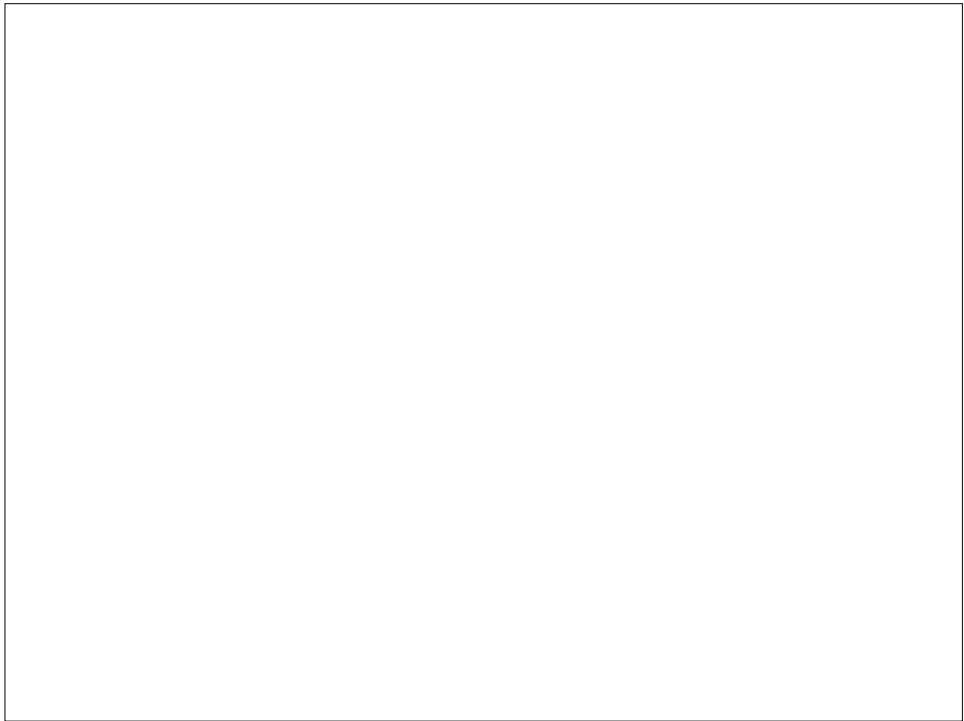


Canup & Asphaug (2001), *Nature*, Vol. 412.



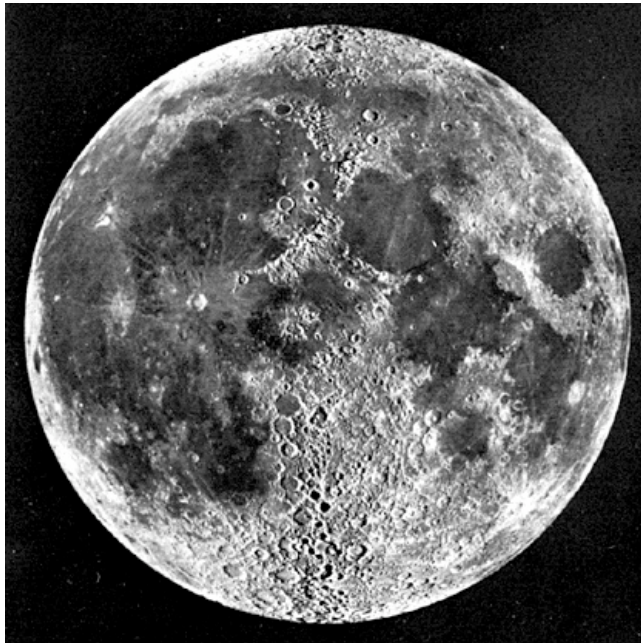


**\*\*\* Ended Here 10/6/08 \*\*\***



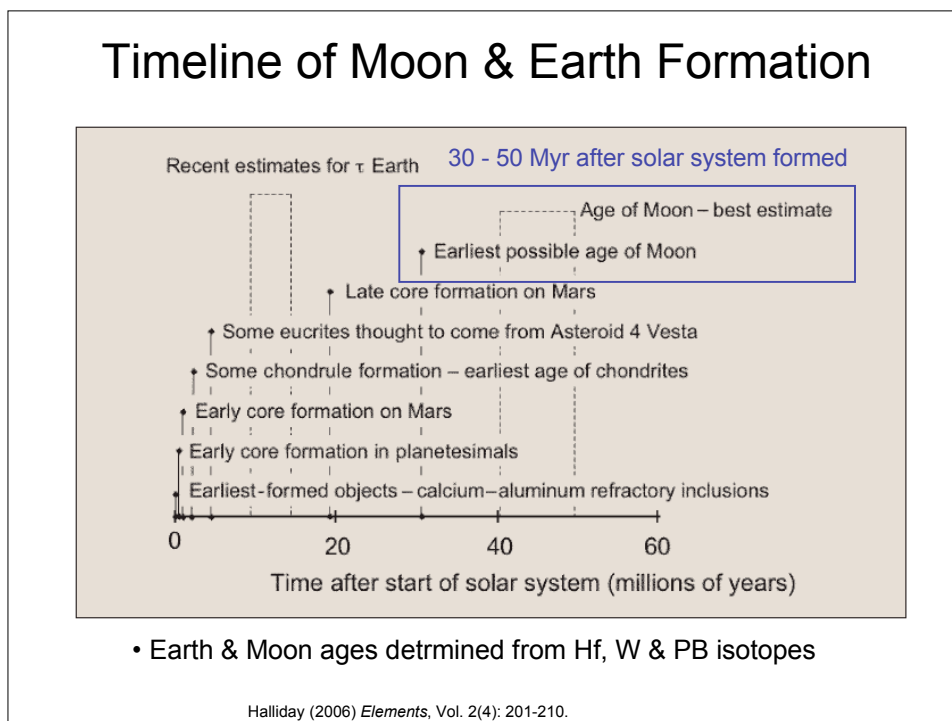
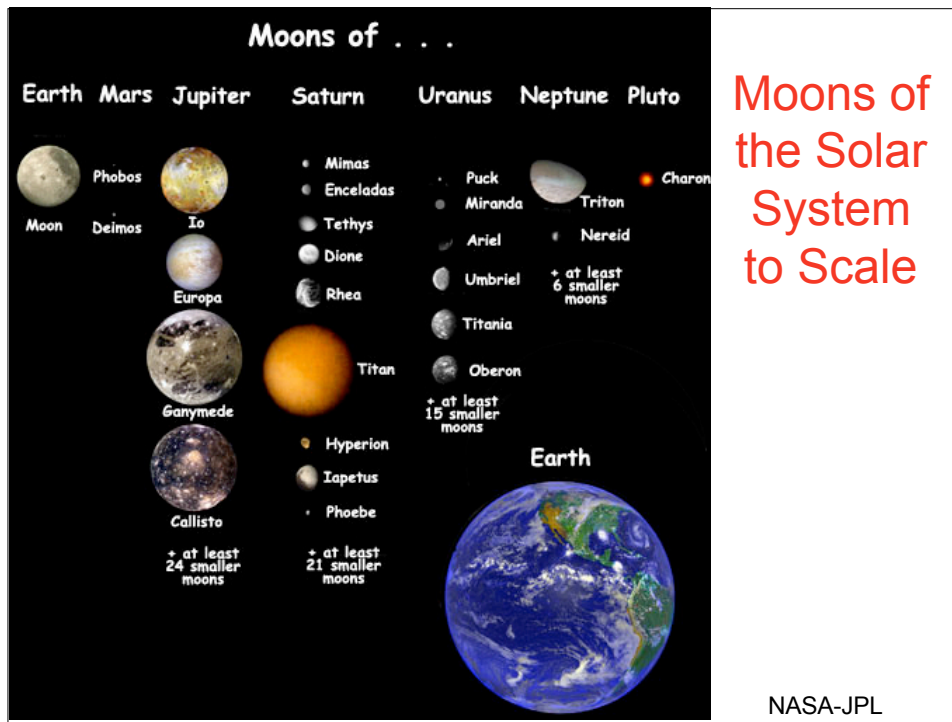
## The Moon

- Critical to life (stabilizes tilt)
- Rocks from crater rims are 4.6-4.0 Ba (heavy bombardment)
- Jupiter's gravity shielded Earth & Moon from 1,000x more impacts!
- Enormous 'seas' ("**maria**" in latin) are large craters; impact caused melting of crust
- Avg crater diam = 200 km



NASA-JPL

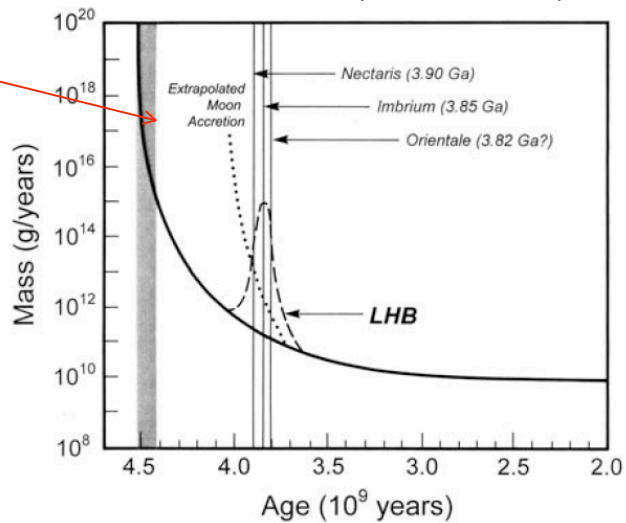






## Timeline of Moon Accretion Rate (= mass flux)

- Formation age determined from recent Hf & W isotope measurements
- LHB = Late Heavy Bombardment period
- Ages of largest impact basins shown
- Solid line is present-day background flux extrapolated back to 4.55 Ga

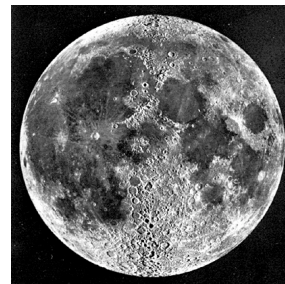


- Dotted line is accretion curve that includes masses of the basin-forming projectiles; unlikely to be correct, b/c implies accretion at 4.1 Ga
- Supports notion of a spike in large impacts (the LHB), possibly associated with migration of Uranus & Neptune towards outer solar system & disturbance of Kuiper Belt objects.

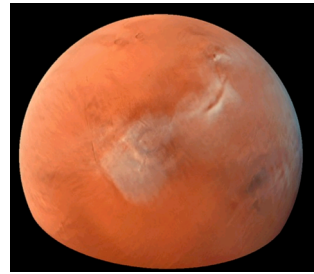
Koeberl (2006) *Elements* Vol. 2(4): 211-216

## Earth's Crustal Evolution: 1. 3 Types of Planetary Crust

1° = original crystalline material to solidify from magma oceans of newly accreted bodies. None of this survives on Earth, but the white highlands of the moon are a good example. Impact that created moon produced 1° crust.



2° = slow heating by radioactive decay melts small quantities of rock in planetary interiors. Results in eruption of basaltic lavas. E.g., Earth's ocean floor, surfaces of Mars & Venus, lunar maria.

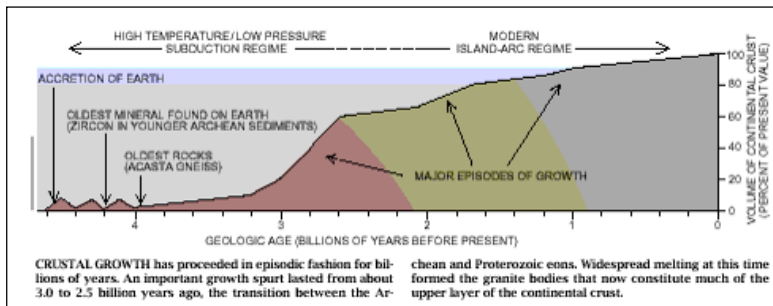


Taylor & McLennan (1996); NASA-JPL

## Earth's Crustal Evolution: 2

3° Crust = Formed from slow, continuous distillation by volcanism on a geologically active planet (I.e., plate tectonics).

- Results in highly differentiated magma distinct from basalt. I.e., the low-density, light-colored granite we see in rocks on the continents.
- Earth may be the only planet where this type of crust exists.
- Unlike 1° & 2° crusts, which form in < 200 Myr, 3° crusts evolve over billions of years.



Taylor & McLennan (1996)

## Igneous Rocks

**Basalt**  
(2° Crust;  
Oceanic  
crust)



**Granite**  
(3° Crust;  
Continental  
Crust)

Stanley (1999)

## The Crust

### Ocean Crust (the cheese)

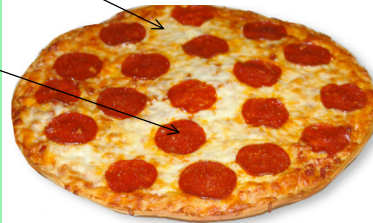
3-15 km thick  
Basaltic rock  
Young (<180 Ma)  
Density ~ 3.0 g/cm<sup>3</sup>

### Continental Crust (the pepperoni)

35 km average thickness  
Granitic rock  
Old (up to 3.8 Ga)  
Density ~ 2.7 g/cm<sup>3</sup>

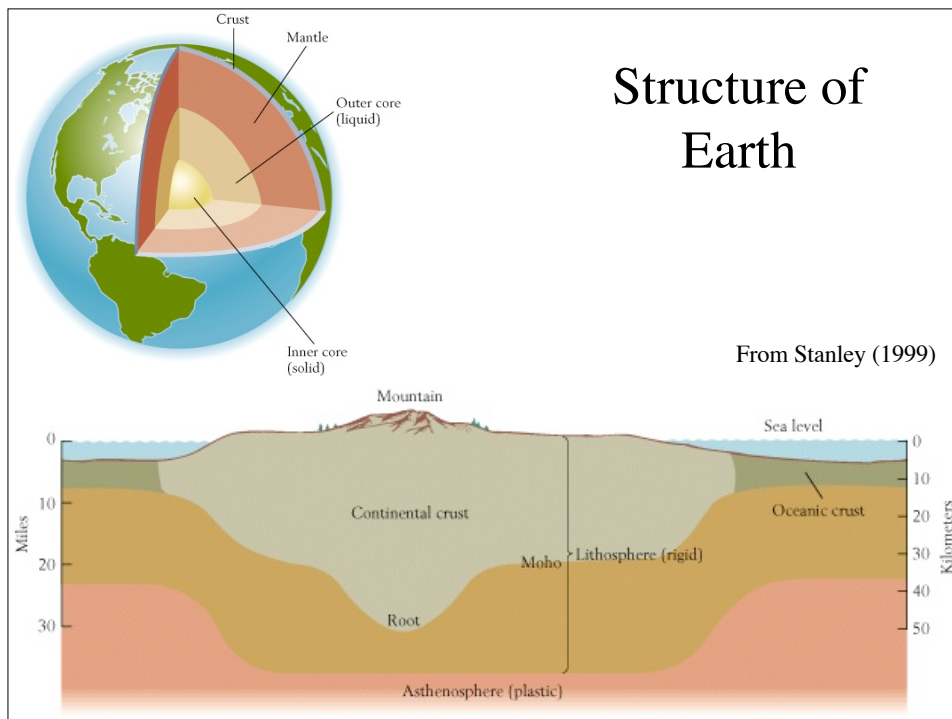
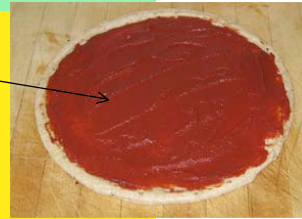
Crust "floating" on "weak" mantle

## The Crust & Mantle

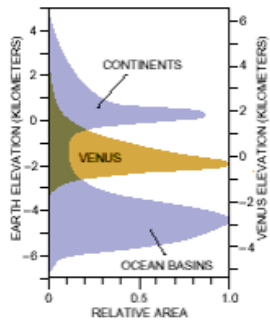


## The Mantle (the sauce)

~2900 km thick  
Comprises >82% of Earth's volume  
Mg-Fe silicates (rock)  
Two main subdivisions:  
Upper mantle (upper 660 km)  
Lower mantle (660 to ~2900 km; "Mesosphere")



## Why is Continental Crust “Elevated Relative to Oceanic Crust?”



SURFACE ELEVATIONS are distributed quite differently on the earth (blue) and on Venus (gold). Most places on the earth stand near one of two prevailing levels. In contrast, a single height characterizes most of the surface of Venus. (Elevation on Venus is given with respect to the planet's mean radius.)

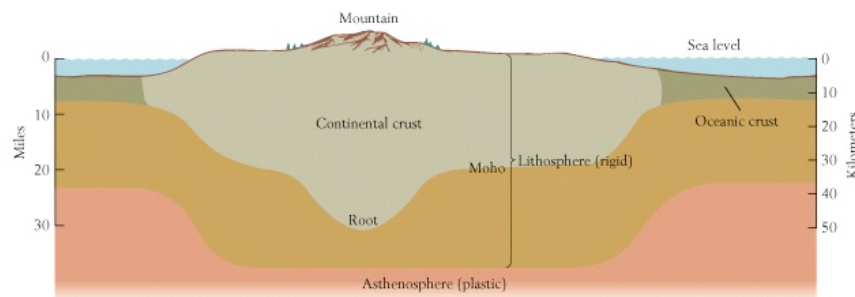
- High-density Basalt sinks into mantle more than low-density Granite.
- Volcanism continually produces highly differentiated continental crust on Earth.
- Venus surface appears to be all basalt.
- Plate tectonics & volcanism do not appear to be happening on Venus (or Mars, Moon).
- So Earth may be unique in Solar System. And plate tectonics & volcanism may be critical in determining habitability.

Taylor & McLennan *Sci. Am.* (1996)

## Lithosphere & Asthenosphere

Lithosphere/Asthenosphere: Outer 660 km divided into 2 layers based on mechanical properties. Includes the Mantle + Crust

- **Lithosphere**  
Rigid outer layer including crust & upper mantle  
Averages 100 km thick; thicker under continents
- **Asthenosphere**  
Weak, ductile layer under lithosphere  
Lower boundary ~660 km (entirely within mantle)



## The Core

- **Outer Core**

~2300 km thick

*Liquid* Fe with Ni, S, O, and/or Si

Magnetic field is evidence of flow

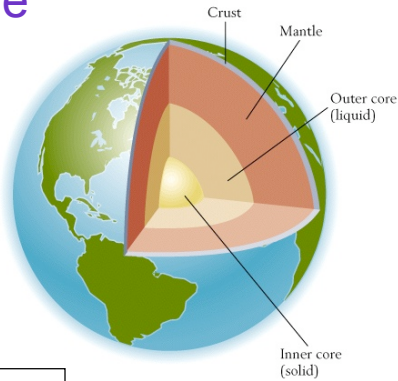
Density ~ 11 g/cm<sup>3</sup>

- **Inner Core**

~1200 km thick

*Solid* Fe with Ni, S, O, and/or Si

Density ~13.5 g/cm<sup>3</sup>



- **Earth's Interior: How do we know its structure?**

- Avg density of Earth (5.5 g/cm<sup>3</sup>)
- Denser than crust & mantle
- Composition of meteorites
- Seismic wave velocities
- Laboratory experiments
- Chemical stability
- Earth's magnetic field

From Stanley (1999)

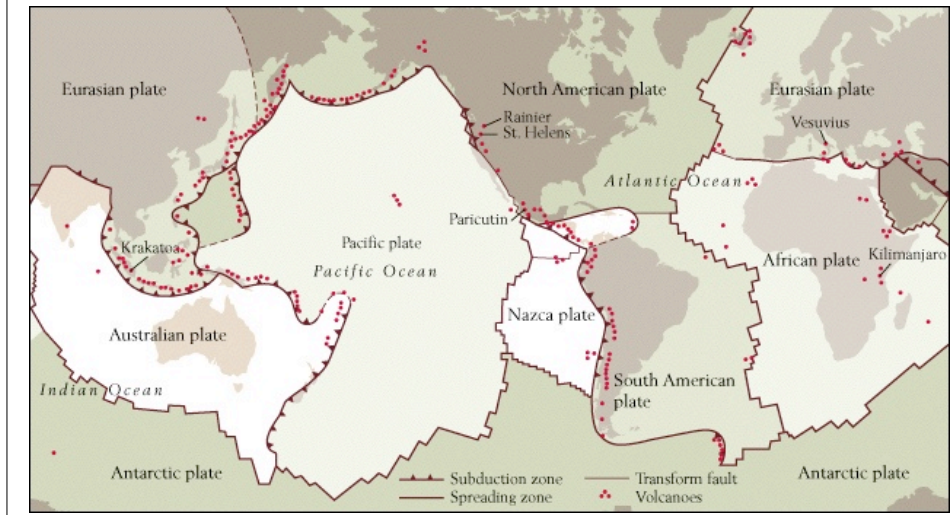
# Basics of Geology



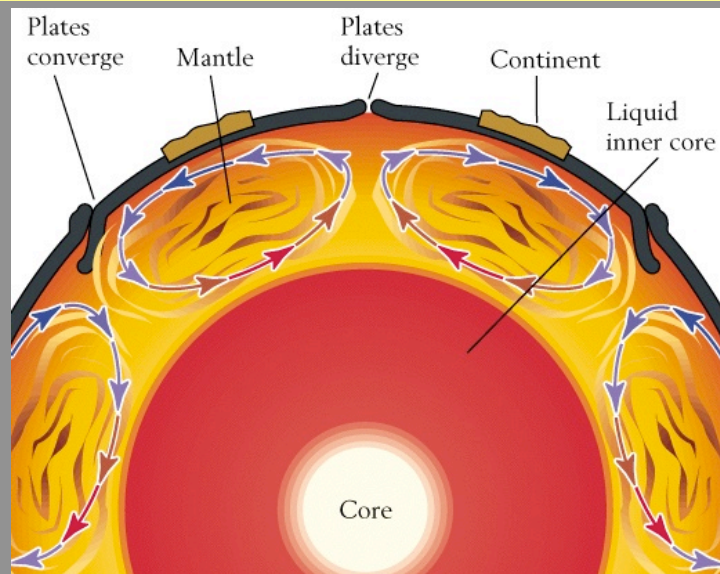
# Lithospheric Plates

From Stanley (1999)

- 8 large plates (+ add'l. small ones)
- Average speed: 5 cm/yr
- 3 types of motion result in 3 types of boundaries: sliding toward (subduction zones), sliding away (ridge axes), sliding along (transform faults)

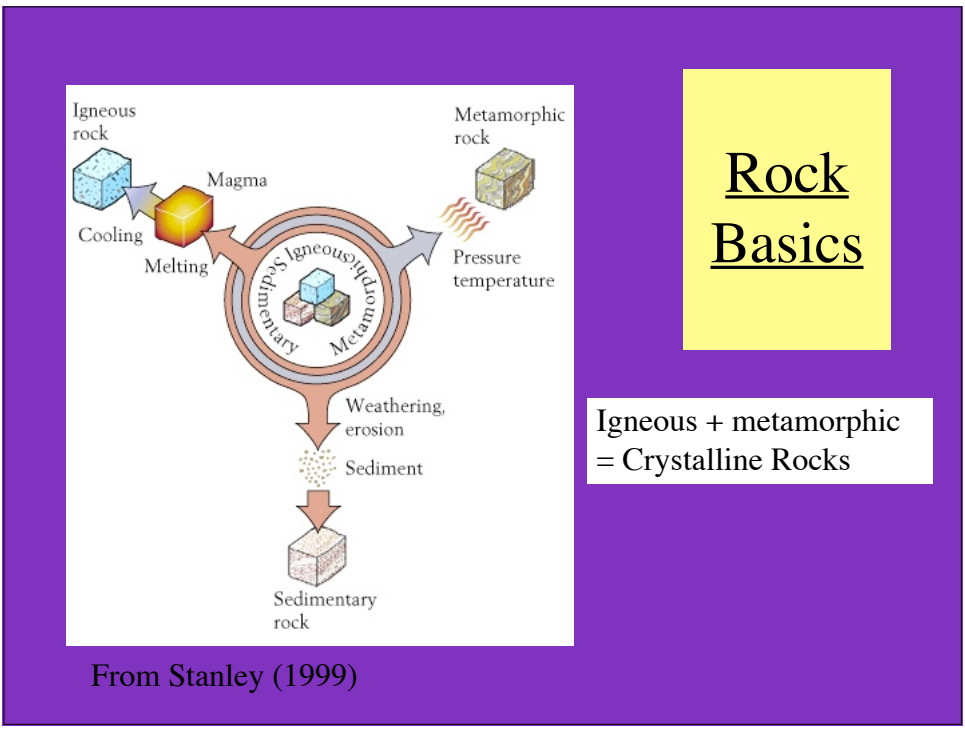
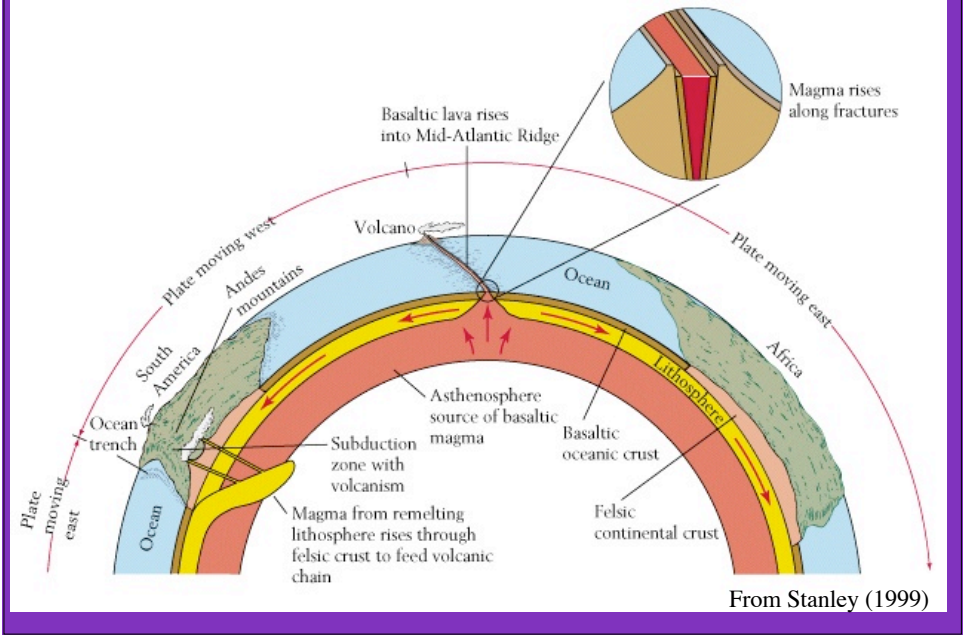


## Convection Drives Plate Movements



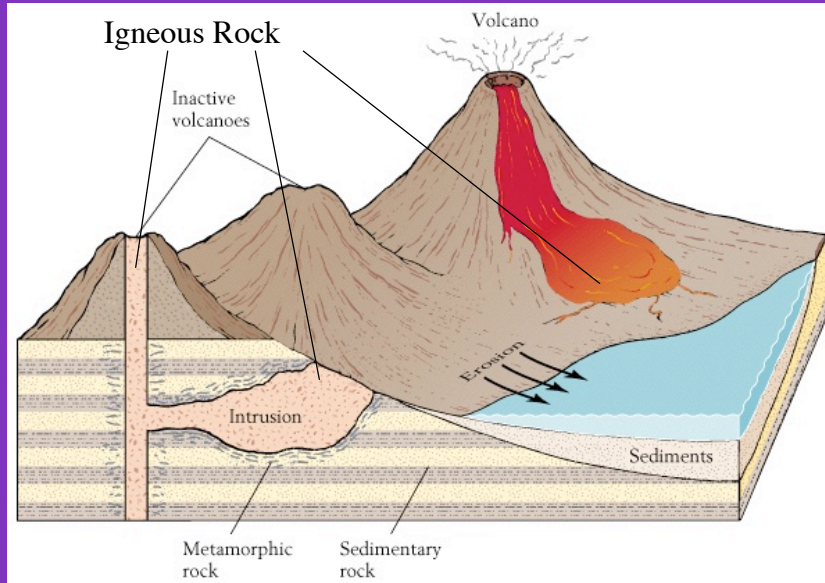
From Stanley (1999)

# Tectonic Activity in the South Atlantic





# The Rock Cycle



From Stanley (1999)

- **Felsic:** Si-, Al-rich. Light-colored, low-density. Feldspar (pink) & quartz (SiO<sub>2</sub>)-rich. Most continental crust. Granite most abundant.
- **Mafic:** Mg-, Fe-rich. Dark-colored, high-density. Most oceanic crust. Ultramafic rock (more dense) forms mantle below crust.
  - **Extrusive:** cools rapidly; small crystals
  - **Intrusive:** cools slowly; large crystals

## Igneous Rocks 101

**Basalt**  
(Oceanic Crust)



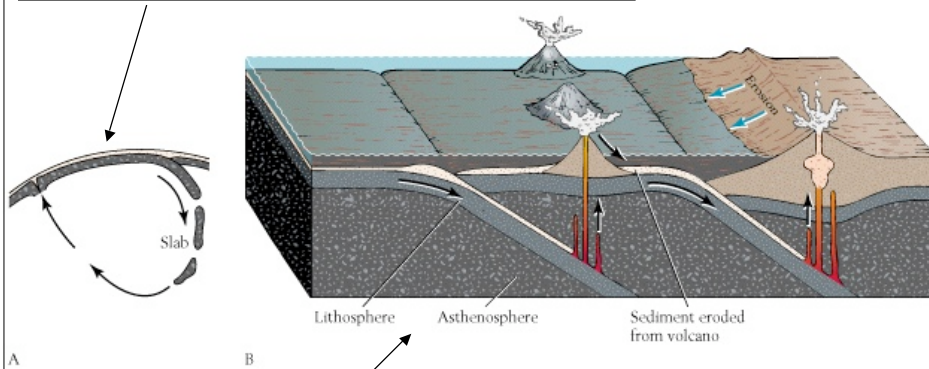
**Granite**  
(Continental Crust)

Stanley (1999)

ME MI  
FE FI

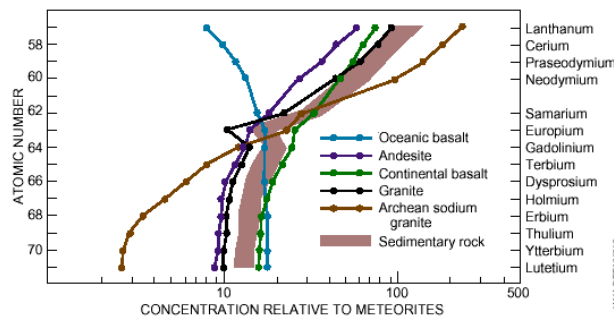
- Slab of lithosphere is subducted, melted & incorporated into asthenosphere
- Convection carries molten material upward where it emerges along a spreading zone as new lithosphere.

## Plate Tectonics & the Rock Cycle



- Subducted sediment melts at a shallower depth where it contributes to magma emitted from an island arc volcano and a mountain chain volcano
- Erosion of volcanic rock provides sediment to complete cycle

From Stanley (1999)



RARE-EARTH ELEMENT abundance patterns provide characteristic chemical markers for the types of rock that have formed the earth's crust. Although igneous rocks (those that solidify from magma) can have highly variable rare-earth element signatures (*dotted lines*), the pattern for most sedimentary rocks falls within a narrow range (*gray band*). That uniformity arises because sediments effectively record the average composition of the upper continental crust.

**Sedimentary  
Rocks  
Represent  
Homogenous  
Mixture of  
Continental  
Crust**

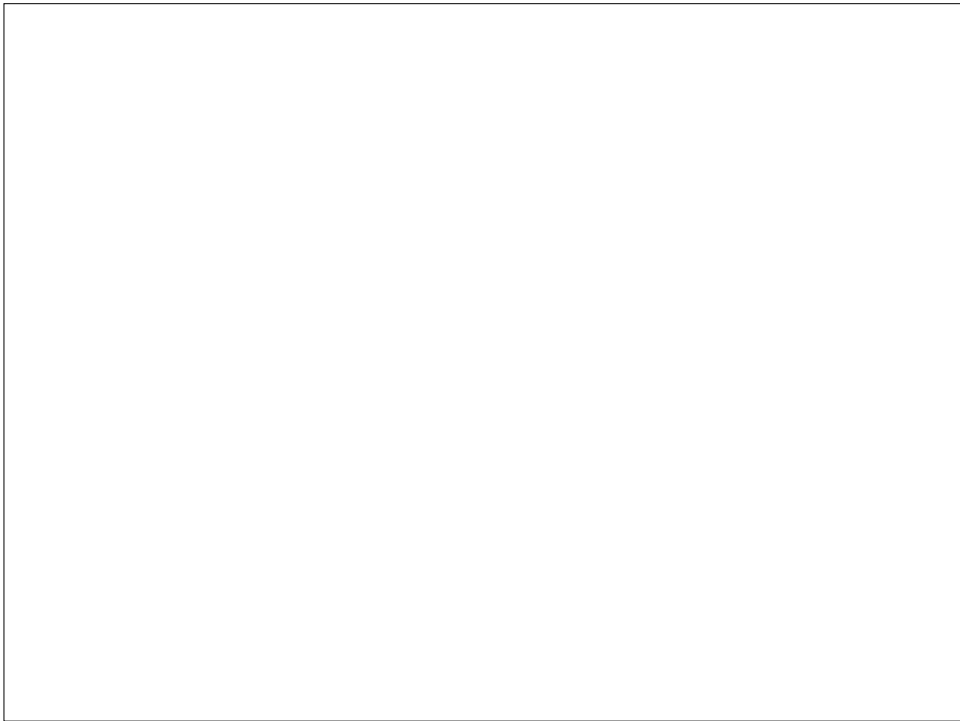
Taylor & McLennan *Sci. Am.* (1996)



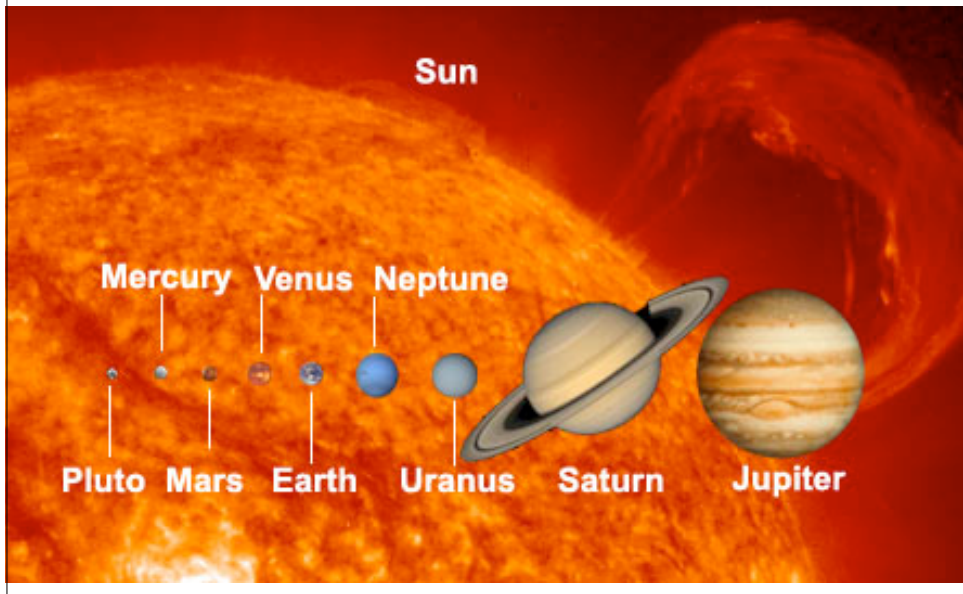
\*\*\* Ended Here - 10/8/08 \*\*\*



Density of lead =  $11.34 \text{ g/cm}^3$   
(at Rm. T)

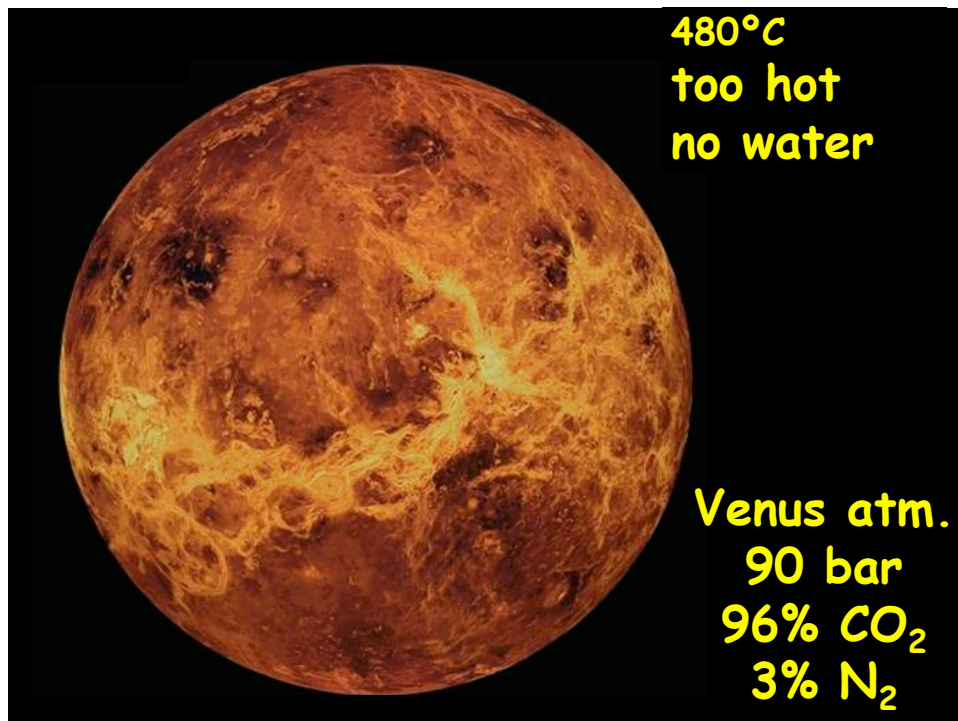
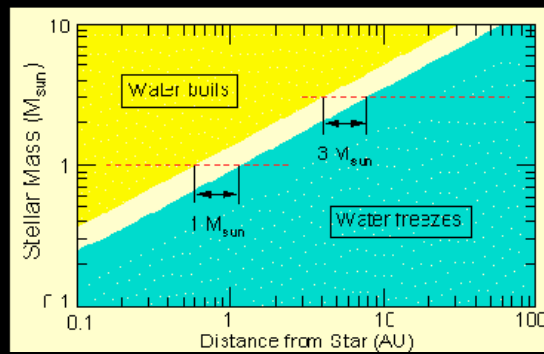


## The Habitable Zone of the Solar System (aka, Top Real Estate Markets outside the Sun)

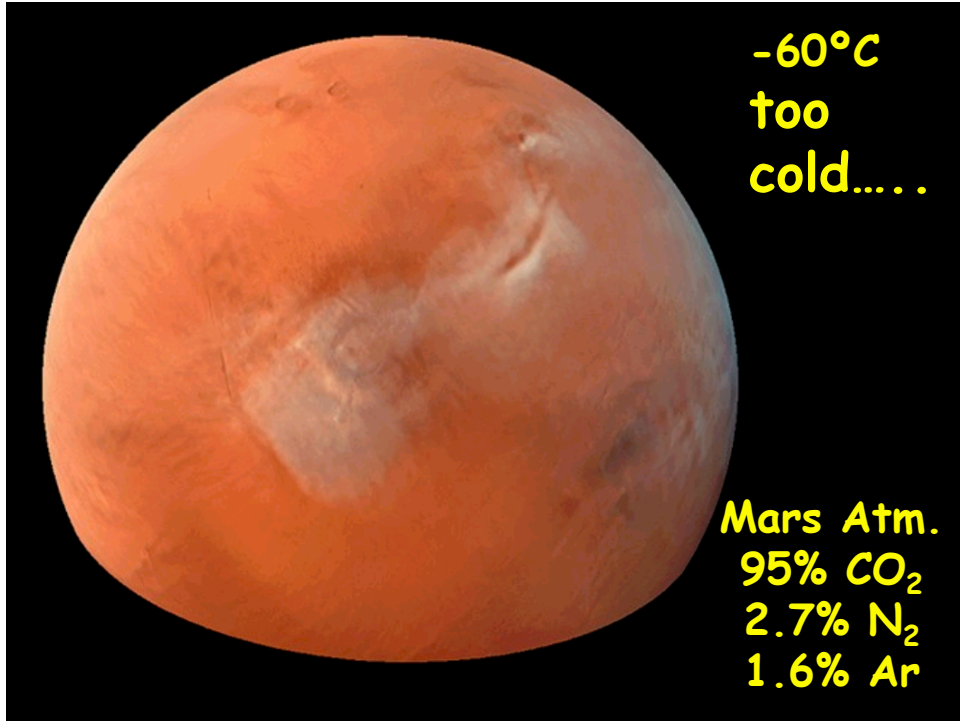


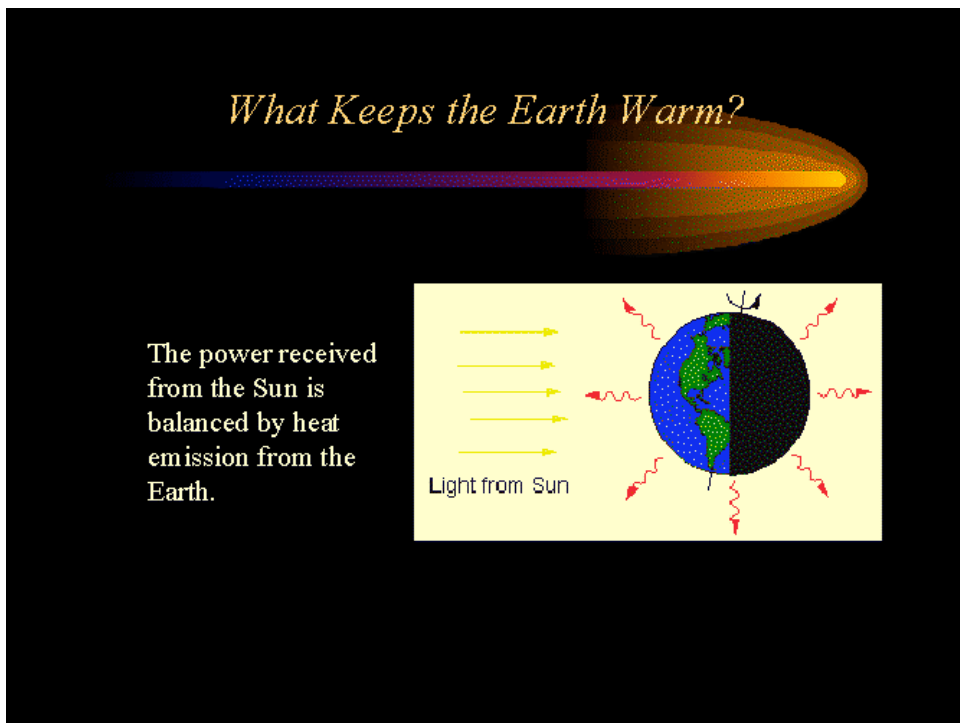
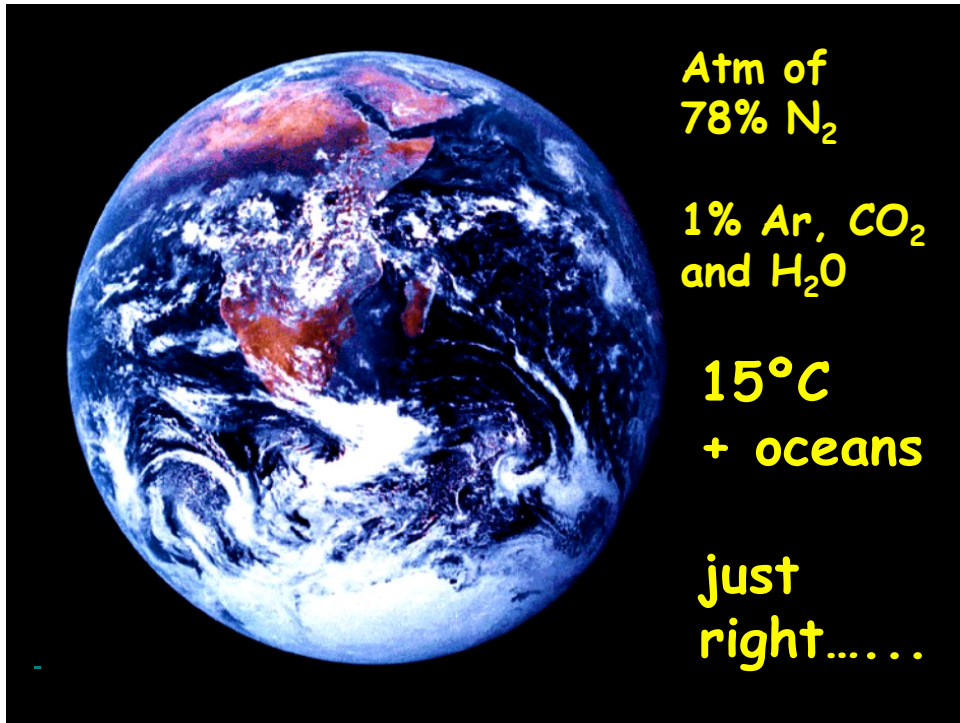
## *Habitable Zone (Ecosphere)*

- Region around a star where planet temperature allows liquid water to exist
- $273 < T_p < 373$  K
- distance of the habitable zone from the star will vary depending on the type of star











## Simple Planetary Energy Balance

$$E_{\text{emitted}} = E_{\text{absorbed}}$$

### ① E<sub>emitted</sub>

- Blackbody w/ effective radiating temperature,  $T_e$

- Stefan-Boltzmann Law

$$E = \sigma T_{\text{eff}}^4 \quad (\sigma = 5.67 \times 10^{-8} \frac{\text{W}}{\text{m}^2 \cdot \text{K}^4})$$

→ Energy emitted per unit area

- For entire surface of Earth

$$E_{\text{emitted}} = 4\pi R_{\text{Earth}}^2 \times \sigma T_{\text{eff}}^4$$

(Multiply by area of sphere)

- Blackbody: emits radiation w/ 100% efficiency @ all  $\lambda$

Adapted from Kump et al. (1999)

### ② Energy Absorbed

$$E_{\text{absorbed}} = E_{\text{intercepted}} - E_{\text{reflected}}$$

Cross section of Earth = area of circle with Earth radius

$$\begin{aligned} &= \pi R_E^2 \times S - \pi R_E^2 \times S \times A \\ &= \pi R_E^2 S (1-A) \end{aligned}$$

$$E_{\text{emitted}} = E_{\text{absorbed}}$$

$$4\pi R_E^2 \times \sigma T_{\text{eff}}^4 = \pi R_E^2 S (1-A)$$

$$\sigma T_{\text{eff}}^4 = \frac{S}{4} (1-A)$$

$$T_{\text{eff}} = \sqrt[4]{\frac{S}{4\sigma} (1-A)}$$

## Energy Balance (cont'd.)

$S$  = solar radiation received at the radius of the planet's orbit around star (so  $S$  is a function of the luminosity of the star and the distance the planet is from the star)

$A$  = albedo; the fraction of solar radiation reflected back to space from clouds, ice, deserts, etc.

Adapted from Kump et al. (1999)

The Surface T ( $T_s$ ) of a Planet can Differ from the Radiating T ( $T_{\text{eff}}$ ) if its Atmosphere Contains Heat-Trapping (“Greenhouse”) Gases

$$T_{\text{eff}} = \sqrt[4]{\frac{S_0}{4\sigma} (1-A)}$$

Today:  $= 255 \text{ K} = -18^\circ \text{C}$

Earth surface Temp =  $15^\circ \text{C}$  ←  $T_s$

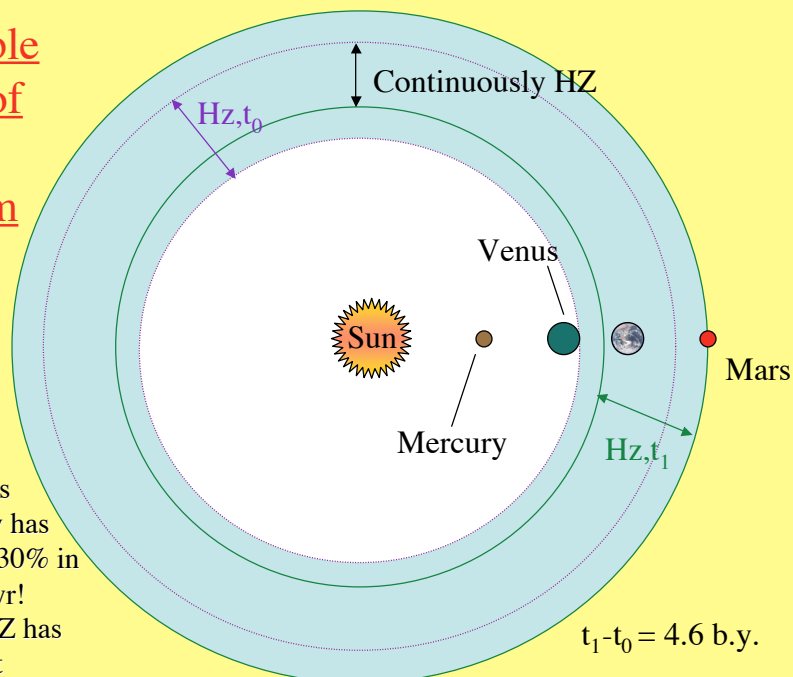
$$T_s - T_{\text{eff}} = \Delta T_g \quad \text{Greenhouse Effect}$$

$$15^\circ - (-18^\circ) = 33^\circ \text{C}$$

Habitability requires:  
 $0^\circ \text{C} < T_s < 100^\circ \text{C}$

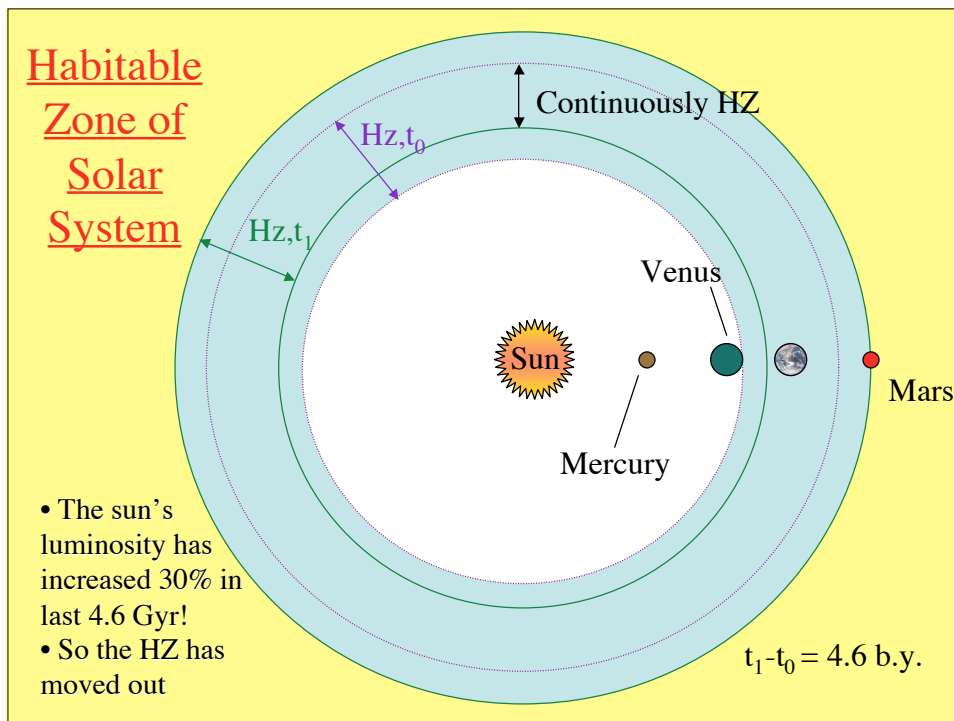
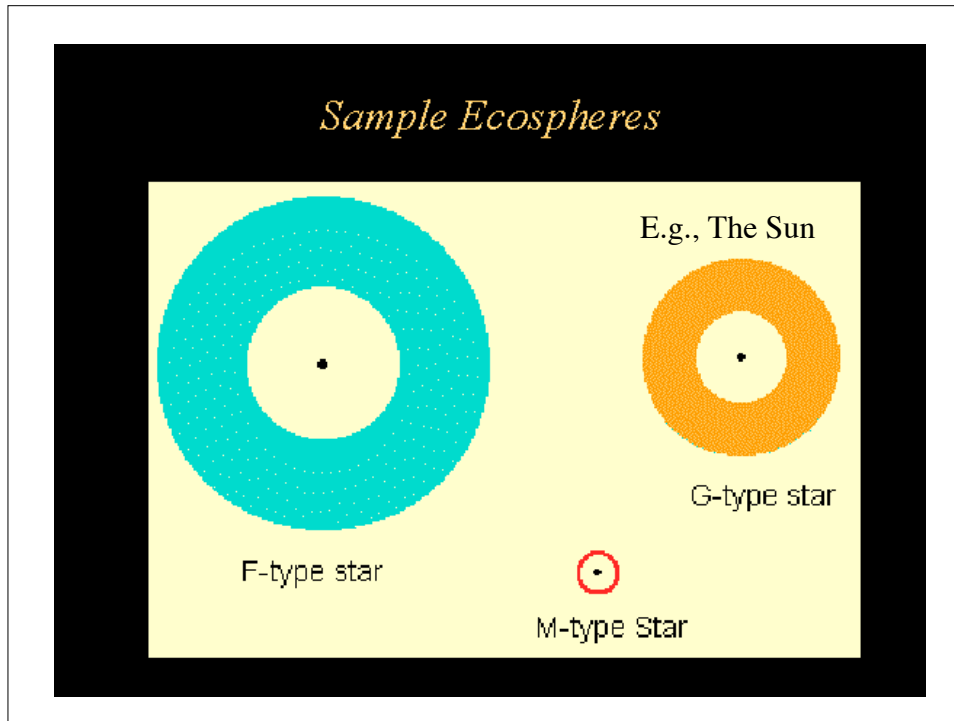
Adapted from Kump et al. (1999)

Habitable Zone of Solar System



- The sun’s luminosity has increased 30% in last 4.6 Gyr!
- So the HZ has moved out

Julian Sachs, Univ. of Washington



## Other Considerations Influencing HZ

*Caveat:* We are relegated to only considering life as we know it & to considering physical conditions similar to Earth

- Greenhouse effect: Increases surface T  
(e.g., Venus, at 0.72 AU, is within HZ, but  $T_s \sim 745$  K!)
- Lifetime of star: larger mass = shorter lifetime  
(must be long enough for evolution)
- UV radiation emission: larger mass = more UV  
(deleterious to life... as we know it)
- Habitable zone moves outward with time  
(star luminosity increases with age)

## **Further Characteristics of the Habitable Zone**

- Liquid water
- Sources of carbon and energy
  - $\text{CO}_2$ , organic matter
  - energy from chemistry of rocks + water
  - energy from the sun
- Mechanisms of renewal and recycling
  - Nutrients limited
  - Space = habitat limited (continents...)
  - **Mechanism = Tectonism. Is it that simple?**

# The Galactic Habitable Zone

“The Galactic habitable zone (GHZ), analogous to the concept of the circumstellar habitable zone, is an annular region lying in the plane of the Galactic disk possessing the heavy elements necessary to form terrestrial planets and a sufficiently clement environment over several billion years to allow the biological evolution of complex multicellular life.”

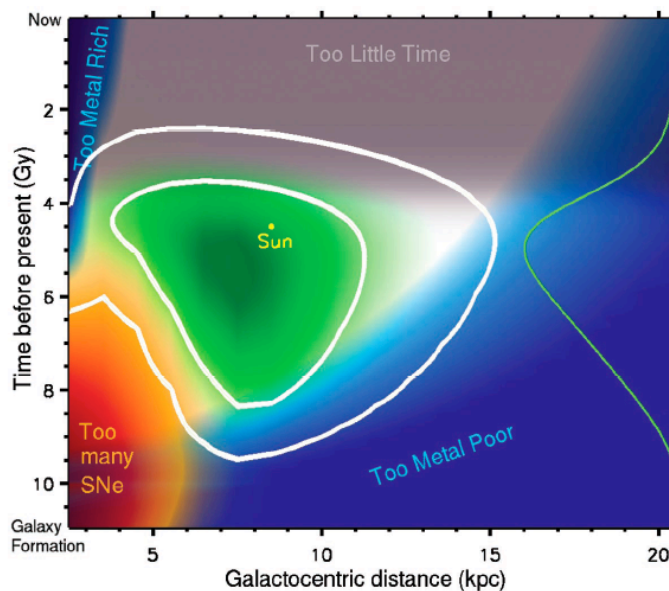
Lineweaver et al. (2004) *Science* Vol. 303: 59-62.

## Habitable Zone of the Milky Way Galaxy

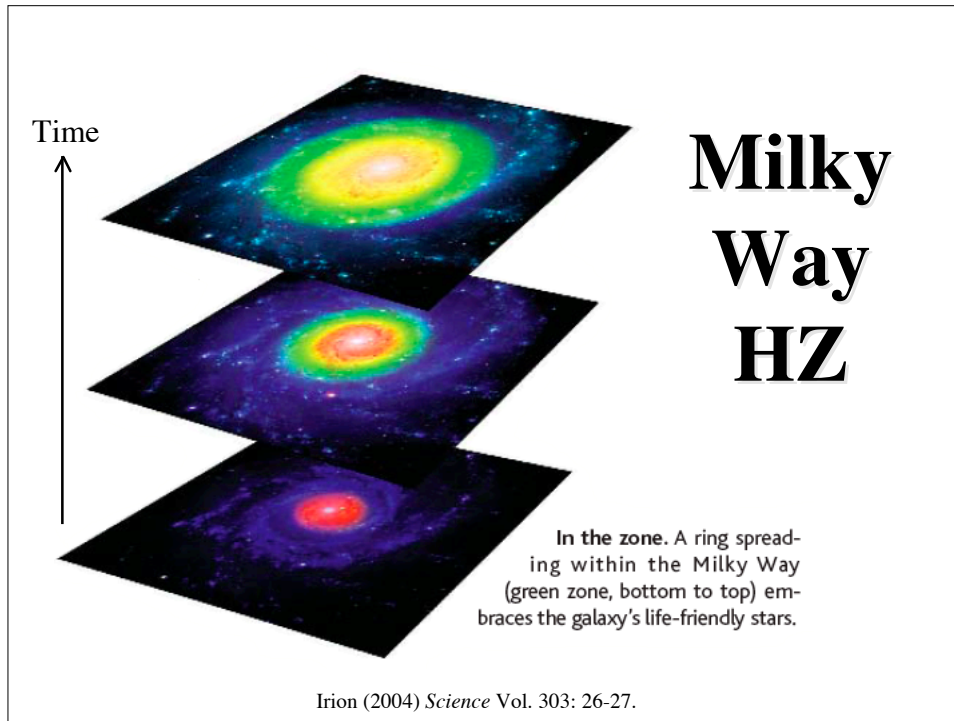
### Requirements

- Metals (Fe)
- Protection from supernovae
- Time for evolution

1 Parsec = 3.26



Lineweaver et al. (2004) *Science* Vol. 303: 59-62.



*Caveat:* If Stars (& therefore solar systems) move around the galaxy then the concept of a Galactic Habitable Zone is less tenable....

Roskar et al. (2008) Riding the Spiral Waves: Implications of Stellar Migration for the Properties of Galactic Disks. *The Astrophysical Journal Letters*, 684(2), L79-L82



## Stars Appear to Migrate Long Distances In Spiral Galaxies like the Milky Way

- The sun might have traveled far from where it formed, contradicting a belief that stars generally remain static
- According to UW astronomers, 9/16/08, using “*N*-body + smooth particle hydrodynamics simulations of disk formation” (100,000 hrs of computer time!)
- May challenge idea of “habitable zones” in galaxies -- where metal abundances, radiation, water, etc. are amenable to life



Roskar et al. (2008) Riding the Spiral Waves: Implications of Stellar Migration for the Properties of Galactic Disks. *The Astrophysical Journal Letters*, 684(2), L79-L82.

Immigrant Sun: Our Star Could Be Far From Where It Started In Milky Way. *Science Daily*, 9/16/08  
Sun might be a long-distance traveler. *UPI.com*, 9/16/08

Simulation of Spiral Galaxy formation: [http://www.astro.washington.edu/roskar/astromy/12M\\_hr\\_rerun\\_angle.mpg](http://www.astro.washington.edu/roskar/astromy/12M_hr_rerun_angle.mpg)

## Views of The Early Earth



- Artist’s view of Hadean Earth, 1952 cover of *Life*
- While such energetic conditions prevailed for ~50 Myr after Earth formed, recent studies indicate a more clement & cool climate within 200–300 Myr after formation

From *Elements Magazine*, August 2006, p. 201.

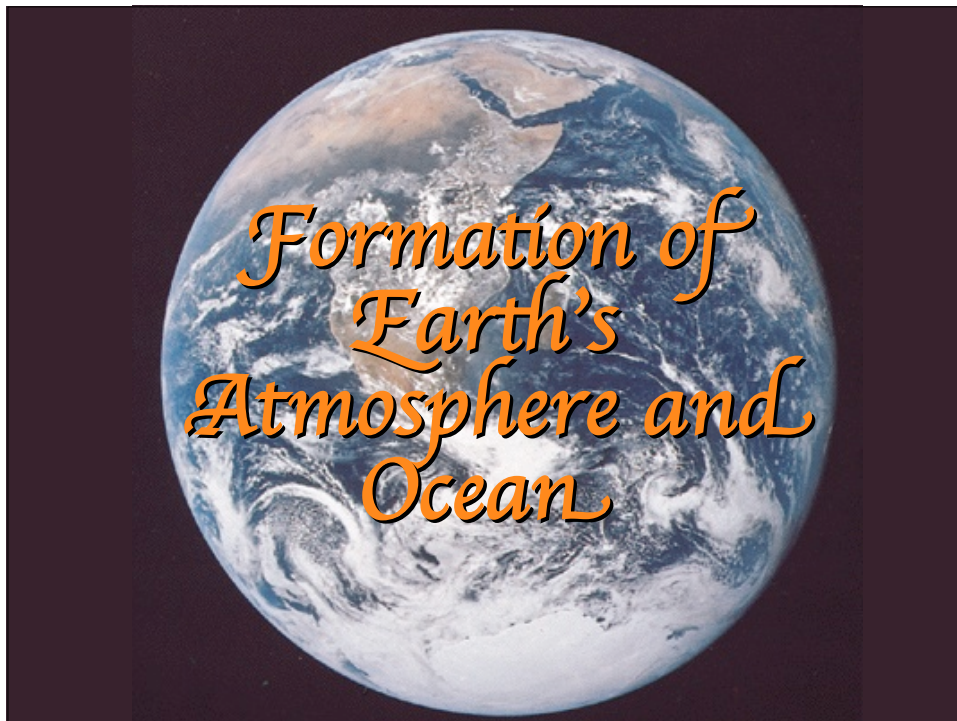
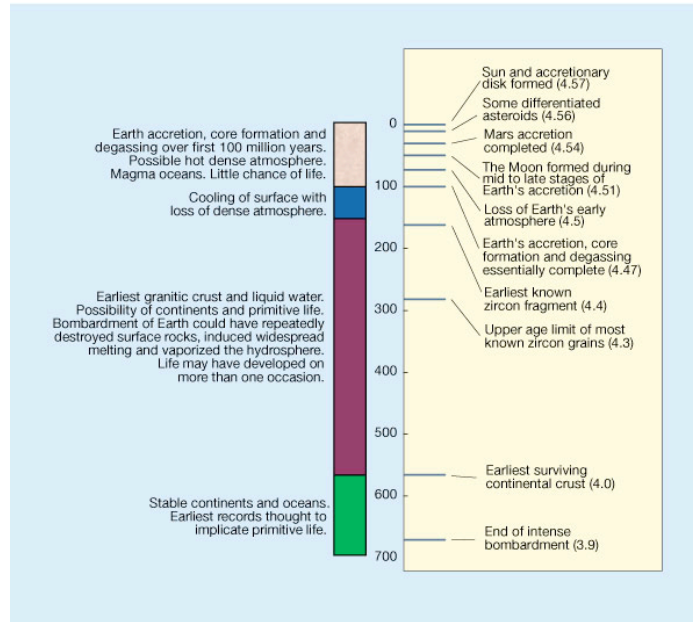


- Artist’s view of cool early Earth ca. 4.2 Ga
- Meteorite impacts still common on Earth & Moon
- Moon in closer orbit & dark maria not yet formed
- Volcanism & other magmatic processes on Earth start to form granitic rocks & proto-continental crust
- Liquid water covered much of Earth

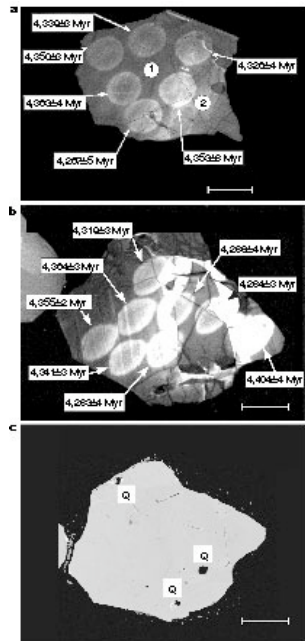
From Valley JW (2005) A cool early Earth? *Scientific American* October 2005: 40-47. IMAGE COURTESY OF DON DIXON, in *Elements* (cover), Aug. 2006.



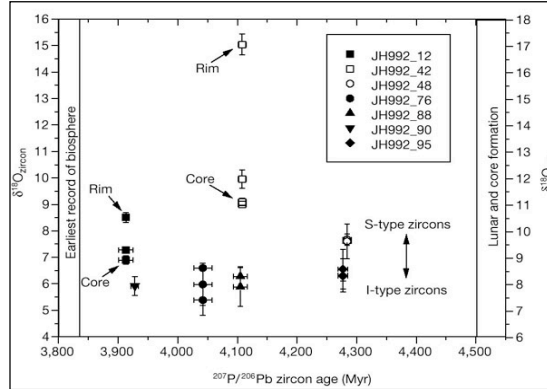
# Early Earth History



## Evidence from Zircons for Liquid Water 4.3 Ga

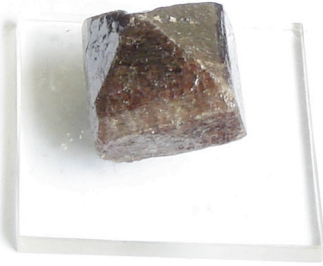


Wilde et al., *Nature* (2001)



- Heavy oxygen isotope ratios ( $^{18}\text{O}/^{16}\text{O}$ ) are produced by low-temperature interactions between rock & liquid water.
- 4.3 Ga zircons have high  $^{18}\text{O}/^{16}\text{O}$ , implying the rocks that were melted to form the magma from which the zircons crystallized included material that had been at the surface in the presence of liquid water.

\*\*\* Ended here - 10/10/08 \*\*\*



Zircon is a mineral belonging to the group of nesosilicates. Its chemical name is zirconium silicate and its corresponding chemical formula is  $ZrSiO_4$ .

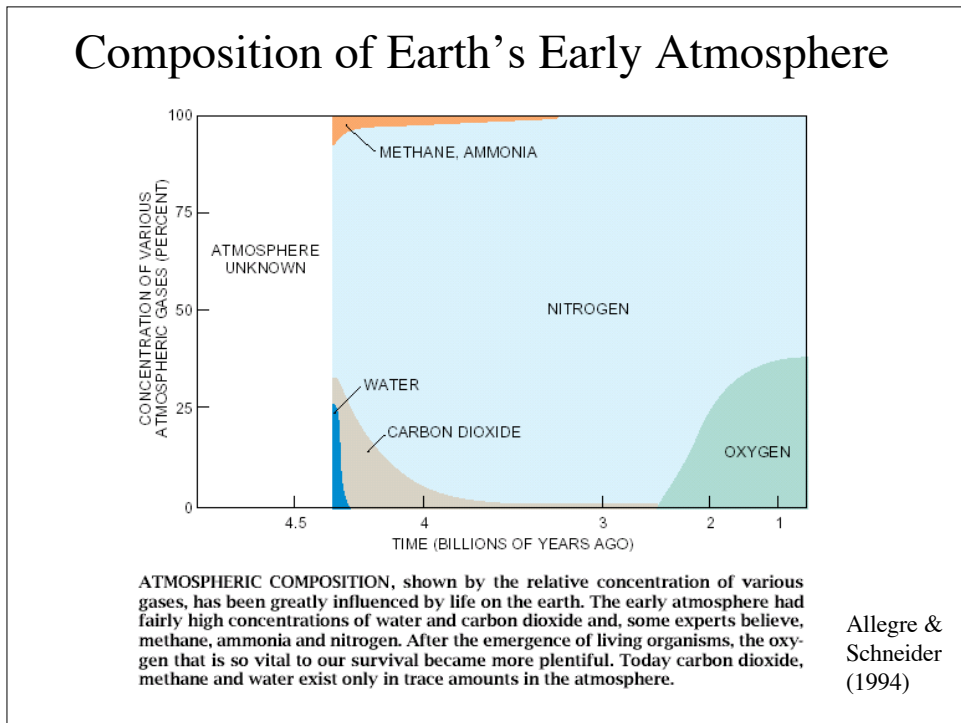
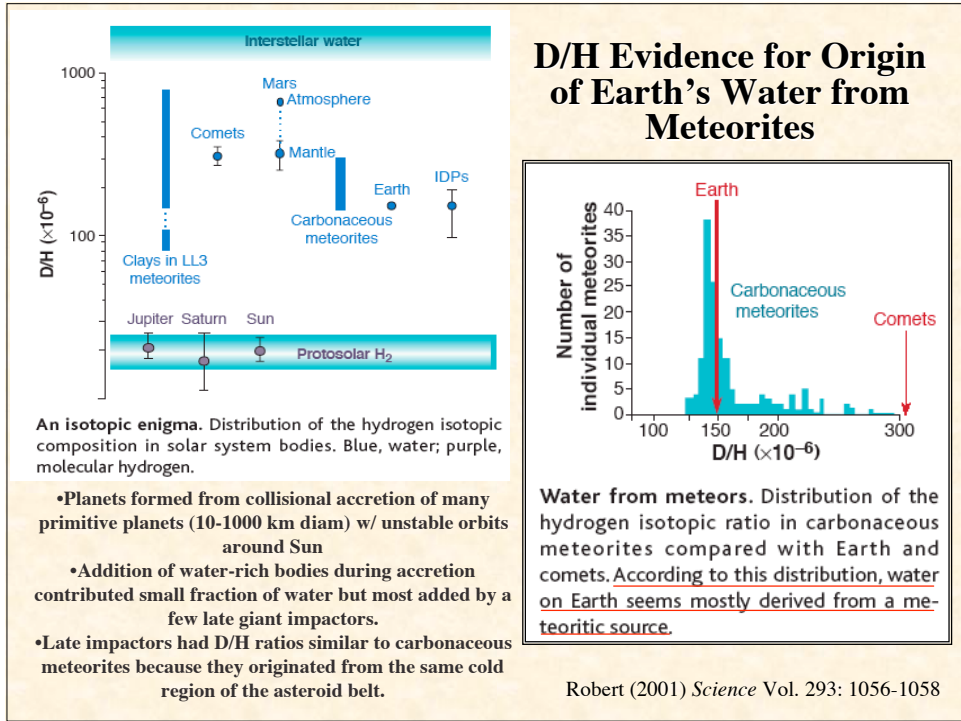
Wikipedia, 10/10/08

## Theories for Origin of Earth's Volatile Components: Atmosphere & Oceans

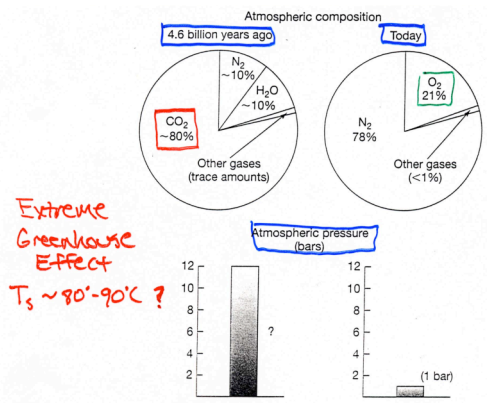
- Arrived with the planetesimals, partly survived the accretion process and outgassed during volcanic activity (Hogbom 1894, Rubey 1951-5). Volcanic gases vary in composition; not primordial and may have been recycled many times. No record of the time and conclusive answers about this scenario (Turekian, 1972; Delsemme, 1997).
- Arrived with comets during the late bombardment - late veneer hypothesis (Delsemme, 1997)
- Arrived with one or more hydrated planetesimals from the outer asteroid belt (Morbidelli, 2001)
- Arrived with meteorites (seems likely....)

## Composition of Comet Halley Volatiles (modeled)

<b>78.5 % H<sub>2</sub>O</b>	<b>2.6% N<sub>2</sub></b>	<b>1.5% C<sub>2</sub>H<sub>4</sub></b>	<b>0.1% H<sub>2</sub>S</b>
<b>4.0% H<sub>2</sub>CO</b>	<b>0.8% NH<sub>3</sub></b>	<b>0.5% CH<sub>4</sub></b>	<b>0.05% S<sub>2</sub></b>
<b>4.5% HCO-OH</b>	<b>1.0% HCN</b>	<b>0.2% C<sub>3</sub>H<sub>2</sub></b>	<b>0.05% CS<sub>2</sub></b>
<b>1.5% CO</b>	<b>0.8% N<sub>2</sub>H<sub>4</sub></b>		
	<b>0.4% C<sub>4</sub>H<sub>4</sub>N<sub>2</sub></b>		
<b>92% with O</b>	<b>5.6% with N</b>	<b>2.6% H/C</b>	<b>0.2% S</b>



# Earth's Early Atmosphere



Extreme Greenhouse Effect  
 $T_s \sim 80^\circ\text{--}90^\circ\text{C}?$

FIGURE 8-7 Possible atmospheric composition and pressure during the heavy bombardment period, compared with today.

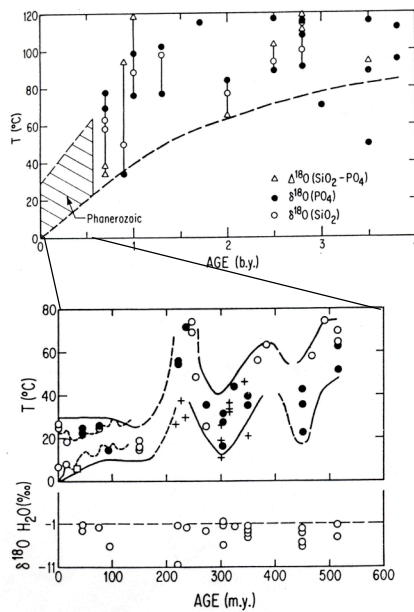
	Prebiotic Atmosphere	Archean Atmosphere	Modern Atmosphere
Surface pressure	1–10 bars	1–2 bars	1 bar
N <sub>2</sub>	10–80%	50–80%	78%
O <sub>2</sub>	about 0	about 0	21%
CO <sub>2</sub>	30–90%	10–20%	0.036%
CH <sub>4</sub>	10–100 ppm	1000–10,000 ppm	1.6 ppm

Constant →

CO 100–1000 ppm 0.1–0.2 ppm  
 H<sub>2</sub> 100–1000 ppm 0.5 ppm

Kump et al. (1999) *The Earth System*, Prentice Hall, New Jersey, 351 p.

## Evidence for High Surface T on early Earth



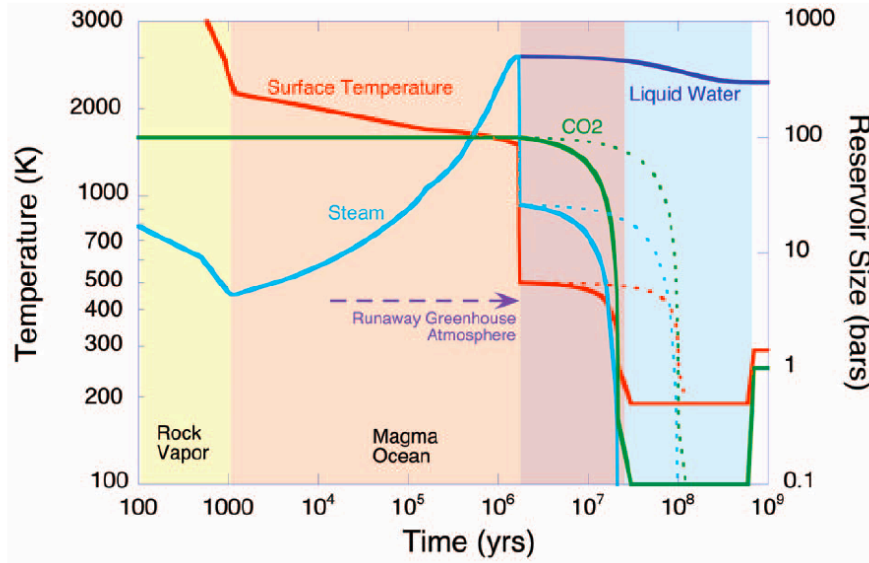
- Oxygen isotope ratios in cherts & phosphates indicate surface temperatures may have exceeded 60°C for most of the period 4–1 Gyr ago



USGS image

Karhu & Epstein (1986) *Geochimica et Cosmochimica Acta* Vol. 50: 1745–1756

## history of temperature, water & CO<sub>2</sub> during the Hadean

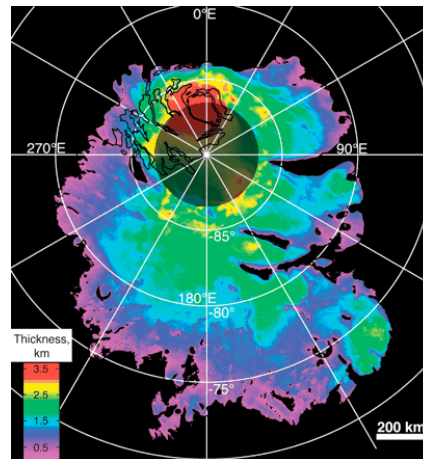
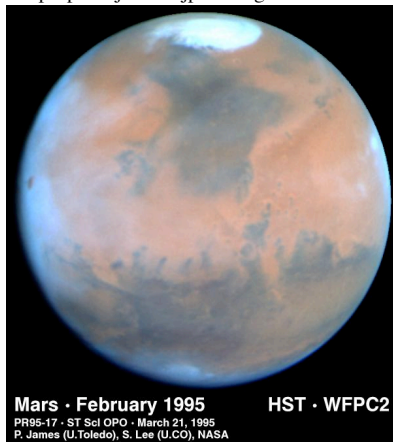


Zahnle (2006) *Elements* Vol. 2(4): 217-222.

## Water Elsewhere in Solar System:

### Water Ice on Mars

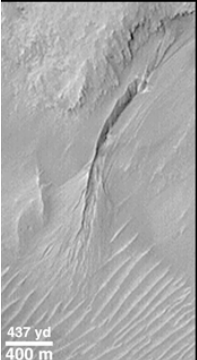
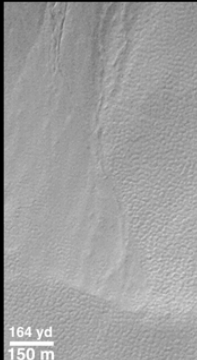
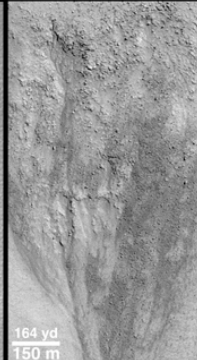
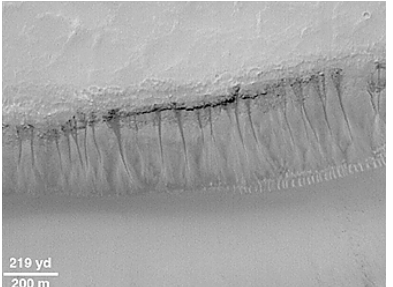

<http://photojournal.jpl.nasa.gov/>



Plaut et al. (2007) *Science*  
 Vol. 316: 92-95.

- South Pole water ice thickness: The total volume is estimated to be  $1.6 \times 10^6$  cubic kilometers, which is equivalent to a global water layer approximately 11 meters thick.

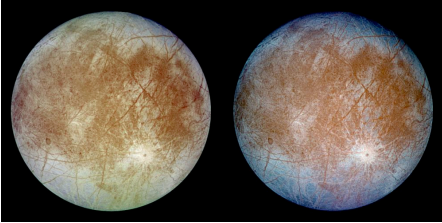


Apron Covering Dunes	Apron on Polygons	Fresh, Dust-free Surfaces
 437 yd 400 m	 164 yd 150 m	 164 yd 150 m
 219 yd 200 m		

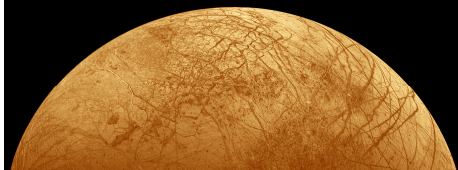
[http://www.msss.com/mars\\_images/moc/june2000/age/index.html](http://www.msss.com/mars_images/moc/june2000/age/index.html)

## Evidence of Recent Water flow on Mars

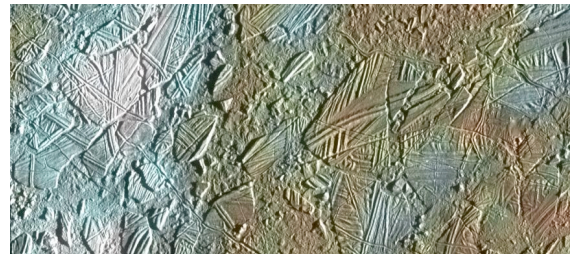
- Martian gullies proposed to have formed by seepage & runoff of liquid water in recent martian times



- Crust composed of water & ice



- Fragmented chunks of water ice on Europa's surface



[http://science.nasa.gov/newhome/headlines/ast09sep99\\_1.htm](http://science.nasa.gov/newhome/headlines/ast09sep99_1.htm); <http://www.solarviews.com/eng/europa.htm>

## Water on Europa

- One of Jupiter's 4 large (Galilean) satellites
- 25% of Earth's radius