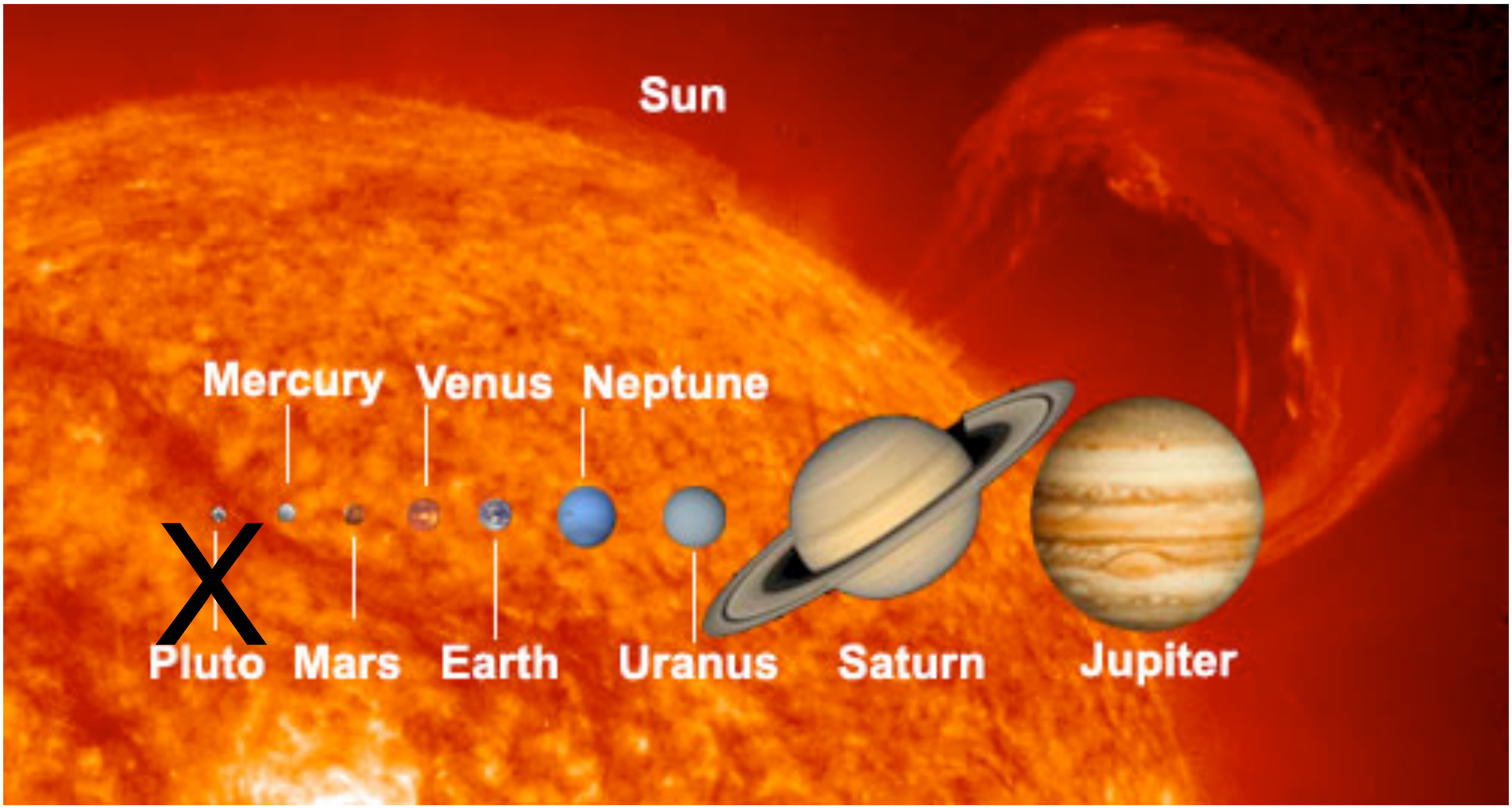
Formation of the Solar System & the Structure of Earth

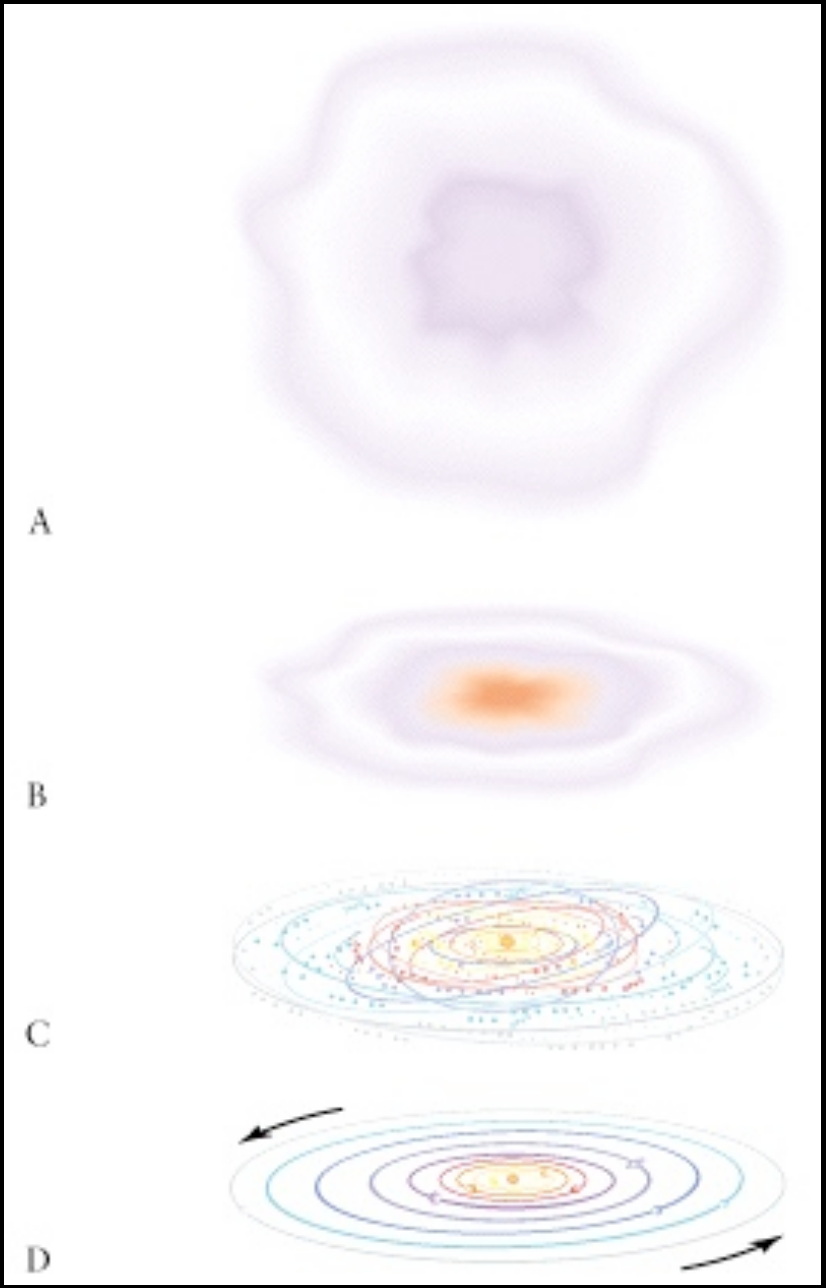
OCEAN 355 Lecture Notes #2



NASA-JPL

Origin of Solar System from solar nebula

- Slowly rotating cloud of gas & dust
- Gravitational contraction
- High P=High T (PV=nRT)
- Rotation rate increases (conserve angular momentum--e.g., skaters)
- Rings of material condense to form planets (Accretion)

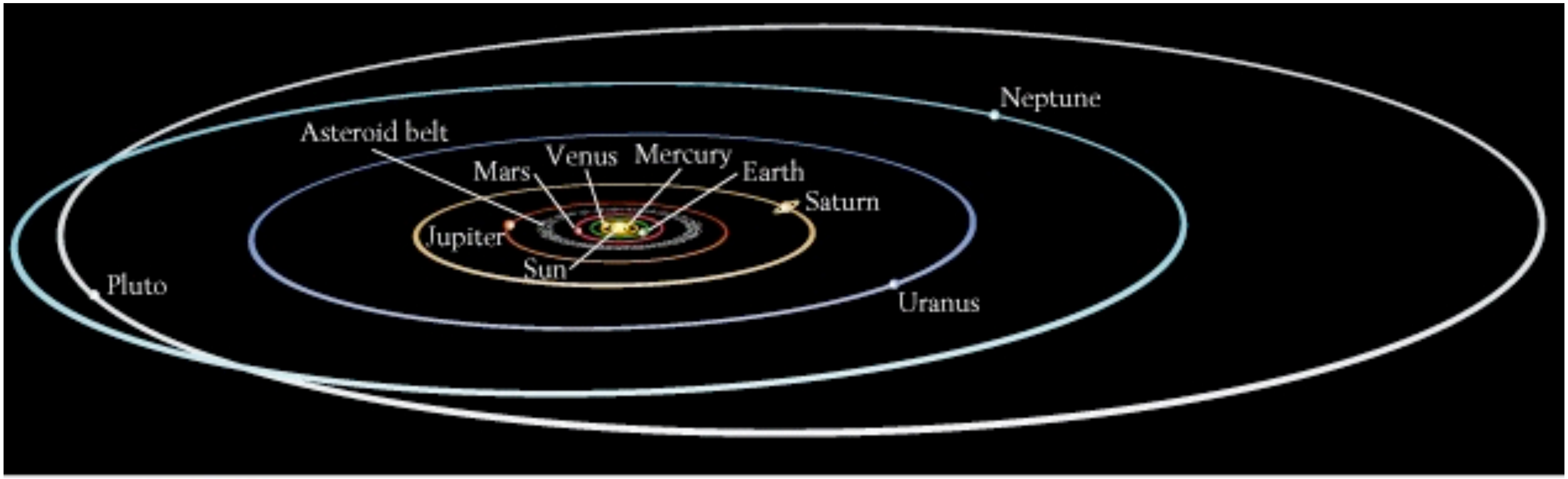


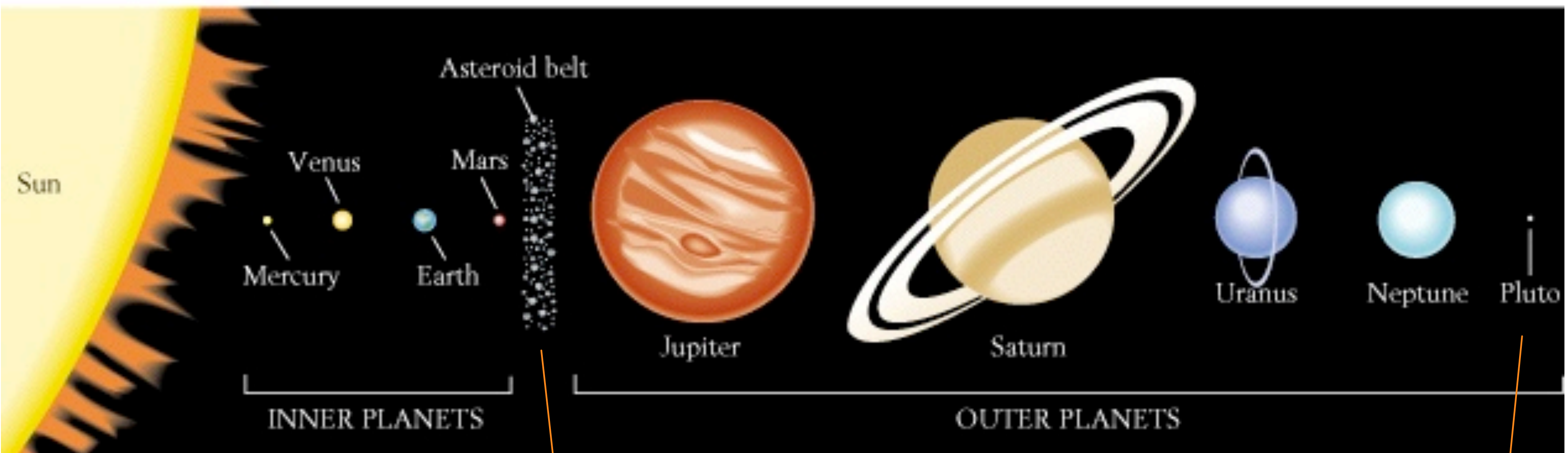
Stanley (1999)

Observational Clues to the Origin of the Planets

- Inner planets are small and dense
- Outer planets are large and have low density
- Satellites of the outer planets are made mostly of ices
- Cratered surfaces are everywhere in the Solar System
- Saturn has such a low density that it can't be solid anywhere

The Solar System





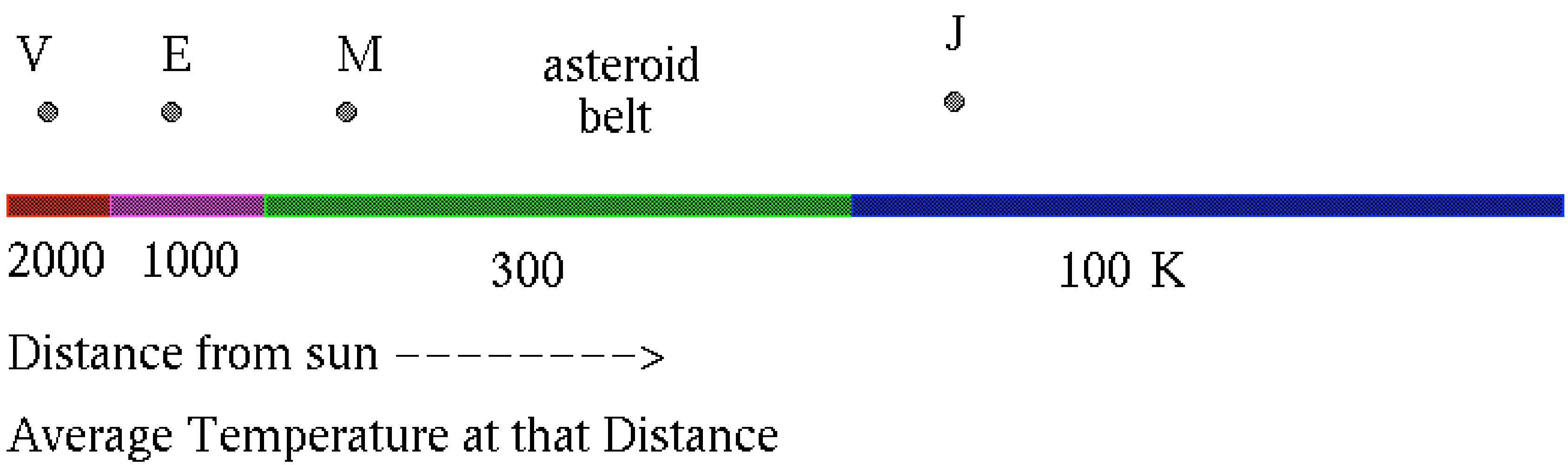
"Failed Planet"

Anomalous Planet

Stanley (1999)

Formation of the Earth by Accretion: 1

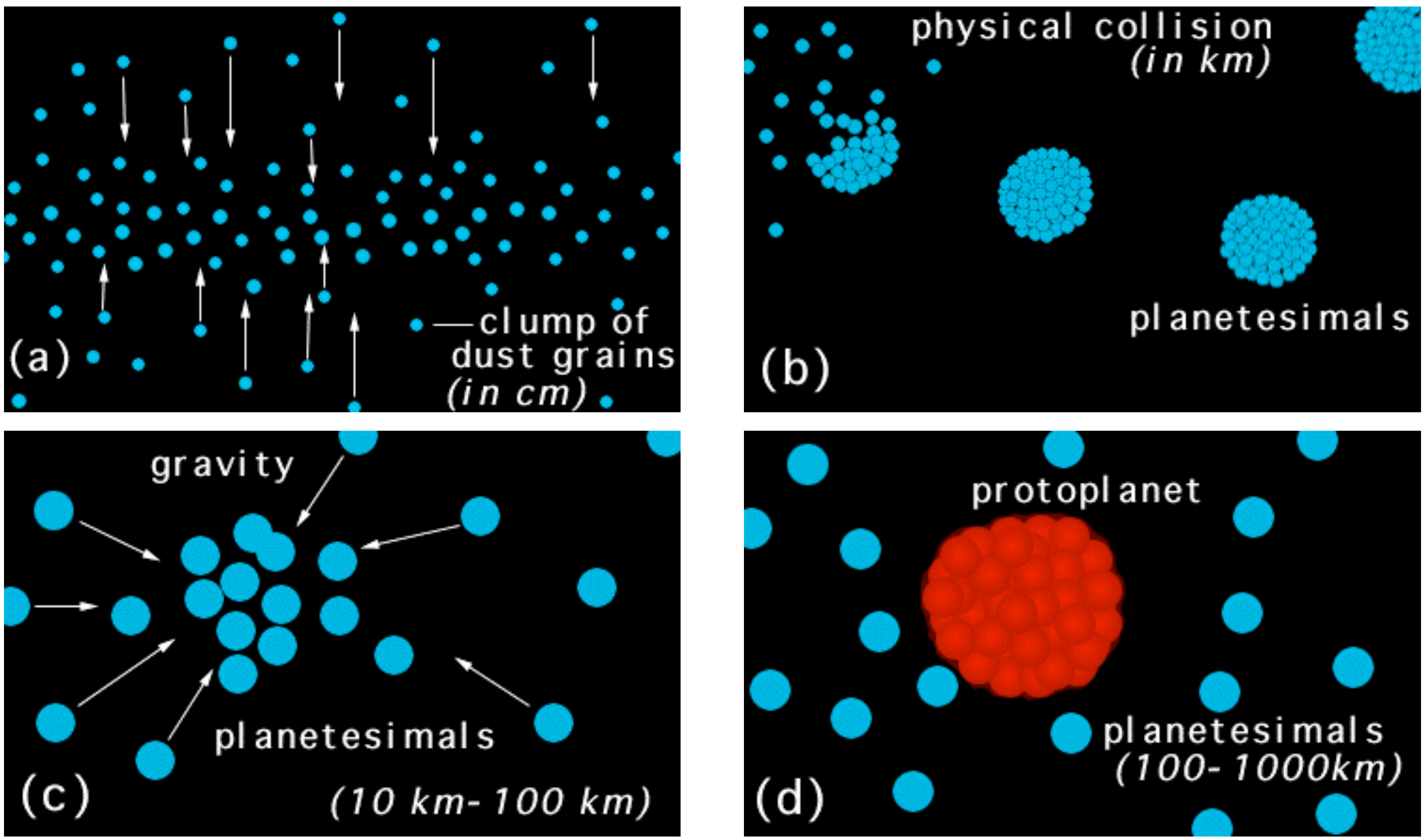
•**Initial solar nebula consisted of cosmic dust & ice with least volatile material condensing closest to the Sun and most volatile material condensing in outer solar system.**



http://zebu.uoregon.edu/ph121/l7.html

Formation of the Earth by Accretion: 2

- Step 1: **accretion of cm sized particles**
- Step 2: **Physical Collision on km scale**
- Step 3: **Gravitational accretion on 10-100 km scale**
- Step 4: **Molten protoplanet from the heat of accretion**



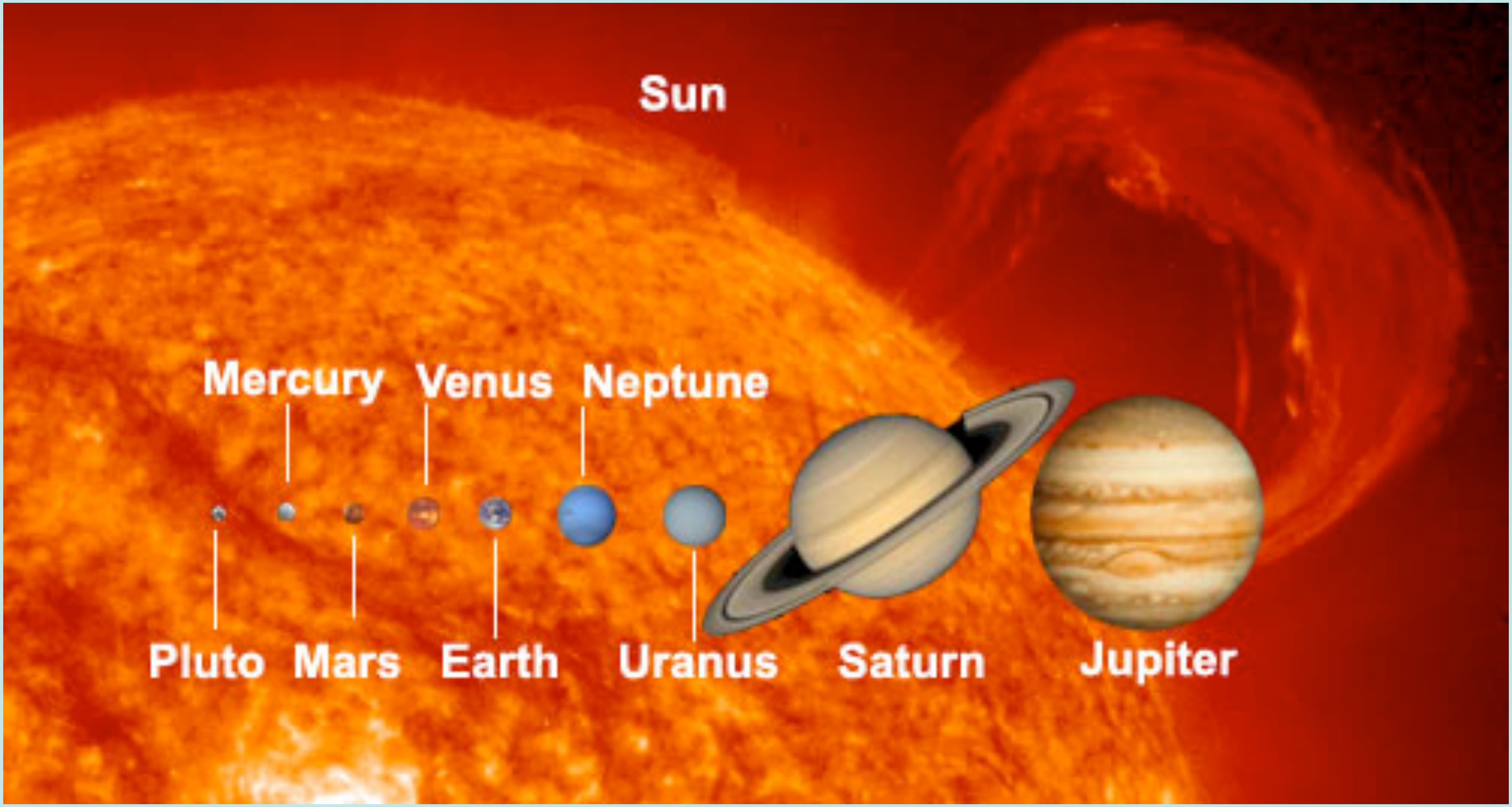
http://zebu.uoregon.edu/ph121/l7.html

Formation of the Earth by Accretion: 3

- Tremendous heat generated in the final accretion process resulted in initially molten objects.
- Any molten object of size greater than about 500 km has sufficient gravity to cause gravitational separation of light and heavy elements thus producing a ***differentiated*** body.
- The accretion process is inefficient, there is lots of left over debris.
- In the inner part of the solar system, leftover rocky debris cratered the surfaces of the newly formed planets *(Heavy Bombardment*, 4.6-3.8 Ga).
- In the outer part of the solar system, the same 4 step process of accretion occurred but it was accretion of ices (cometisemals) instead of grains.

http://zebu.uoregon.edu/ph121/l7.html

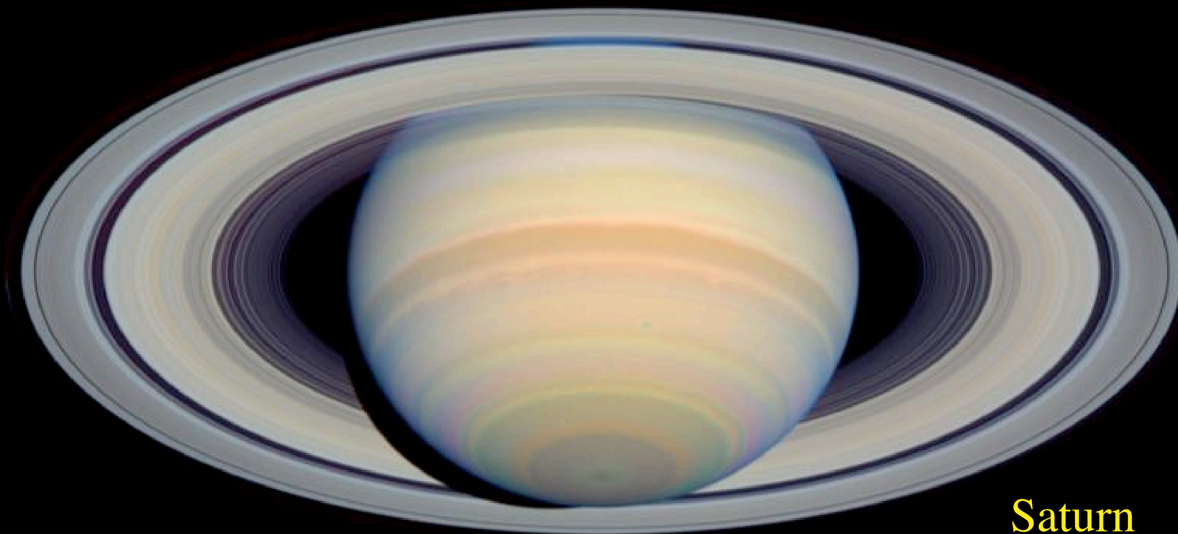
The Sun & Planets to Scale



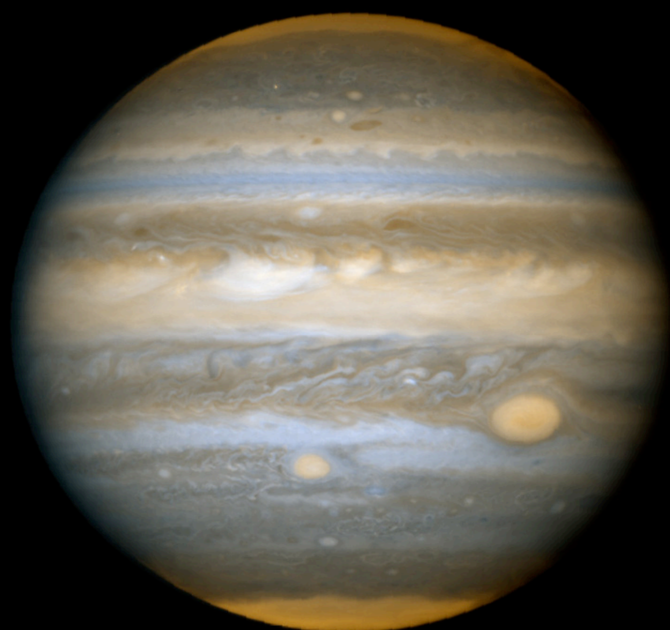
NASA-JPL

More Planetary Eye Candy from Hubble

http://hubblesite.org/gallery/album/solar_system_collection/+1



Saturn



Jupiter

Uranus

Mars

After much outcry, NASA plans a final space shuttle servicing mission to HST in Aug. ‘08 to extend capabilities through 2013

In 2004, NASA administrators canceled a final servicing mission to the telescope that had been scheduled for 2006. The officials based the decision on concern for the safety of astronauts after the loss of the space shuttle Columbia in 2003. Scientists, politicians, and astronomy enthusiasts protested that the decision would bring an early end to the telescope's observations. NASA officials then agreed to study the possibility of sending a robotic craft to perform needed repairs.

STS-125: Final Shuttle Mission to Hubble Space Telescope

Image above: The astronauts selected for the final shuttle mission to perform work on the Hubble Space Telescope pose for a group photo. From left to right are astronauts Megan McArthur, Michael Good, Gregory C. Johnson, Scott Altman, John Grunsfeld, Michael Massimino and Andrew Feustel. Image credit: NASA

HOUSTON - NASA managers officially are targeting August 7, 2008, for the launch of the fifth and final space shuttle servicing mission to the Hubble Space Telescope. During the 11-day flight, Atlantis' seven astronauts will repair and improve the observatory's capabilities through 2013.

Mission planners have been working since last fall, when the flight was announced, to determine the best time in the shuttle manifest to support the needs of Hubble while minimizing the impact to International Space Station assembly. NASA also will support a "launch on need" flight during the Hubble mission. In the unlikely event a rescue flight becomes necessary, shuttle Endeavour currently is planned to lift off from Launch Pad 39-B at NASA's Kennedy Space Center, Fla. However, managers constantly are evaluating the manifest to determine the best mission options.

Shuttle missions beyond the Hubble flight, designated STS-125, still are being assessed. Shuttle and station program officials will continue to consider options for the remainder of the shuttle flights to complete construction of the space station by 2010, when the fleet will be retired. Those target launch dates are subject to change.

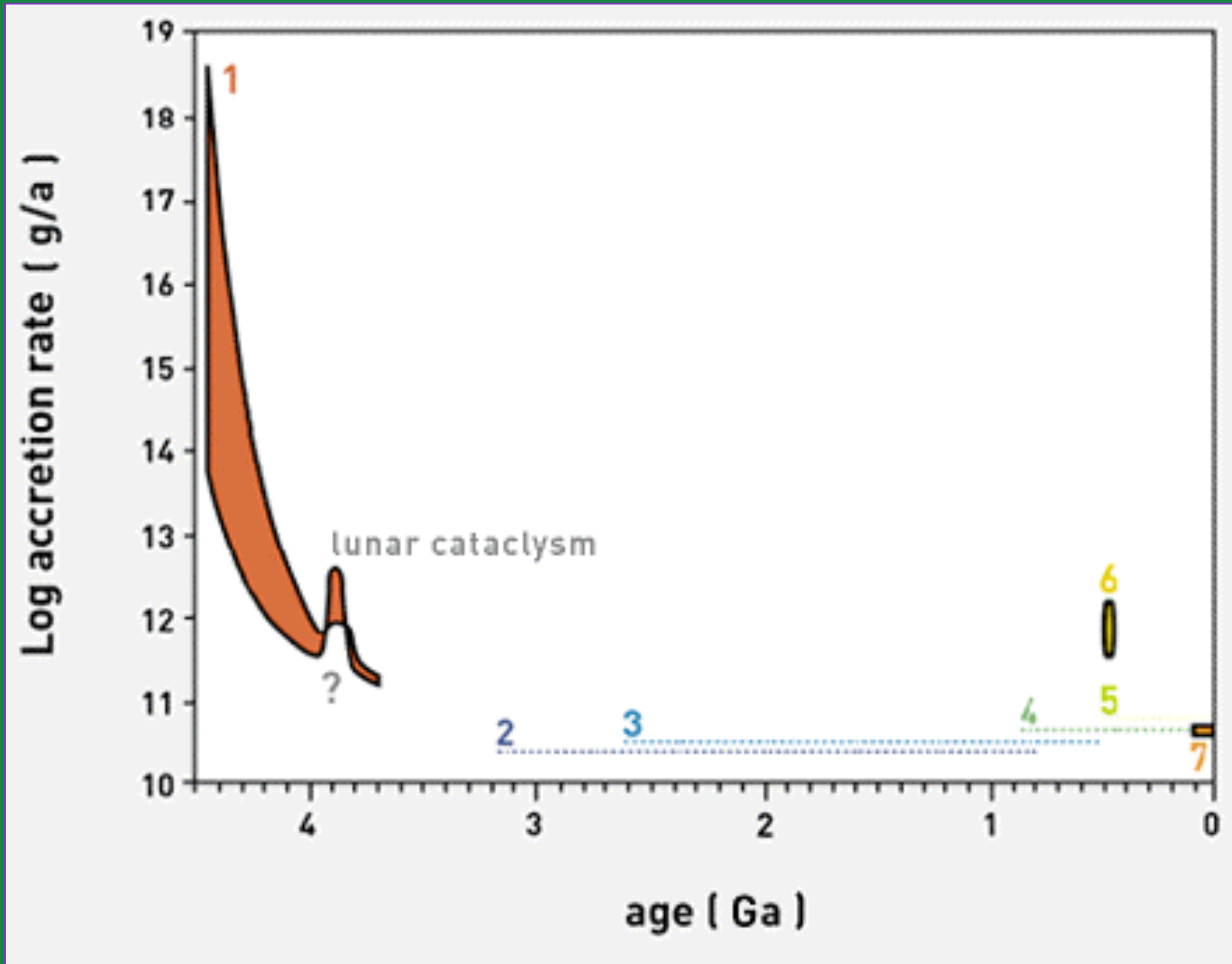
http://www.nasa.gov/mission_pages/hubble/servicing/index.html

http://spaceflight.nasa.gov/gallery/images/shuttle/sts-109/hires/sts109-331-010.jpg

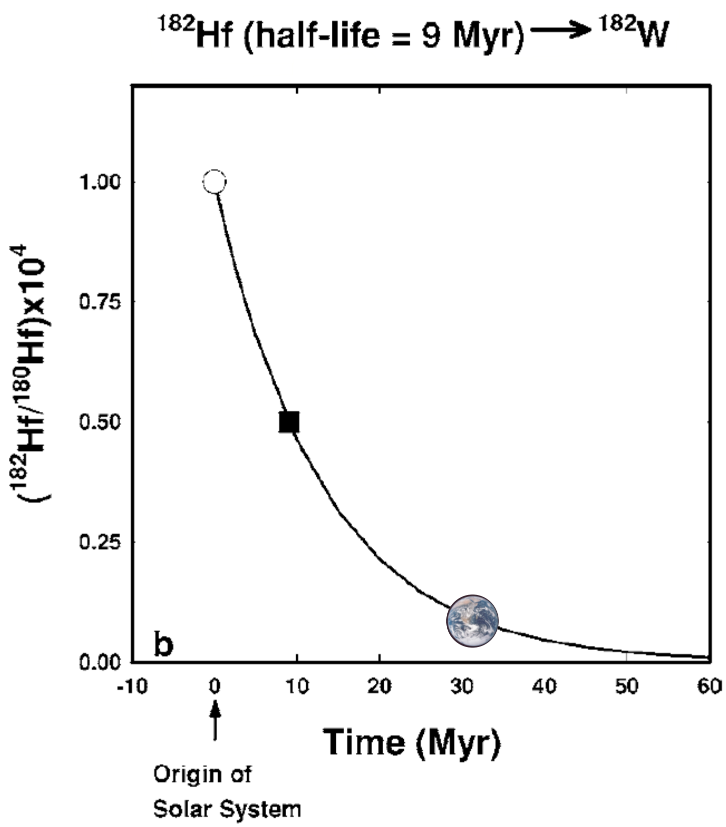
*** Ended Here - 10/3/08 ***

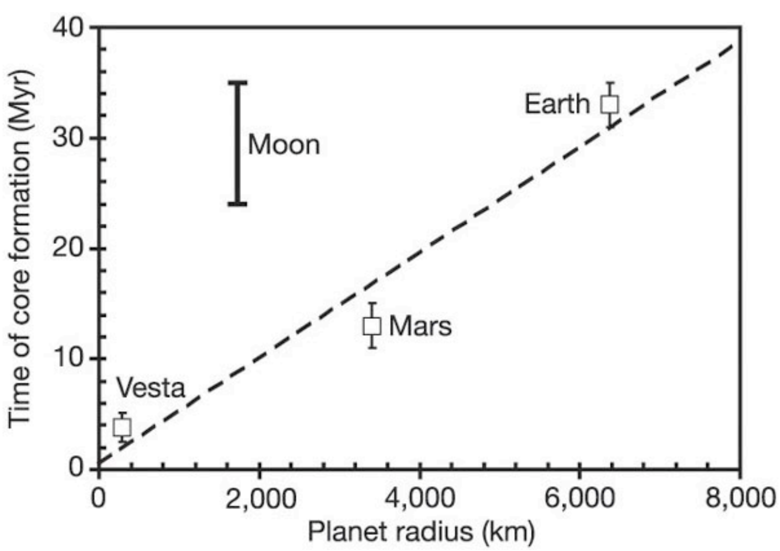
Update Hubble repair mission info

Earth Accretion Rate Through Time



Schmitz et al.. (1997) *Science*, Vol. 278: 88-90, and references therein.
http://www.whoi.edu/science/MCG/pge/project4.html





Terrestrial Planets Formed Rapidly

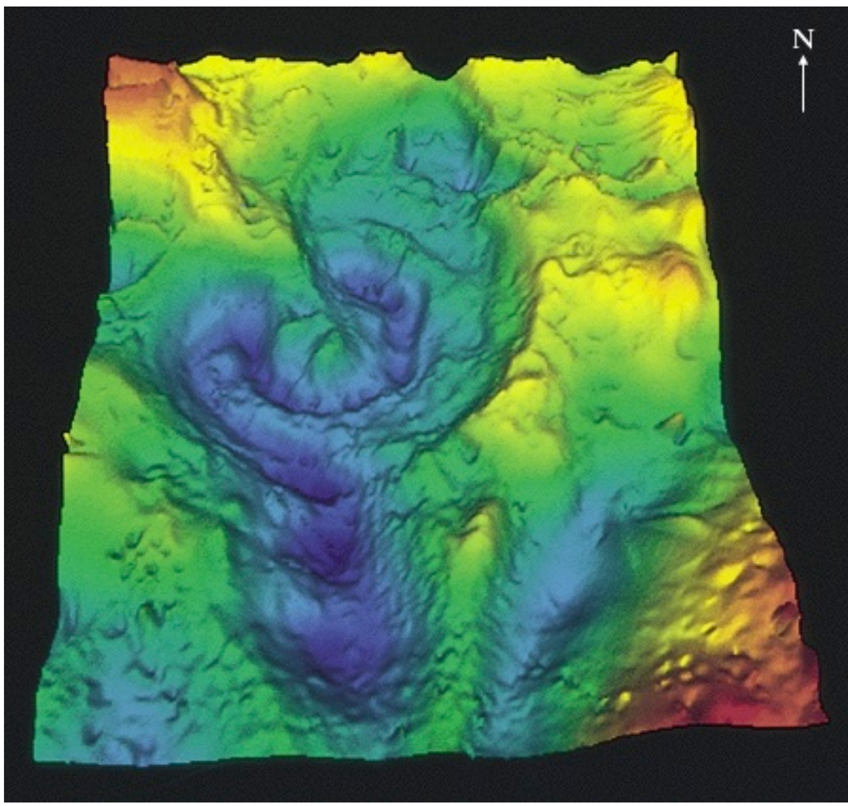
- Hafnium-182 decays to Tungsten-182 with a half-life of 9 Myr
- ~60 Myr after the beginning of the Solar System all ^{182}Hf would have decayed
- **Hf** is retained entirely in the **mantle** during core-mantle segregation while **W** is preferentially partitioned into the **core**.
- If core formation occurred during the lifetime of ^{182}Hf, an excess of ^{182}W should develop in the mantle as a consequence of its enhanced Hf/W ratio
- This excess of ^{182}W has been measured & indicates that core formation on Earth occurred ~30 Myr after the beginning of the Solar System
- So inner planets formed very rapidly (<30 Myr) with smaller bodies forming faster
- Moon must have formed by another mechanism…

Jacobsen (2005) *Ann. Rev. Earth & Planetary Sci.*, 33: 531-570
Kleine et al. (2002) *Nature* 418: 952-955
Yin et al. (2002) *Nature* 418: 949-952

Accretion continues…

Chicxulub Crater, Gulf of Mexico

- 200 km crater
- 10-km impactor
- 65 Myr BP
- Extinction of 75% of all species!



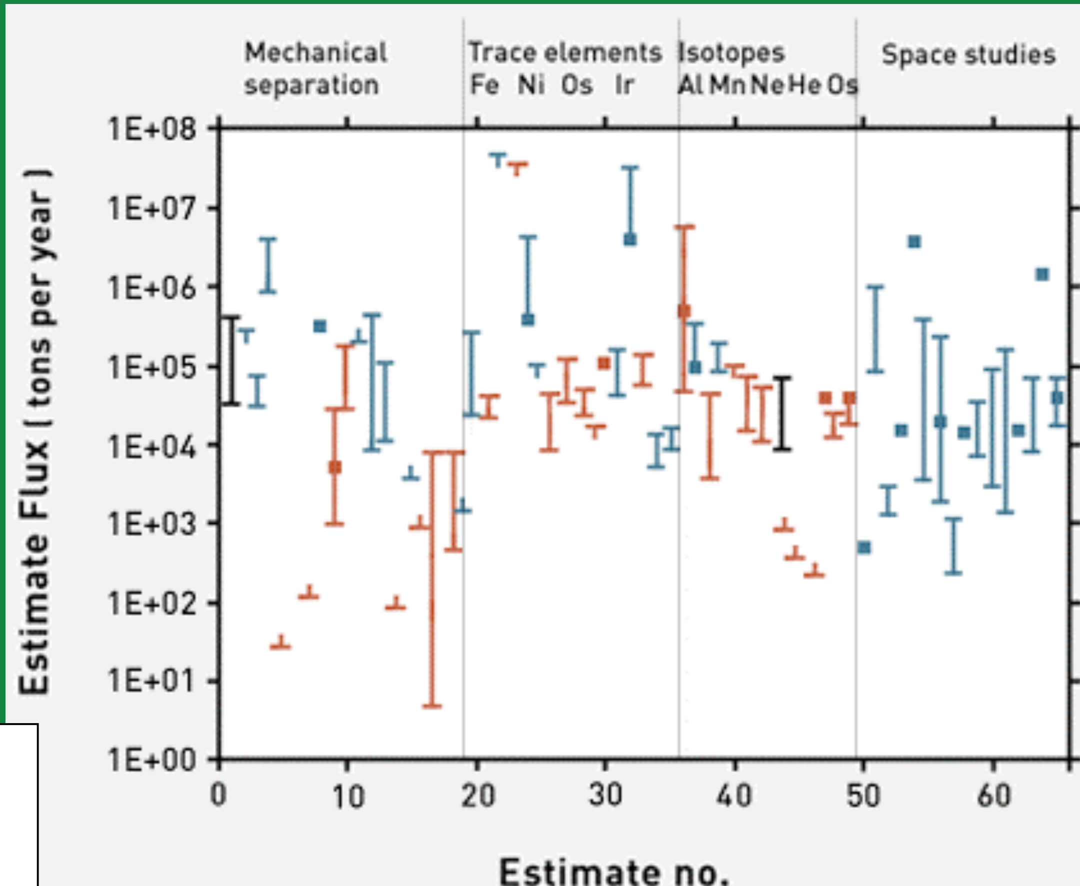
Meteor (Barringer) Crater, Arizona

- 1 km diam. Crater
- 40-m diam Fe-meteorite
- 50 kyr BP
- 300,000 Mton
- 15 km/s

http://www.gi.alaska.edu/remsense/features/impactcrater/imagexplain.htm

Interplanetary Dust Accumulation

http://presolar.wustl.edu/work/idp.html



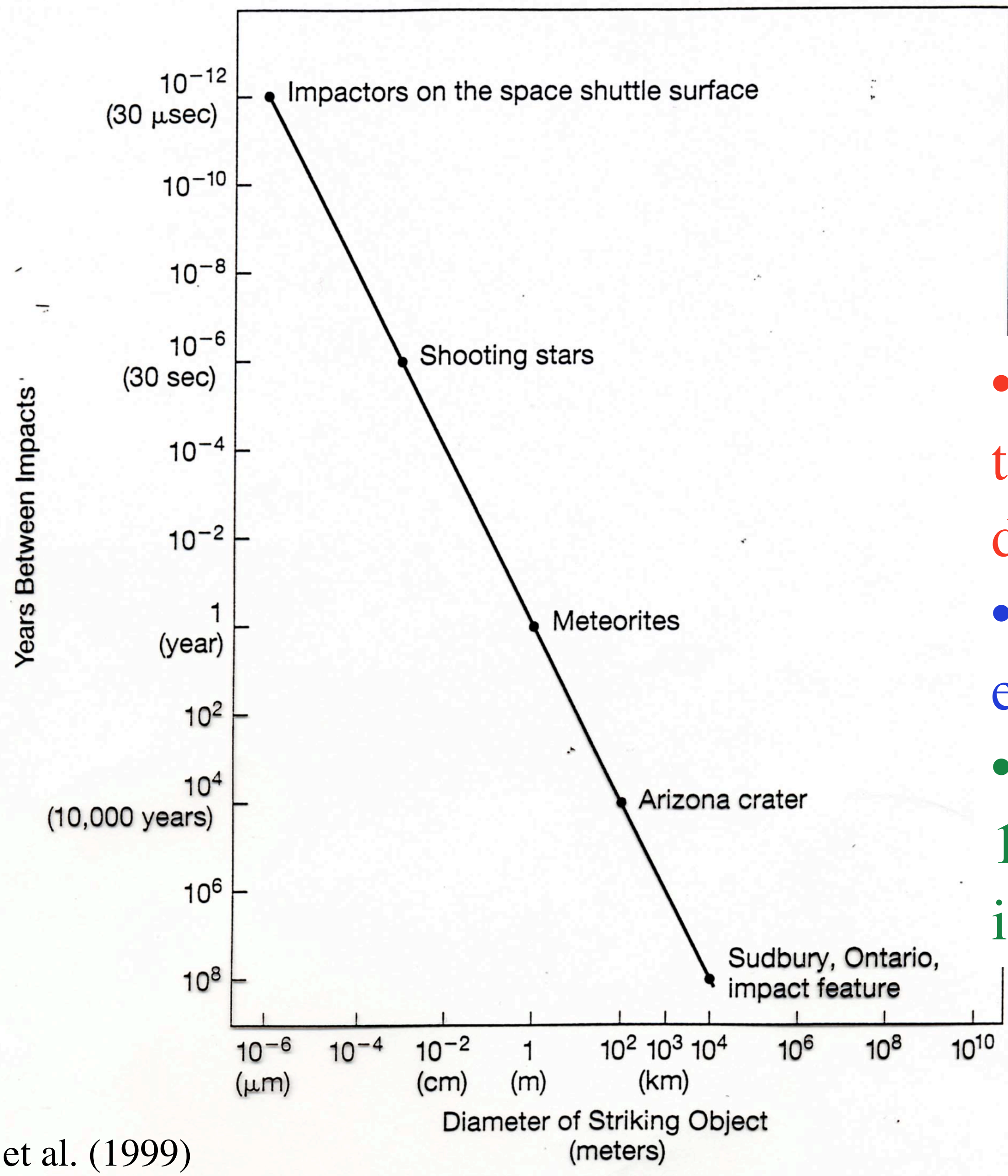
http://www.whoi.edu/science/MCG/pge/project4.html

40±20 $x10^4$ metric tons/ yr ($40\ x10^{10}$ g) interplanetary dust accretes every yr!

= 78.4 g/km^2/yr

= 524.1 Ford Explorers/day

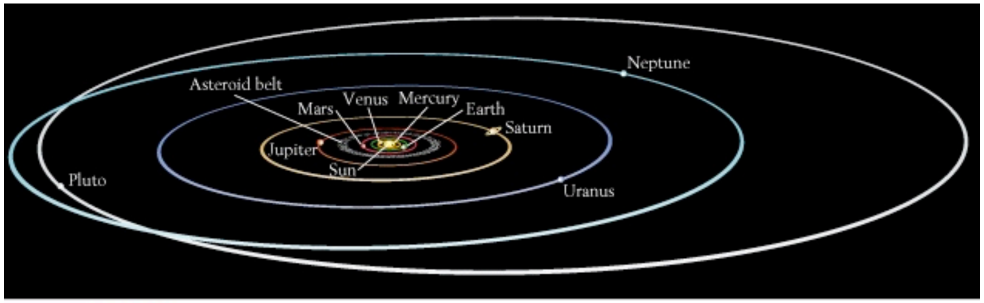
Size & Frequency of Impacts

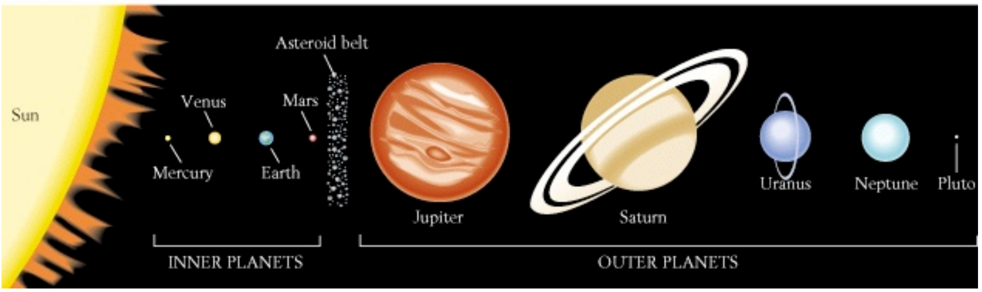


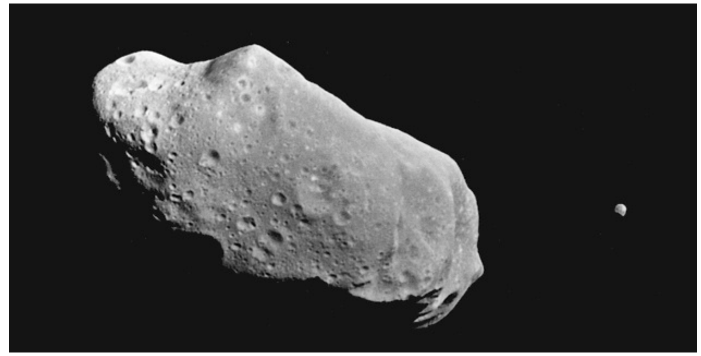
Kump et al. (1999)

- 40±20 x 10^4 metric tons/ yr interplanetary dust accretes every yr!
- 100 m object impacts every 10 kyr
- 10 km object every 100 Myr (e.g., K/T impact)

The Asteroid Belt



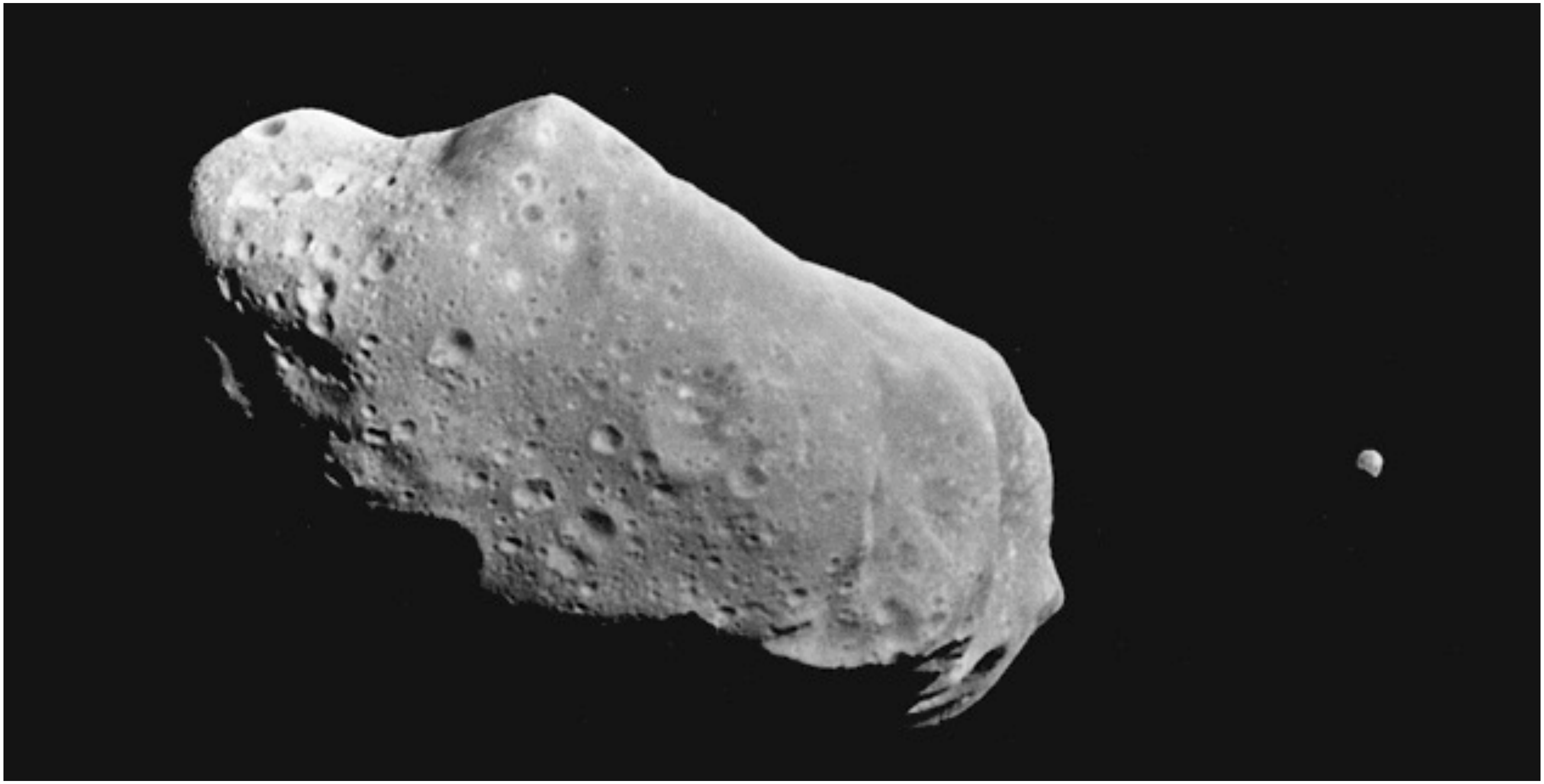




- A relic of the accretion process. A failed planet.
- Gravitational influence of Jupiter accelerates material in that location to high velocity.
- High-velocity collisions between chunks of rock shatter them.
- The sizes of the largest asteroids are decreasing with time.

Total mass (Earth = 1)	0.001
Number of objects > 1 km	~100,000
Number of objects > 250 km	~12
Distance from Sun	2-4 AU
Width of asteroid belt (million km)	180

Asteroid 243 IDA

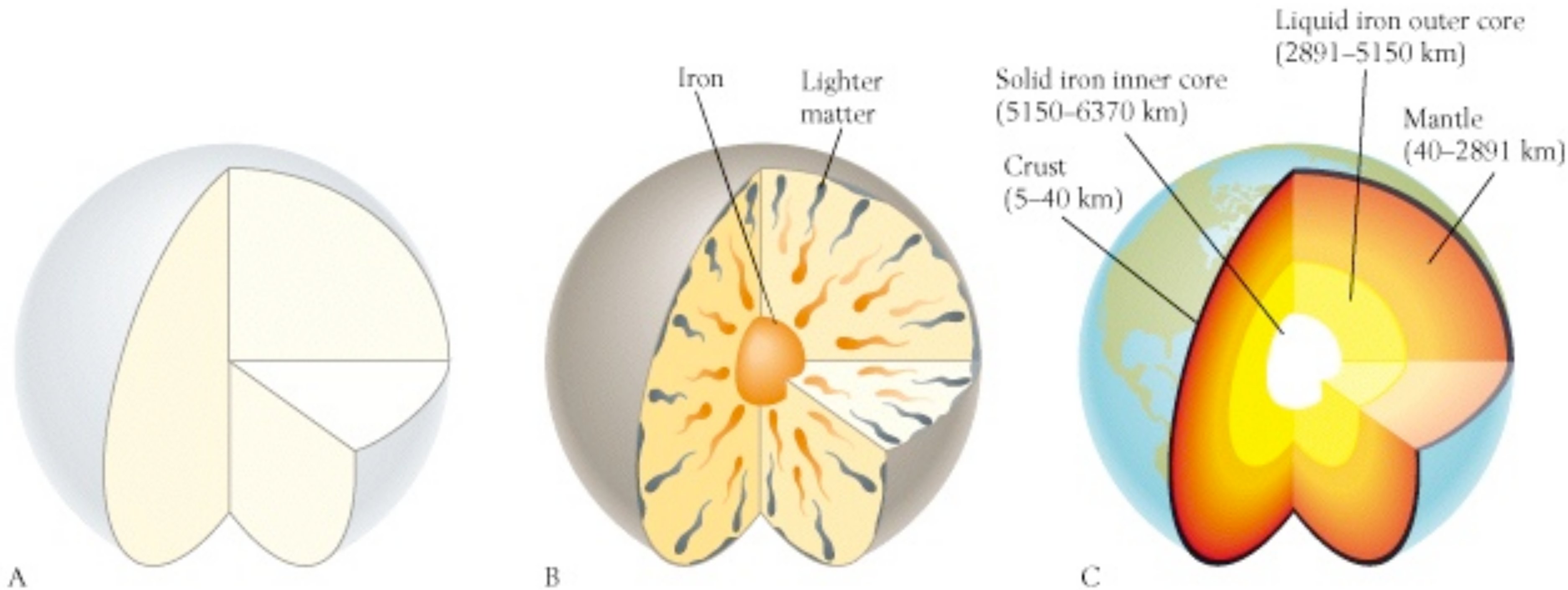


- Meteorite = asteroid that has landed on earth
- All chondrites (meteorites) date to ~4.5 B.y.
- Cratering indicates early origin

Differentiation of the Earth: 1

- VM Goldschmidt (1922) published landmark paper "Differentiation of the Earth":
  1. Earth has a chondritic (meteoritic) elemental composition.
  2. Surface rocks are not chemically representative of solar abundances, therefore must be differentiated.
- Proto-planet differentiated early into a dense iron-rich core surrounded by a metal sulfide-rich shell above which floated a low-density silicate-rich magma ocean.
- Cooling of the magma caused segregation of dense silicate minerals (pyroxenes & olivines) from less dense minerals (feldspars & quartz) which floated to surface to form crust.
- In molten phase, elements segregate according to affinities for: Fe = siderophile, sulfide = chalcophile & silicate = lithophile.

Differentiation of Earth: 2



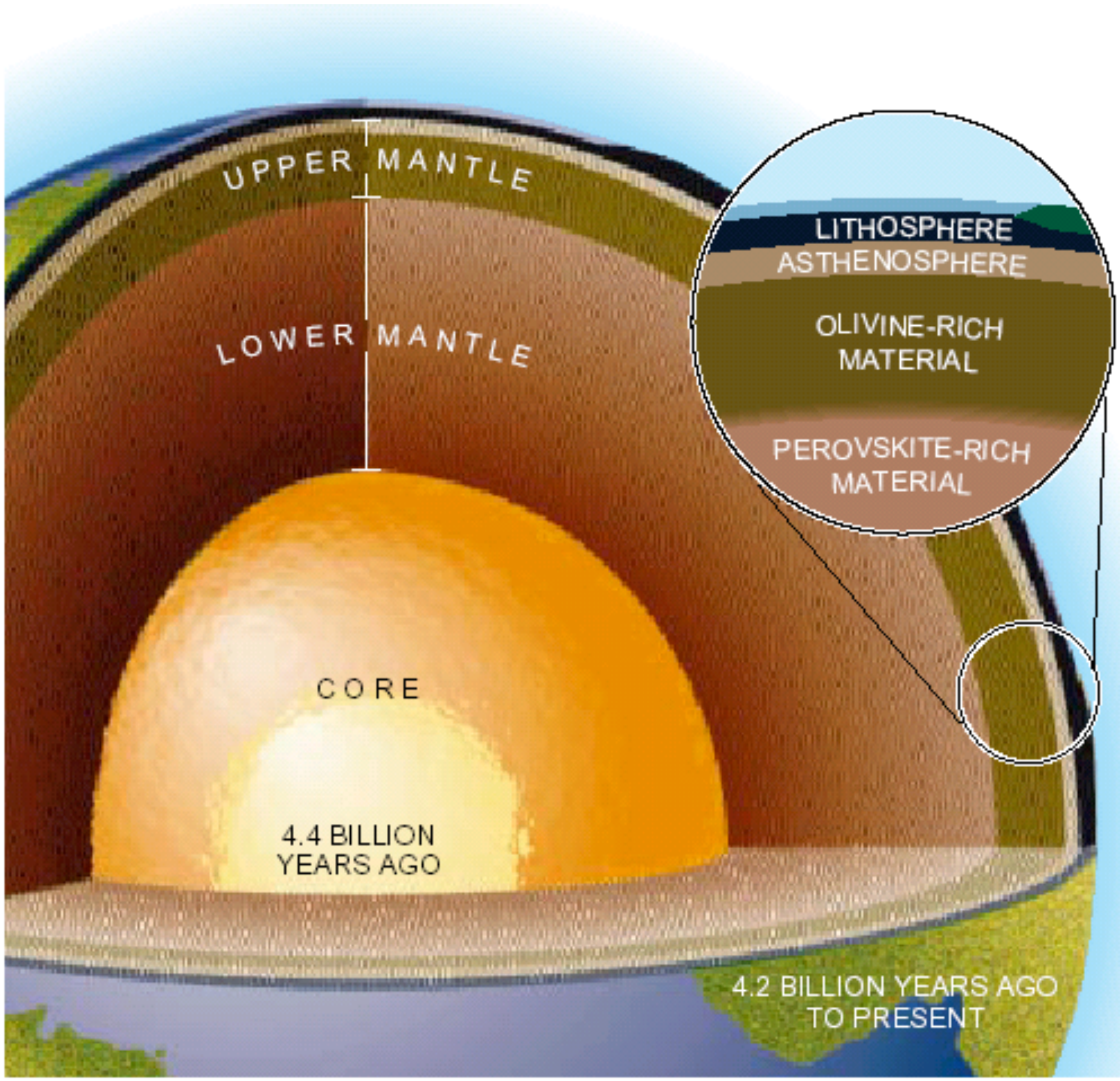
Undifferentiated protoplanet (chondritic composition)

Early differentiation; magma ocean

Fully differentiated Earth

- Driven by density differences
- Occurred in ~30 Myr ($^{182}Hf/^{180}Hf$, c.f., Jacobsen (2005) *Ann. Rev. Earth & Planetary Sci.*, 33: 531-570)

Differentiation of the Earth: 3



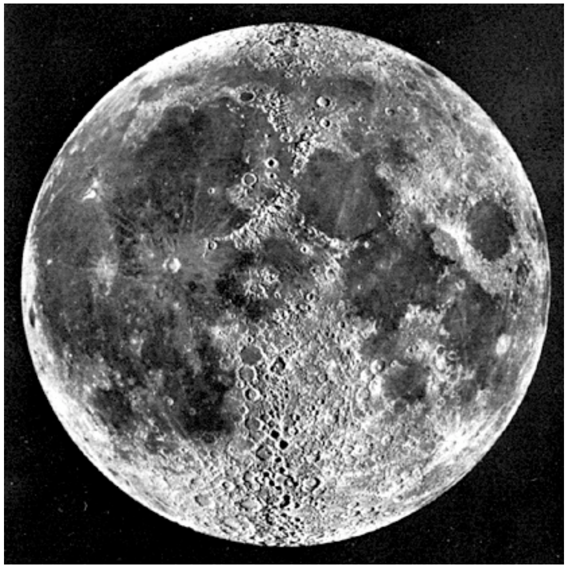
Allegre & Schneider, *Sci. Am.*(1994)

Differentiation of Earth, Continents, Ocean & Atmosphere

- **Differentiation of Earth**
  - Homogenous planetesimal
  - Earth heats up
    - Accretion and *compression* (T~1000°C)
    - *Radioactive decay* (T~2000°C)
  - Iron melts--migrates to center, forming ***core***
    - *Frictional heating* as iron migrates
  - Light materials float--***crust***
  - Intermediate materials remain--***mantle***

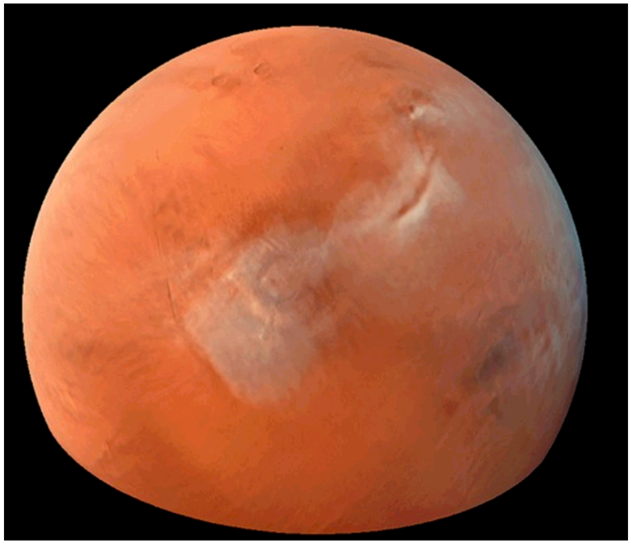
- **Differentiation of Continents, Oceans, and Atmosphere**
  - Continental crust forms from differentiation of primal crust
  - Oceans and atmosphere
  - *Two hypotheses*
    - ***internal***: degassing of Earth's interior (volcanic gases)
    - ***external***: comet impacts add H_2O CO_2, and other gases
  - Early atmosphere rich in H_2, H_2O, N_2, CO_2; deficient in O_2

Earth's Crustal Evolution: 1. 3 Types of Planetary Crust



1° = original crystalline material to solidify from magma oceans of newly accreted bodies. None of this survives on Earth, but the white highlands of the moon are a good example. Impact that created moon produced 1° crust.

2° = slow heating by radioactive decay melts small quantities of rock in planetary interiors. Results in eruption of basaltic lavas. E.g., Earth's ocean floor, surfaces of Mars & Venus, lunar maria.



Taylor & McLennan (1996); NASA-JPL

Origin of the Moon (Artist's Rendition!)

Stanley (1999)

Moon-Forming Impact

Canup R & AspaugE:Eos Trans. AGU, 82(47), Fall Meet. Suppl., Abstract U51A-02, 2001
http://www.swri.edu/9what/releases/canupmoon.htm

Hypothesis for lunar origin - Moon forms from debris ejected as a result of the collision of a roughly Mars-sized impactor with early Earth

- Geophysical simulations use a method known as smooth particle hydrodynamics, or SPH and can achieve resolutions sufficient to study the production of orbit-bound debris necessary to yield the Moon.

- Off-center, low-velocity collisions yield material in bound orbit from which a satellite may then accumulate.

- Simulations must account for mass, angular momentum and compositions of the earth-Moon system.

- Must yield an Earth that retains an iron-rich core and a moon that is appropriately iron-depleted and the right density.

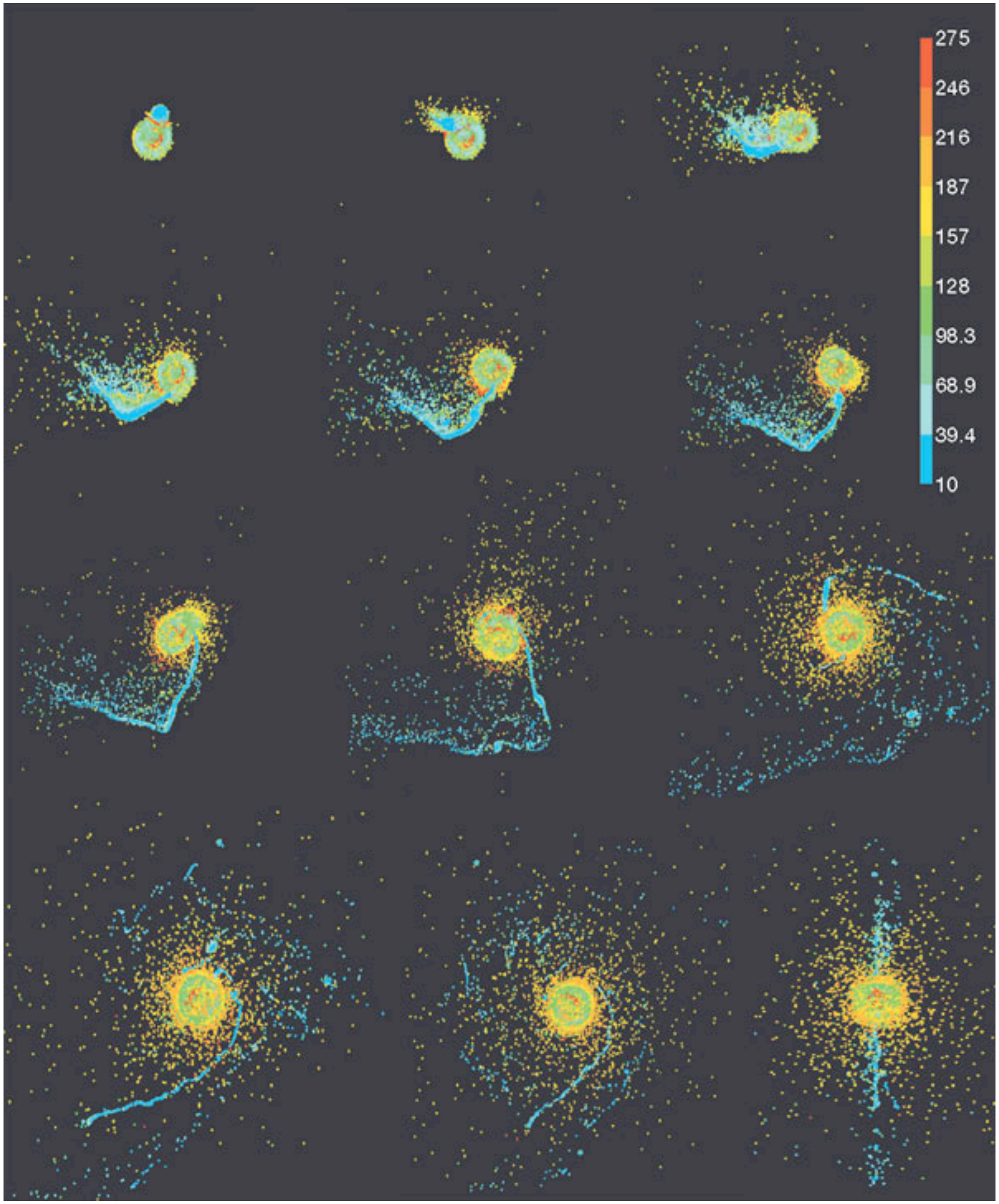
SPH results suggest:

-The object had 10-12% of Earth's mass (Mars-size!)
-Produces a satellite with <3% Fe by mass. Unable to be subsequently captured.
-Happened near end of Earth's accretional history.
-Resulted in melting of Earth crust.

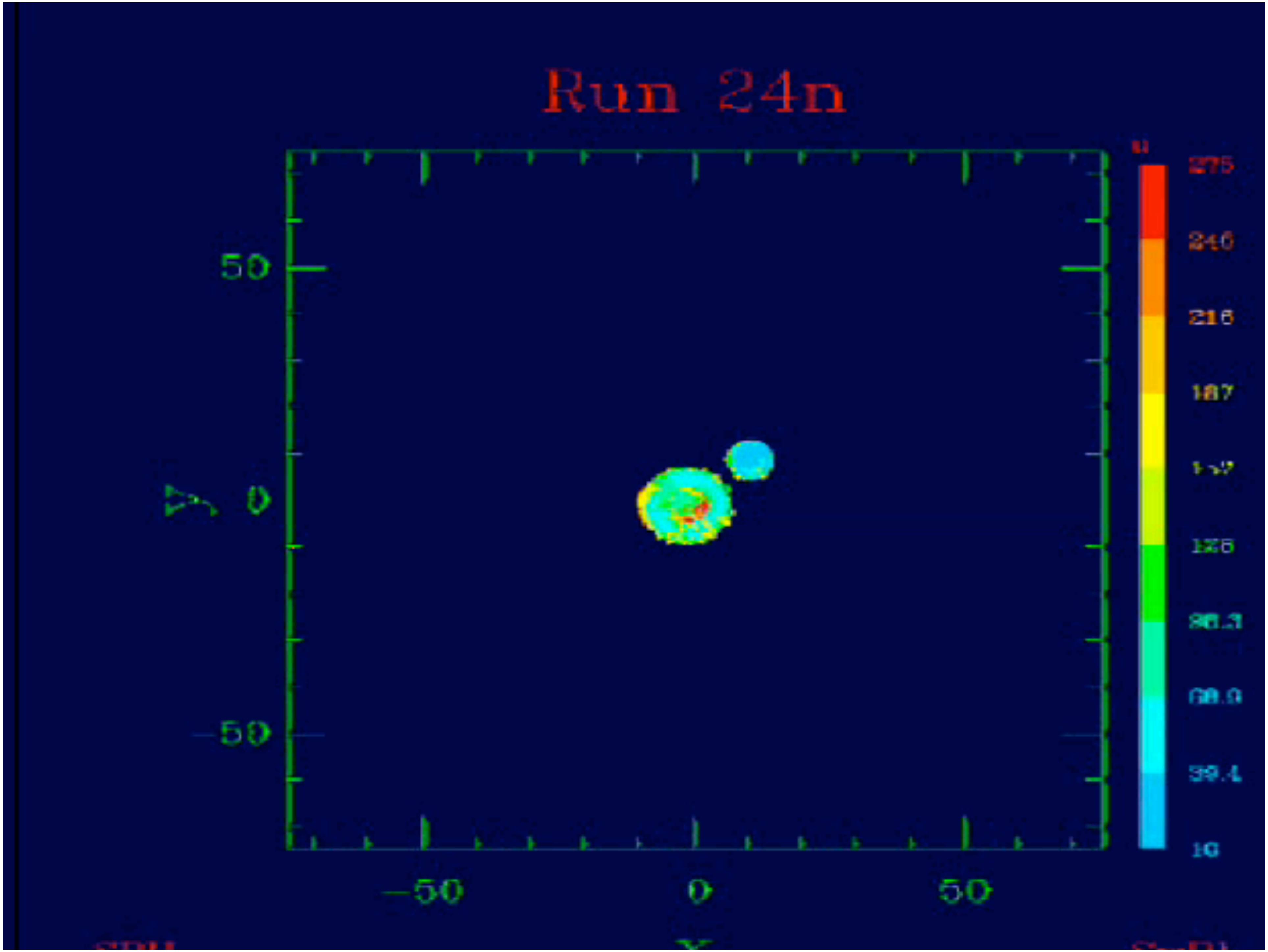
Numerical Simulation of Moon-Formation Event

-Mars-size object (10% M_E) struck Earth
-core merged with Earth
-Moon coalesced from ejected debris of impactor's mantle

-Explains high rotation rate of Earth
-Heat of impact melted crust
-magma ocean #2



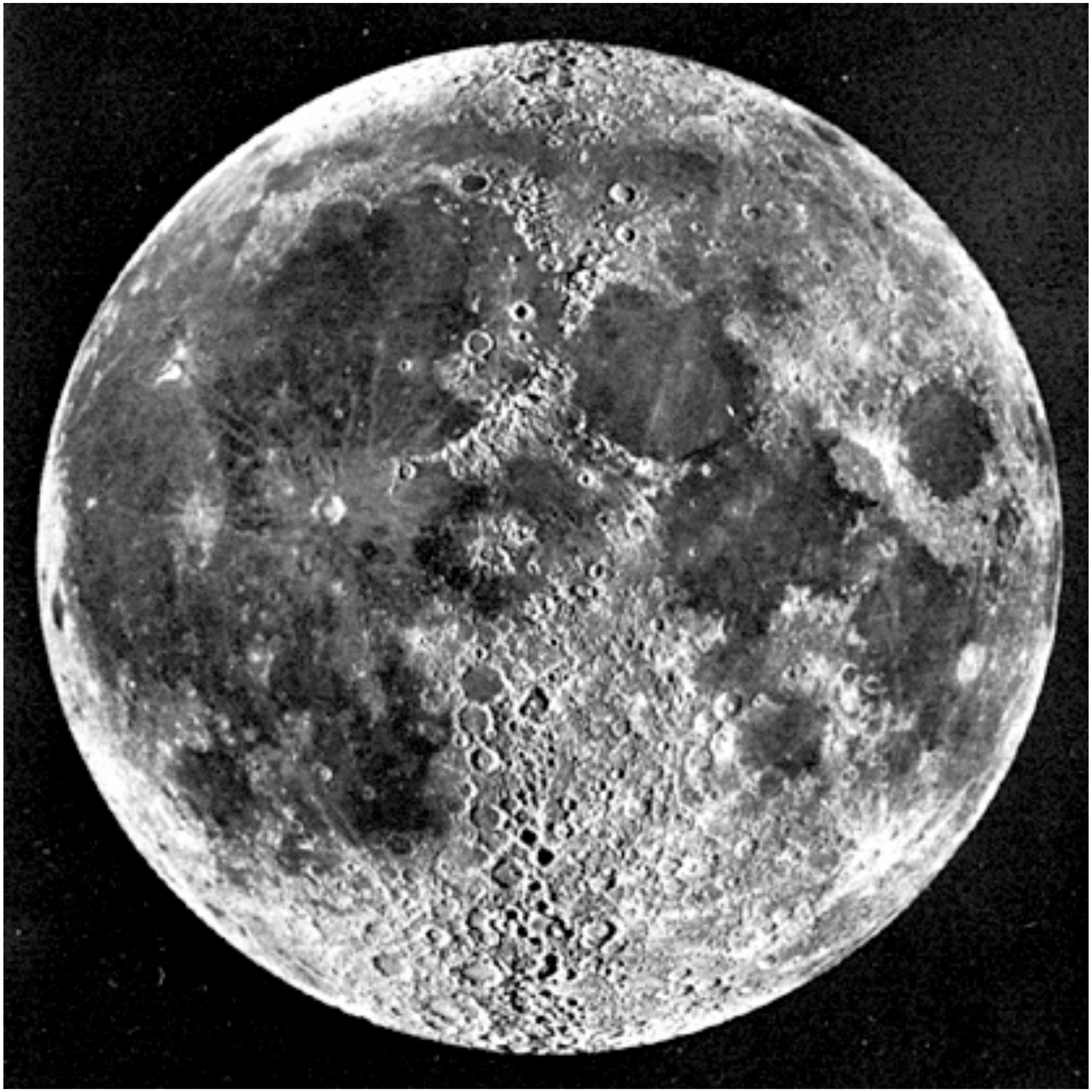
Canup & Asphaug (2001), *Nature*, Vol. 412.


Run 24n
50
0
-50
y
275
216
187

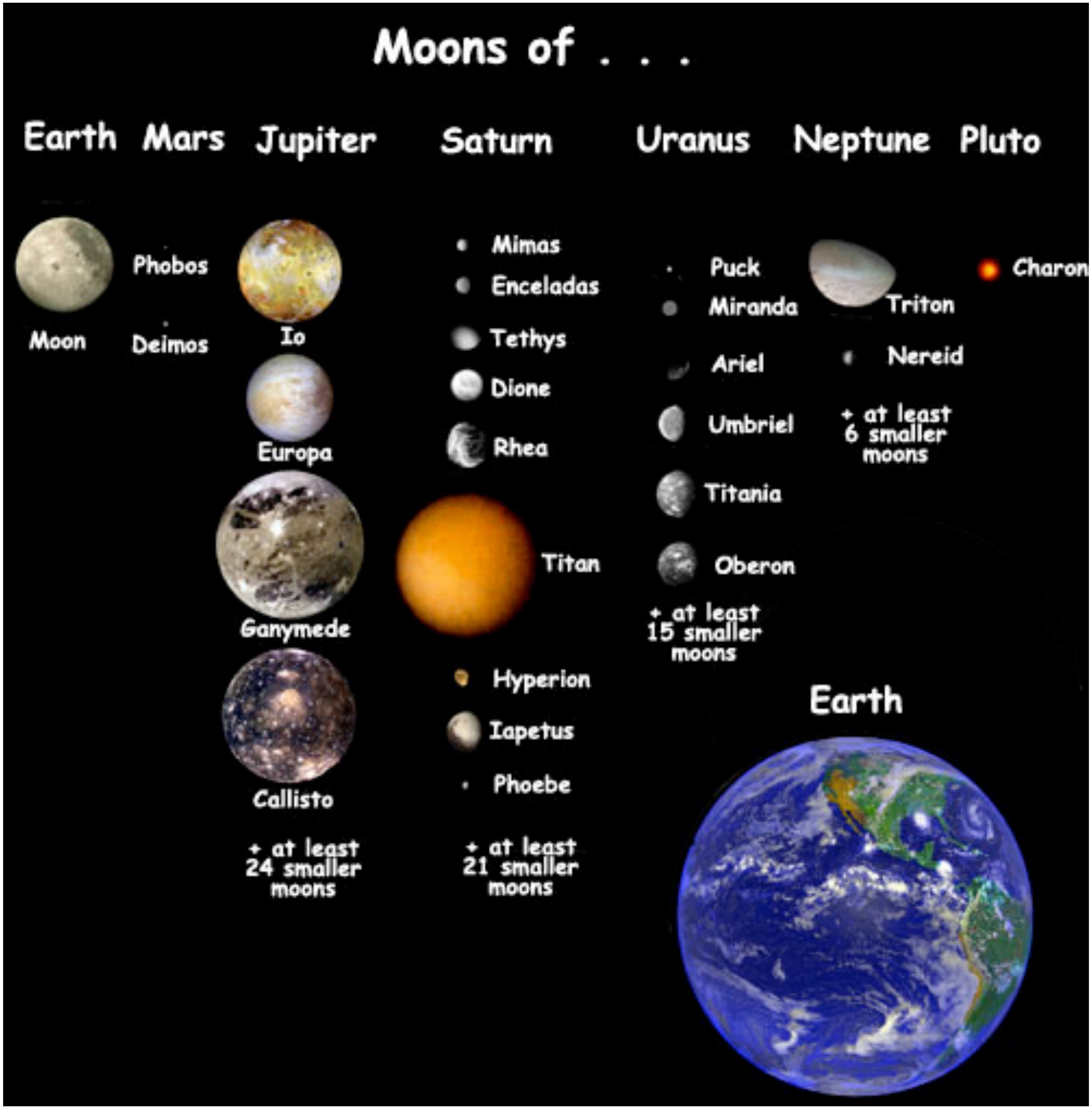
*** Ended Here 10/6/08 ***

The Moon

- Critical to life (stabilizes tilt)
- Rocks from crater rims are 4.6-4.0 Ba (heavy bombardment)
- Jupiter's gravity shielded Earth & Moon from 1,000x more impacts!
- Enormous 'seas' ("**maria**" in latin) are large craters; impact caused melting of crust
- Avg crater diam = 200 km



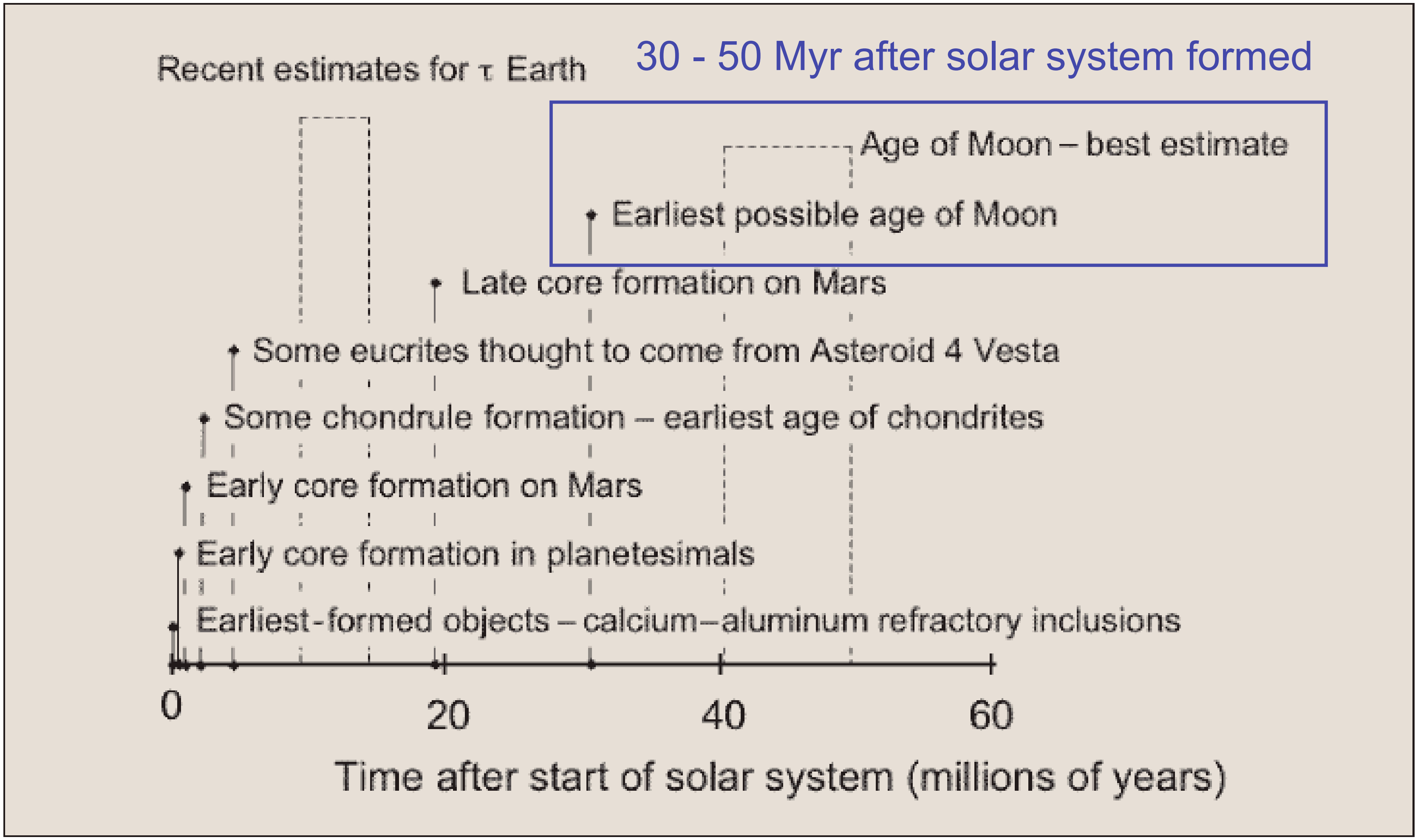
NASA-JPL



Moons of the Solar System to Scale

NASA-JPL

Timeline of Moon & Earth Formation

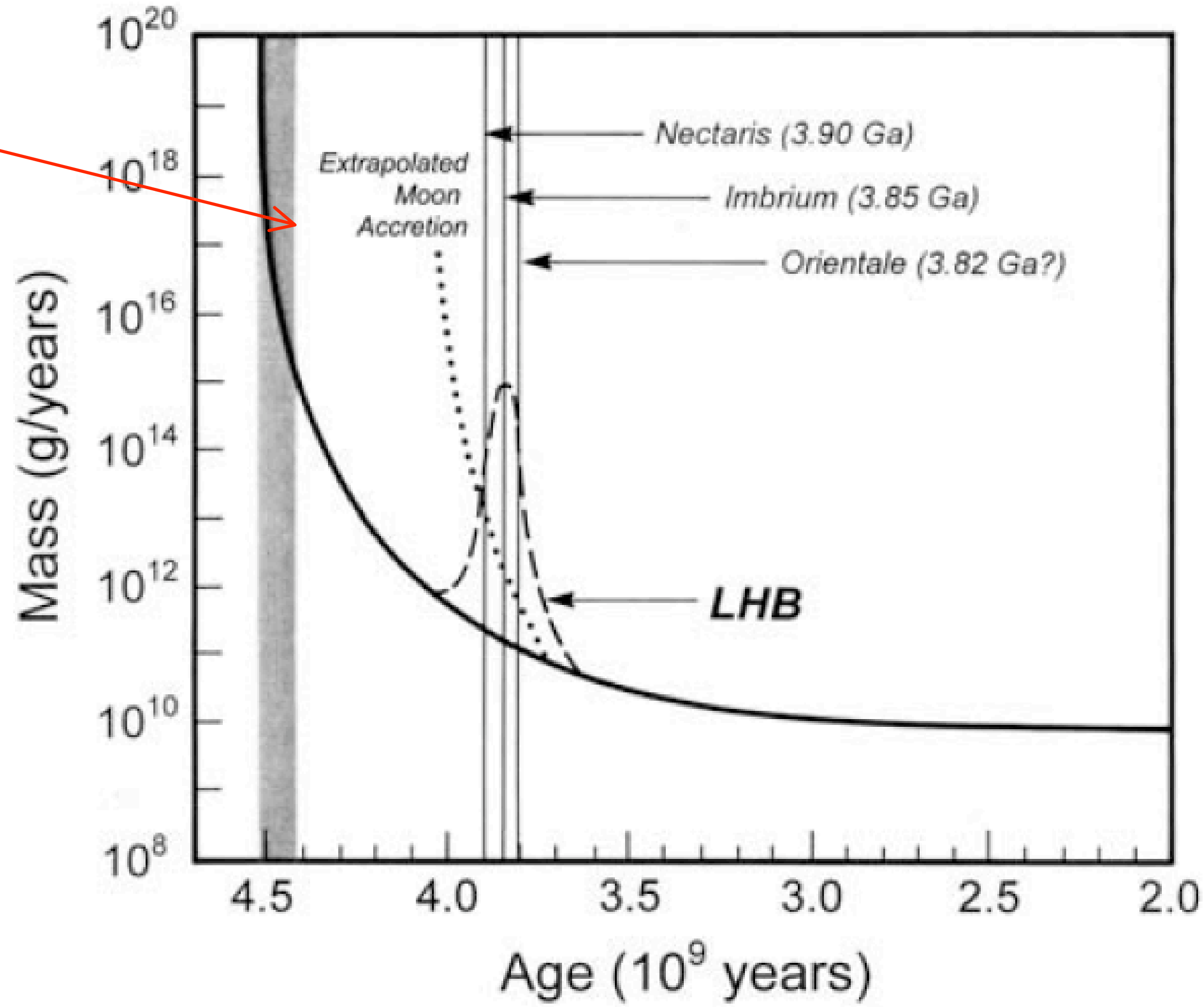


- Earth & Moon ages detrmined from Hf, W & PB isotopes

Halliday (2006) *Elements*, Vol. 2(4): 201-210.

Timeline of Moon Accretion Rate (= mass flux)

- Formation age determined from recent Hf & W isotope measurements
- LHB = Late Heavy Bombardment period
- Ages of largest impact basins shown
- Solid line is present-day background flux extrapolated back to 4.55 Ga

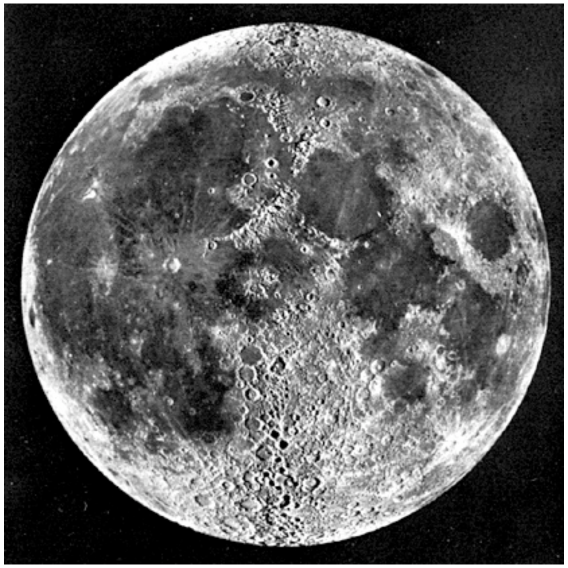


• Dotted line is accretion curve that includes masses of the basin-forming projectiles; unlikely to be correct, b/c implies accretion at 4.1 Ga

• Supports notion of a spike in large impacts (the LHB), possibly associated with migration of Uranus & Neptune towards outer solar system & disturbance of Kuiper Belt objects.

Koeberl (2006) *Elements* Vol. 2(4): 211-216

Earth's Crustal Evolution: 1. 3 Types of Planetary Crust



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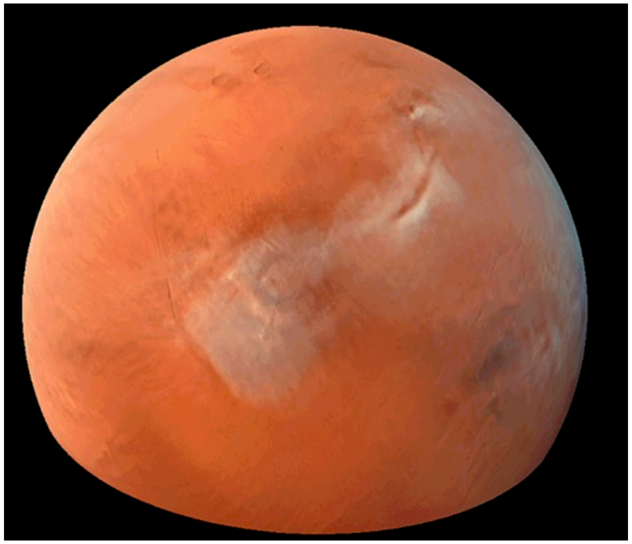
2° = slow heating by radioactive decay melts small quantities of rock in planetary interiors. Results in eruption of basaltic lavas. E.g., Earth's ocean floor, surfaces of Mars & Venus, lunar maria.



Taylor & McLennan (1996); NASA-JPL

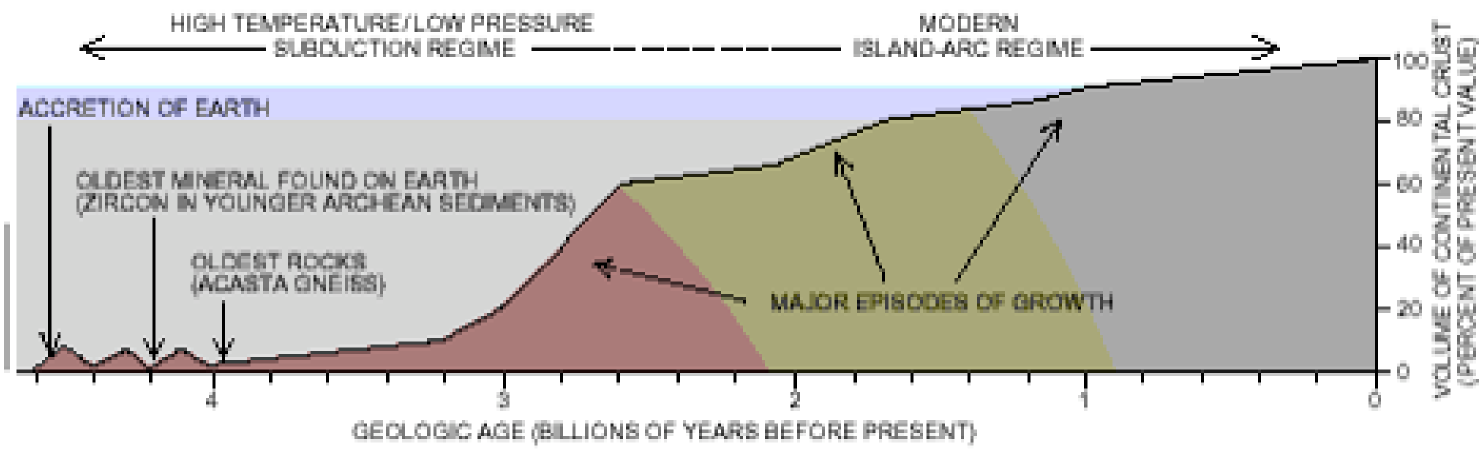
Earth's Crustal Evolution: 2

3° Crust = Formed from slow, continuous distillation by volcanism on a geologically active planet (I.e., plate tectonics).

• Results in highly differentiated magma distinct from basalt. I.e., the low-density, light-colored granite we see in rocks on the continents.

• Earth may be the only planet where this type of crust exists.

• Unlike 1° & 2° crusts, which form in < 200 Myr, 3° crusts evolve over billions of years.



CRUSTAL GROWTH has proceeded in episodic fashion for billions of years. An important growth spurt lasted from about 3.0 to 2.5 billion years ago, the transition between the Archean and Proterozoic eons. Widespread melting at this time formed the granite bodies that now constitute much of the upper layer of the continental crust.

Taylor & McLennan (1996)

Igneous Rocks

Basalt
(2° Crust; Oceanic crust)

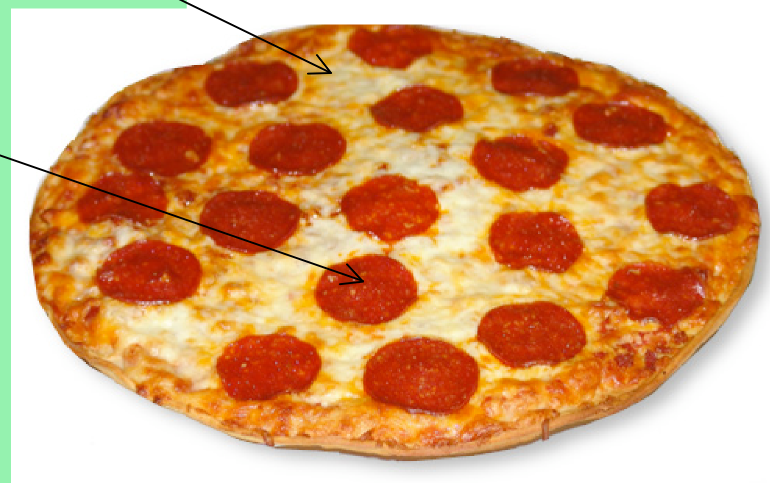
Granite
(3° Crust; Continental Crust)

Stanley (1999)

The Crust & Mantle

<u>The Crust</u>

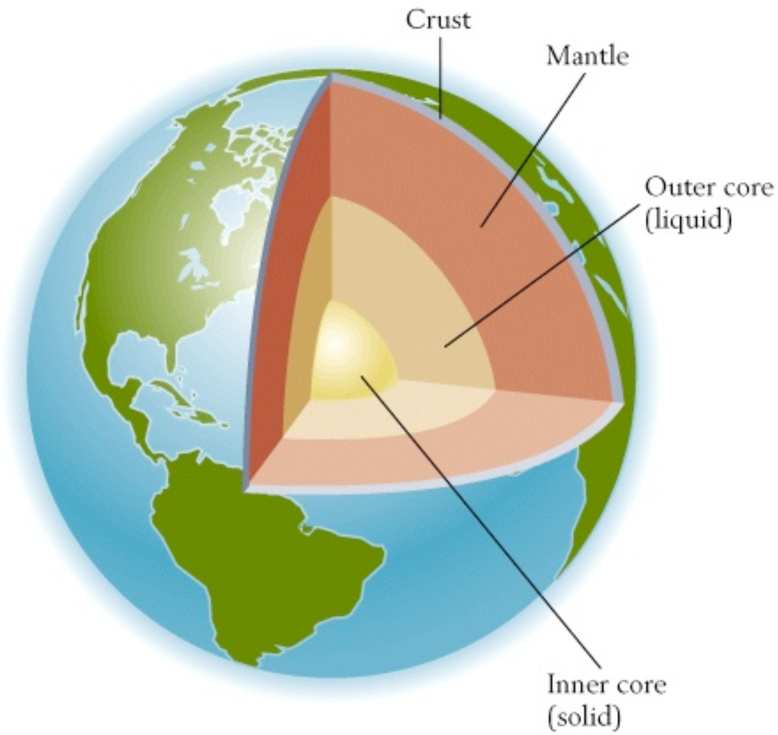
- <u>Ocean Crust</u> (the cheese)
  - 3-15 km thick
  - Basaltic rock
  - Young (<180 Ma)
  - Density ~ 3.0 g/cm3
- <u>Continental Crust</u> (the pepperoni)
  - 35 km average thickness
  - Granitic rock
  - Old (up to 3.8 Ga)
  - Density ~ 2.7 g/cm3
- Crust "floating" on "weak" mantle



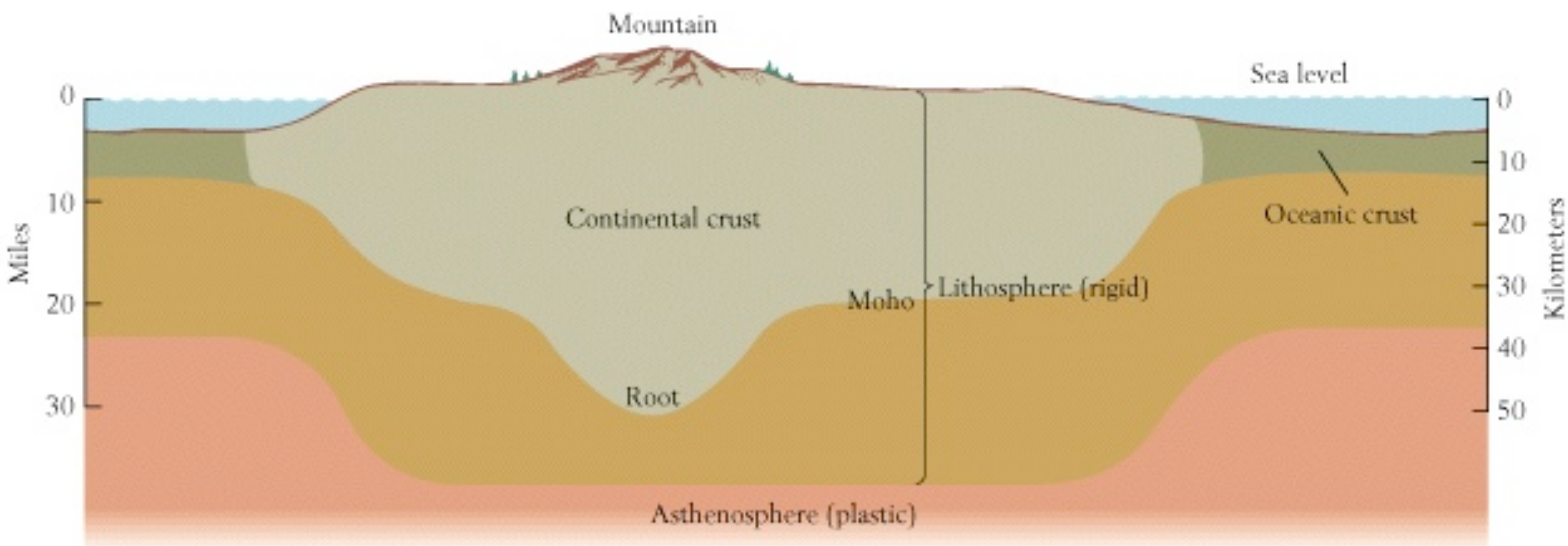
<u>The Mantle</u> (the sauce)

- ~2900 km thick
- Comprises >82% of Earth’s volume
- Mg-Fe silicates (rock)
- Two main subdivisions:
  - Upper mantle (upper 660 km)
  - Lower mantle (660 to ~2900 km; "Mesosphere")

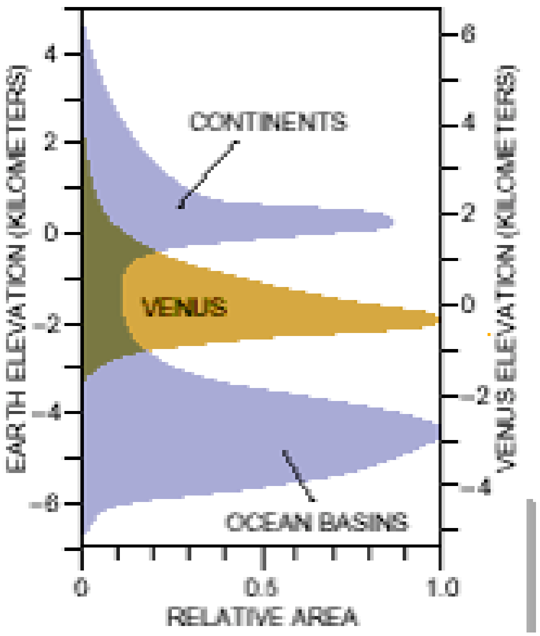
Structure of Earth



From Stanley (1999)



Why is Continental Crust "Elevated Relative to Oceanic Crust?



SURFACE ELEVATIONS are distributed quite differently on the earth (*blue*) and on Venus (*gold*). Most places on the earth stand near one of two prevailing levels. In contrast, a single height characterizes most of the surface of Venus. (Elevation on Venus is given with respect to the planet's mean radius.)

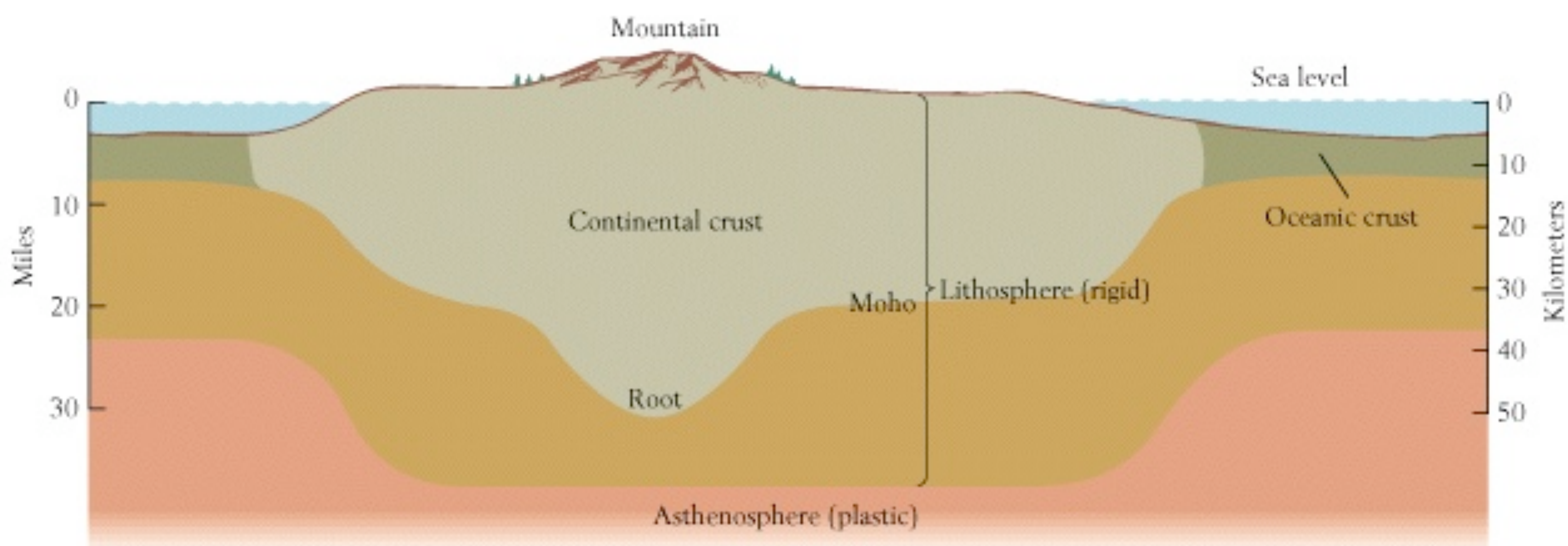
- High-density Basalt sinks into mantle more than low-density Granite.
- Volcanism continually produces highly differentiated continental crust on Earth.
- Venus surface appears to be all basalt.
- Plate tectonics & volcanism do not appear to be happening on Venus (or Mars, Moon).
- So Earth may be unique in Solar System. And plate tectonics & volcanism may be critical in determining habitability.

Taylor & McLennan *Sci. Am.* (1996)

Lithosphere & Asthenosphere

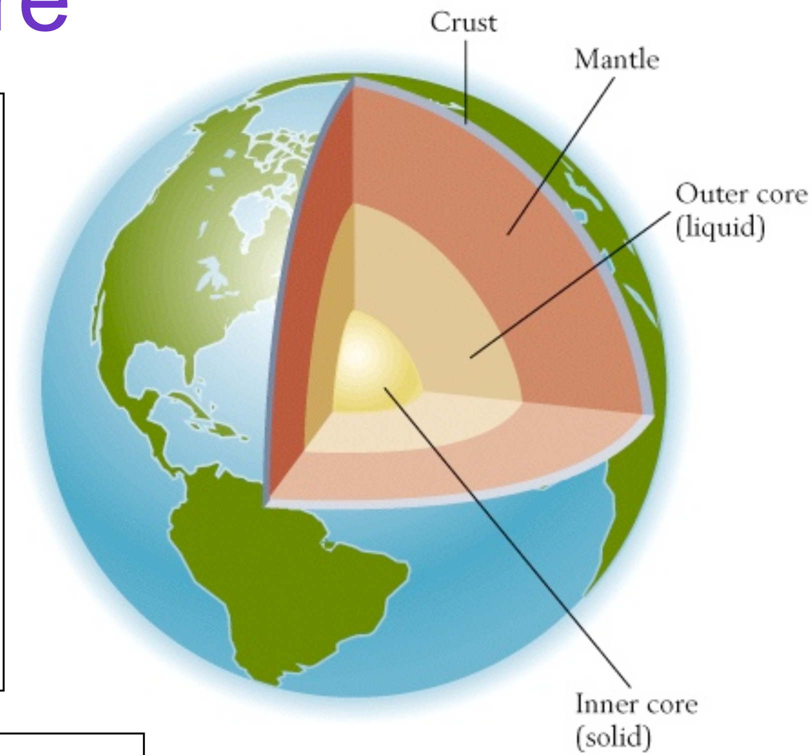
<u>Lithosphere/Asthenosphere</u>: Outer 660 km divided into 2 layers based on mechanical properties. Includes the Mantle + Crust

- ***Lithosphere***
  - Rigid outer layer including crust & upper mantle
  - Averages 100 km thick; thicker under continents
- ***Asthenosphere***
  - Weak, ductile layer under lithosphere
  - Lower boundary ~660 km (entirely within mantle)



The Core

- **<u>Outer Core</u>**
  - ~2300 km thick
  - *Liquid* Fe with Ni, S, O, and/or Si
  - Magnetic field is evidence of flow
  - Density ~ 11 g/cm3
- <u>Inner Core</u>
  - ~1200 km thick
  - *Solid* Fe with Ni, S, O, and/or Si
  - Density ~13.5 g/cm3



▪<u>Earth's Interior</u>: How do we know its structure?

- •Avg density of Earth (5.5 g/cm^3)
- •Denser than crust & mantle
- •Composition of meteorites
- •Seismic wave velocities
- •Laboratory experiments
- •Chemical stability
- •Earth's magnetic field

From Stanley (1999)

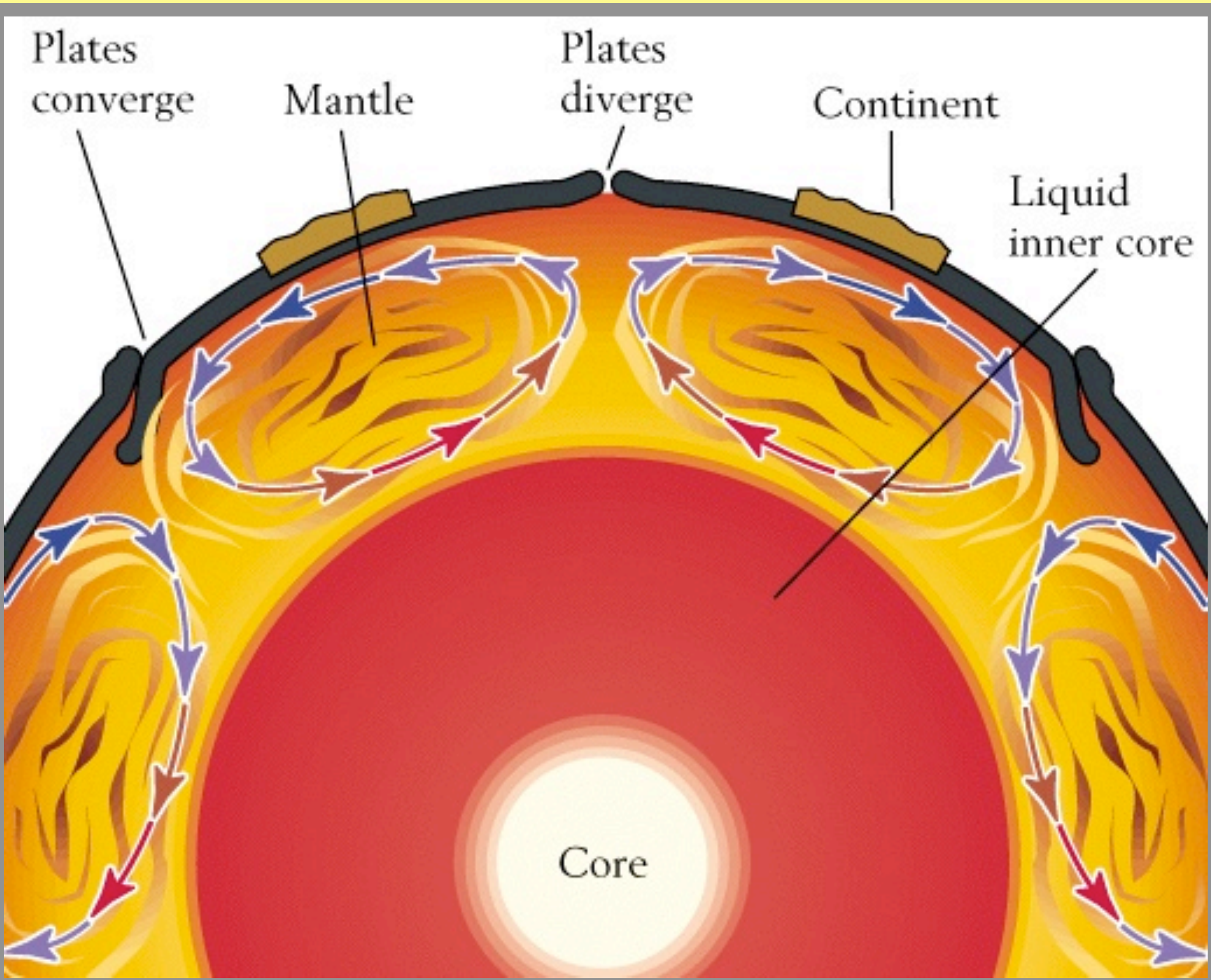
Basics of Geology

Lithospheric Plates

From Stanley (1999)

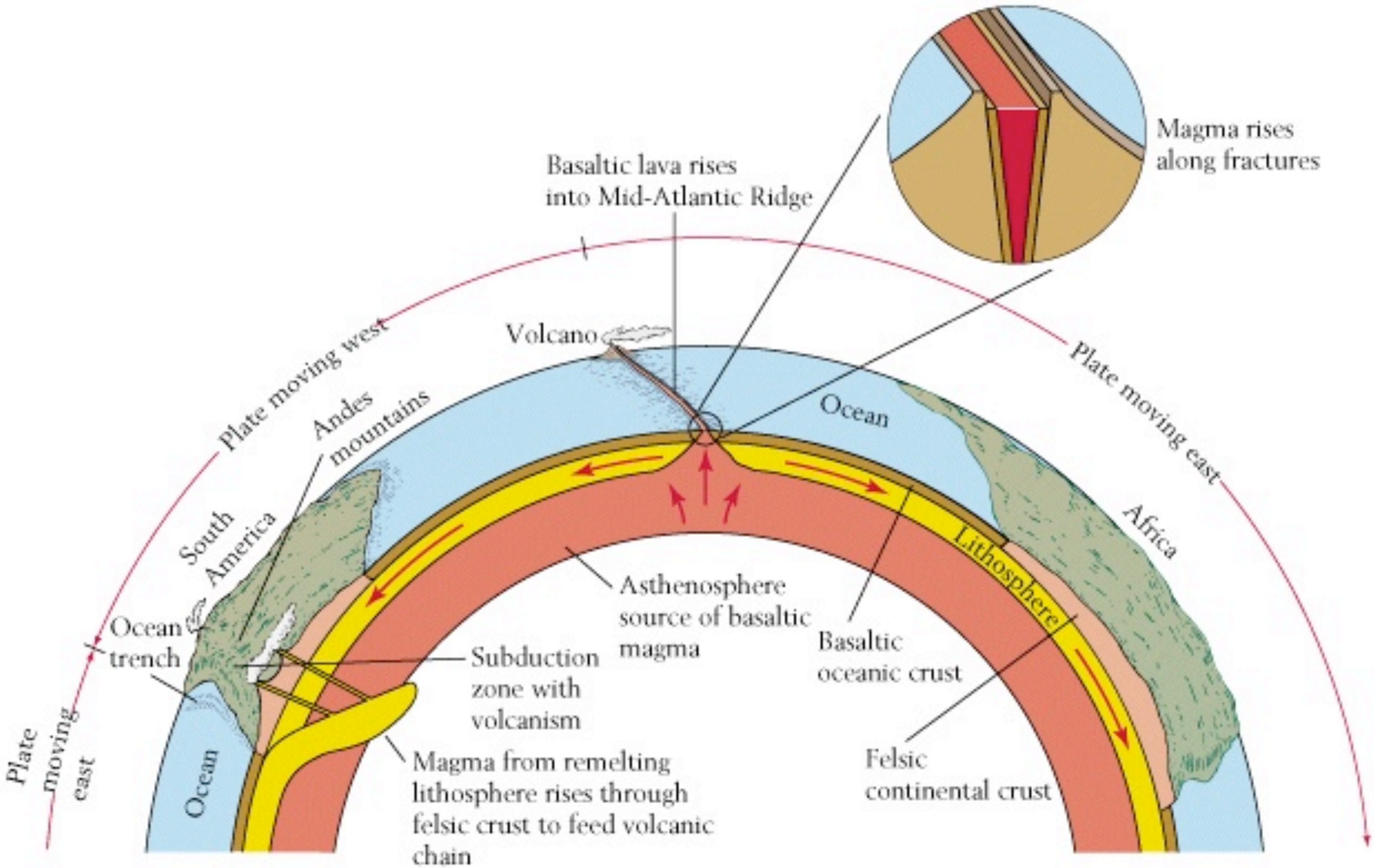
- 8 large plates (+ add'l. small ones)
- Average speed: 5 cm/yr
- 3 types of motion result in 3 types of boundaries: sliding toward (subduction zones), sliding away (ridge axes), sliding along (transform faults)

Convection Drives Plate Movements



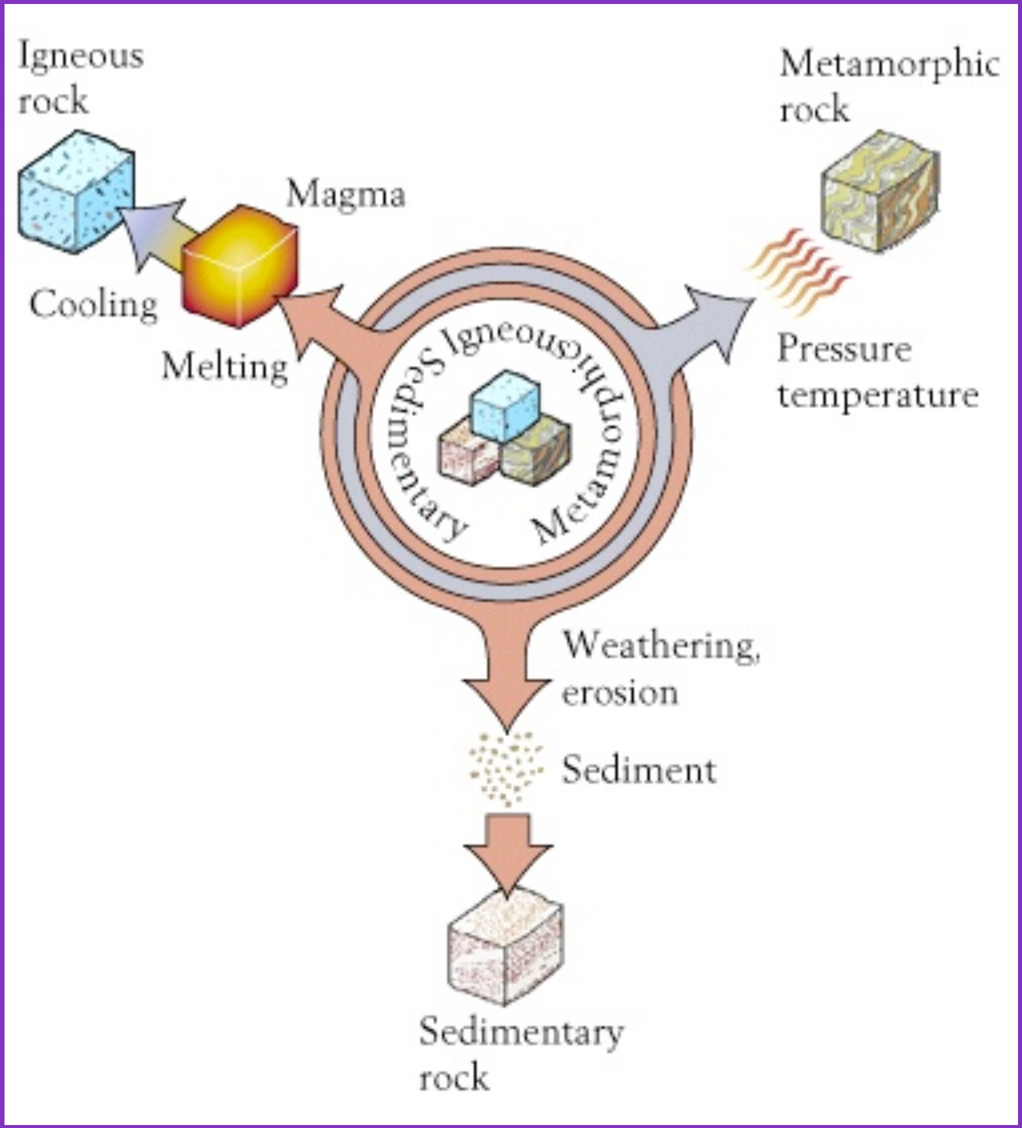
From Stanley (1999)

Tectonic Activity in the South Atlantic



From Stanley (1999)

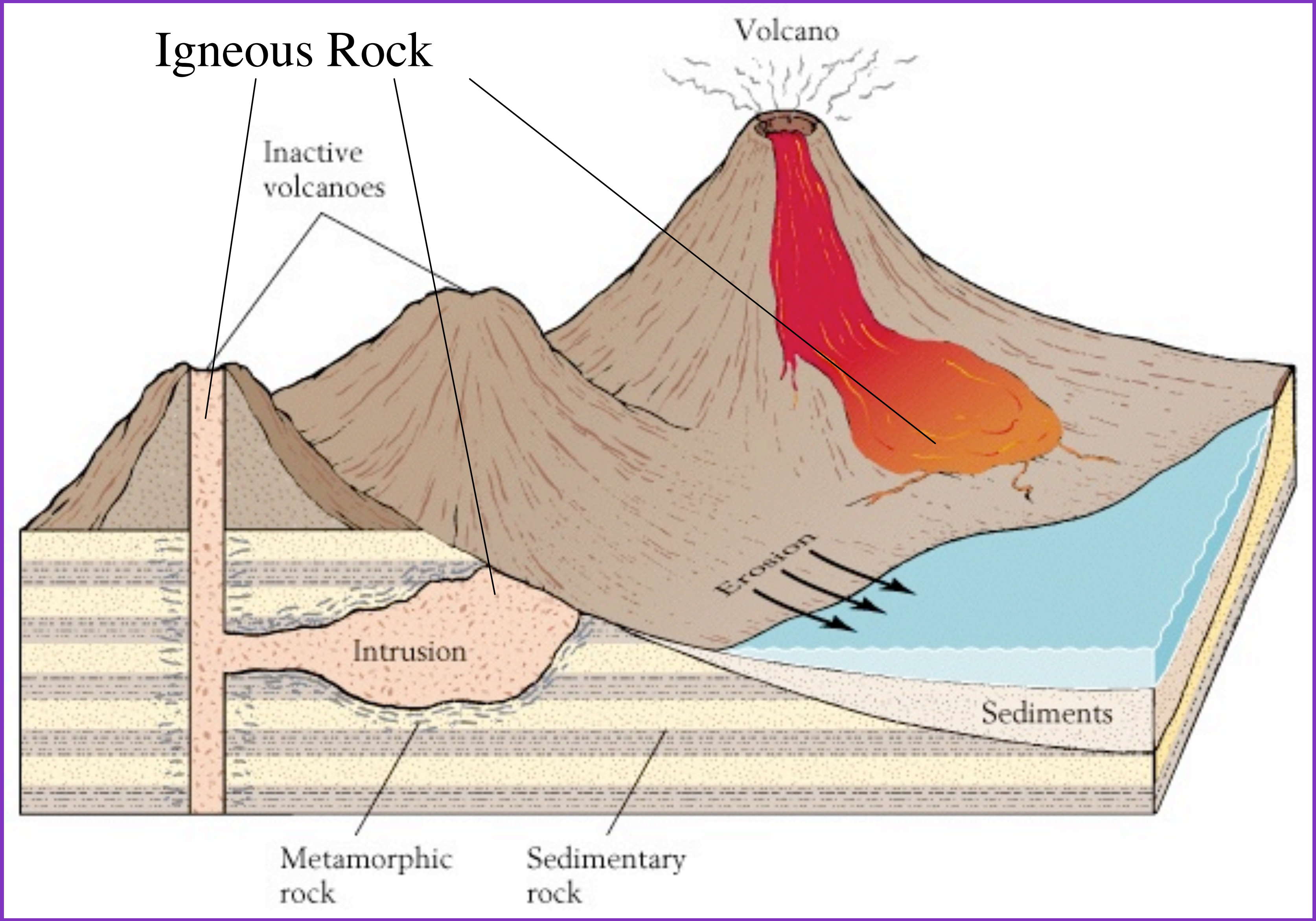
Rock Basics



From Stanley (1999)

Igneous + metamorphic = Crystalline Rocks

The Rock Cycle



From Stanley (1999)

•Felsic: Si-,Al-rich. Light-colored, low-density. Feldspar (pink) & quartz (SiO_2)-rich. Most continental crust. Granite most abundant.
•Mafic: Mg-, Fe-rich. Dark-colored, high-density. Most oceanic crust. Ultramafic rock (more dense) forms mantle below crust.
•Extrusive: cools rapidly; small crystals
•Intrusive: cools slowly; large crystals

Igneous Rocks 101

Basalt
(Oceanic Crust)

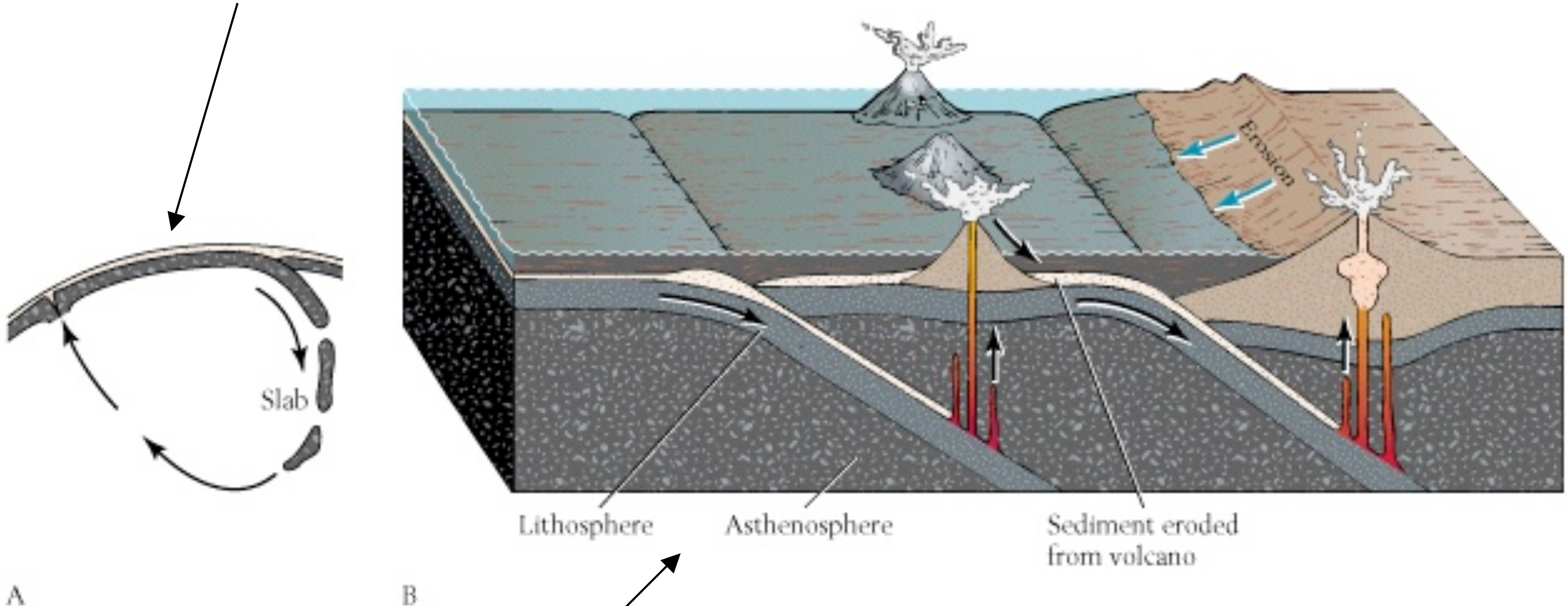
Granite
(Continental Crust)

Stanley (1999)

ME MI
FE FI

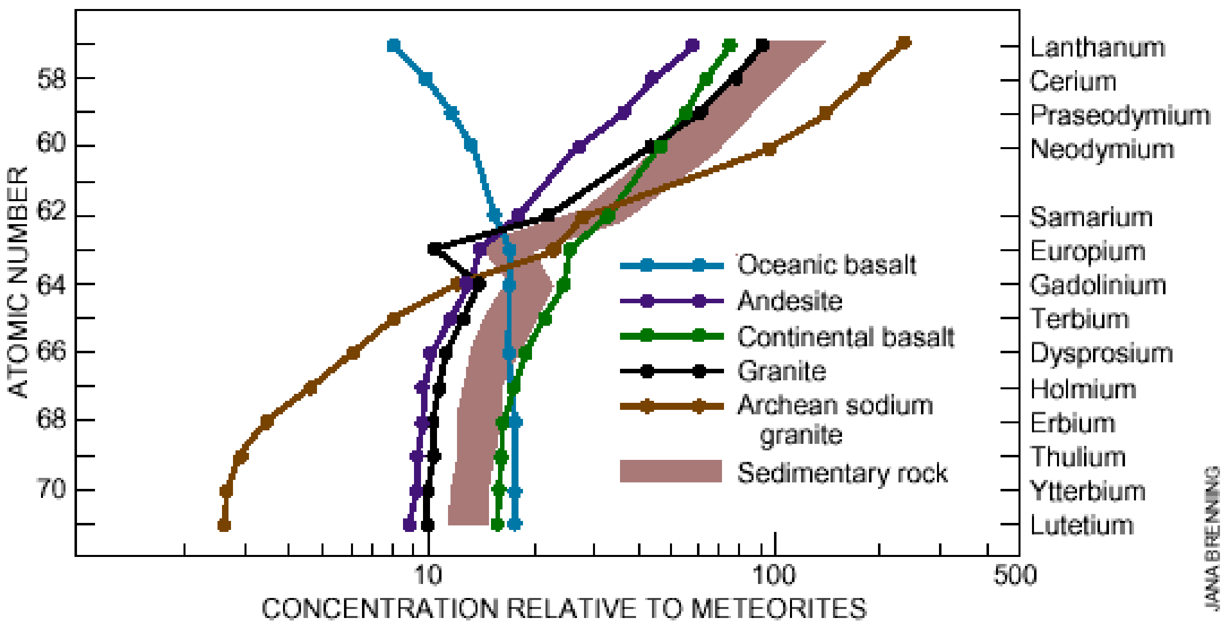
- Slab of lithosphere is subducted, melted & incorporated into asthenosphere
- Convection carries molten material upward where it emerges along a spreading zone as new lithosphere.

Plate Tectonics & the Rock Cycle



•Subducted sediment melts at a shallower depth where it contributes to magma emitted from an island arc volcano and a mountain chain volcano
•Erosion of volcanic rock provides sediment sediment to complete cycle

From Stanley (1999)



RARE-EARTH ELEMENT abundance patterns provide characteristic chemical markers for the types of rock that have formed the earth's crust. Although igneous rocks (those that solidify from magma) can have highly variable rare-earth element signatures (*dotted lines*), the pattern for most sedimentary rocks falls within a narrow range (*gray band*). That uniformity arises because sediments effectively record the average composition of the upper continental crust.

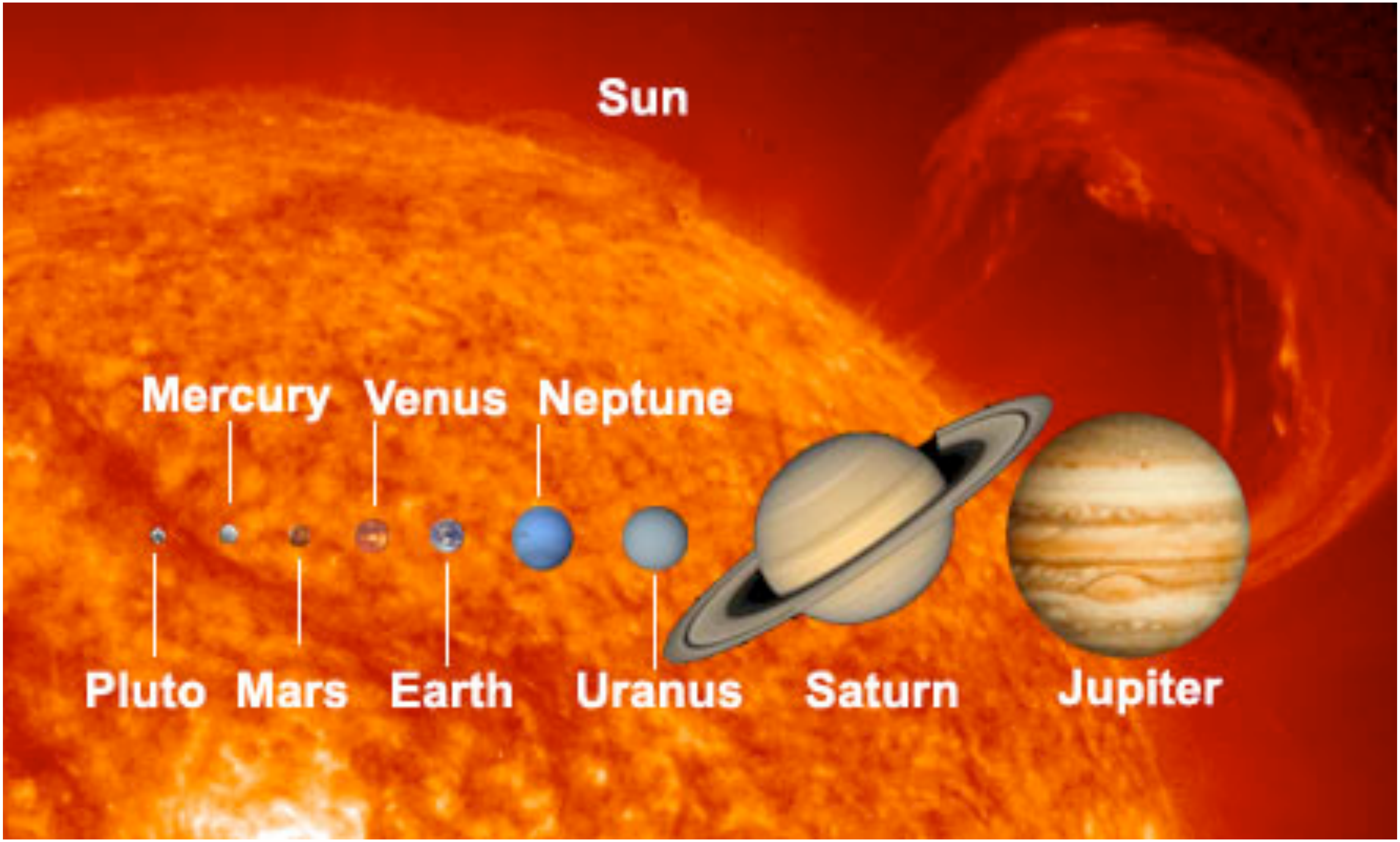
Sedimentary Rocks Represent Homogenous Mixture of Continental Crust

Taylor & McLennan *Sci. Am.* (1996)

*** Ended Here - 10/8/08 ***

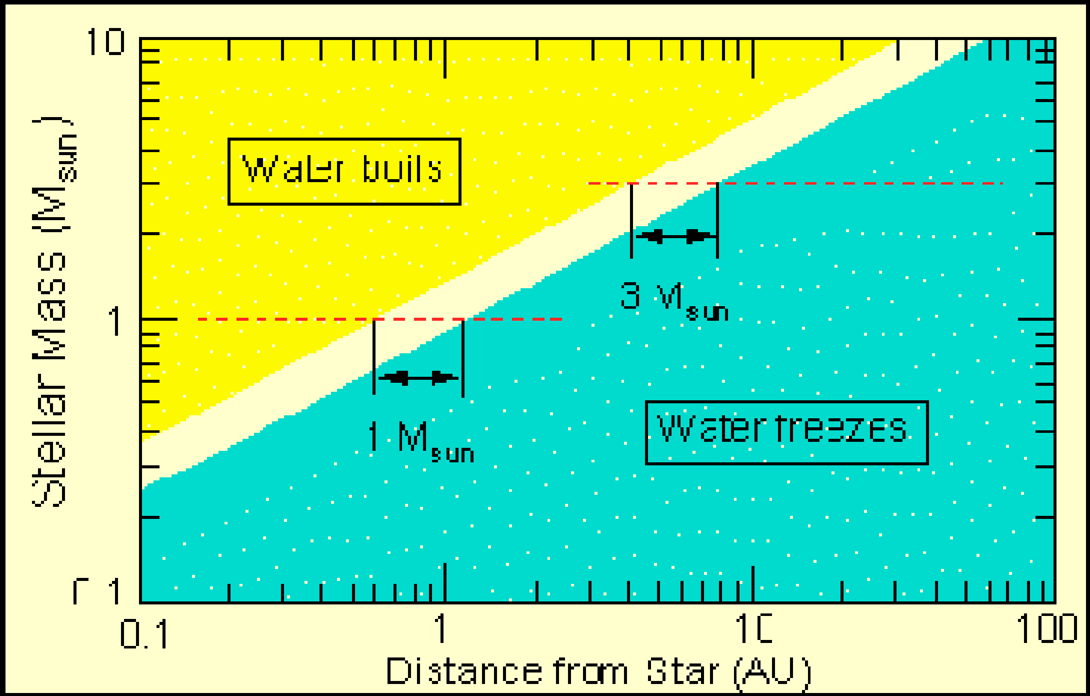
Density of lead = 11.34 g/cm^3
(at Rm. T)

The Habitable Zone of the Solar System
(aka, Top Real Estate Markets outside the Sun)



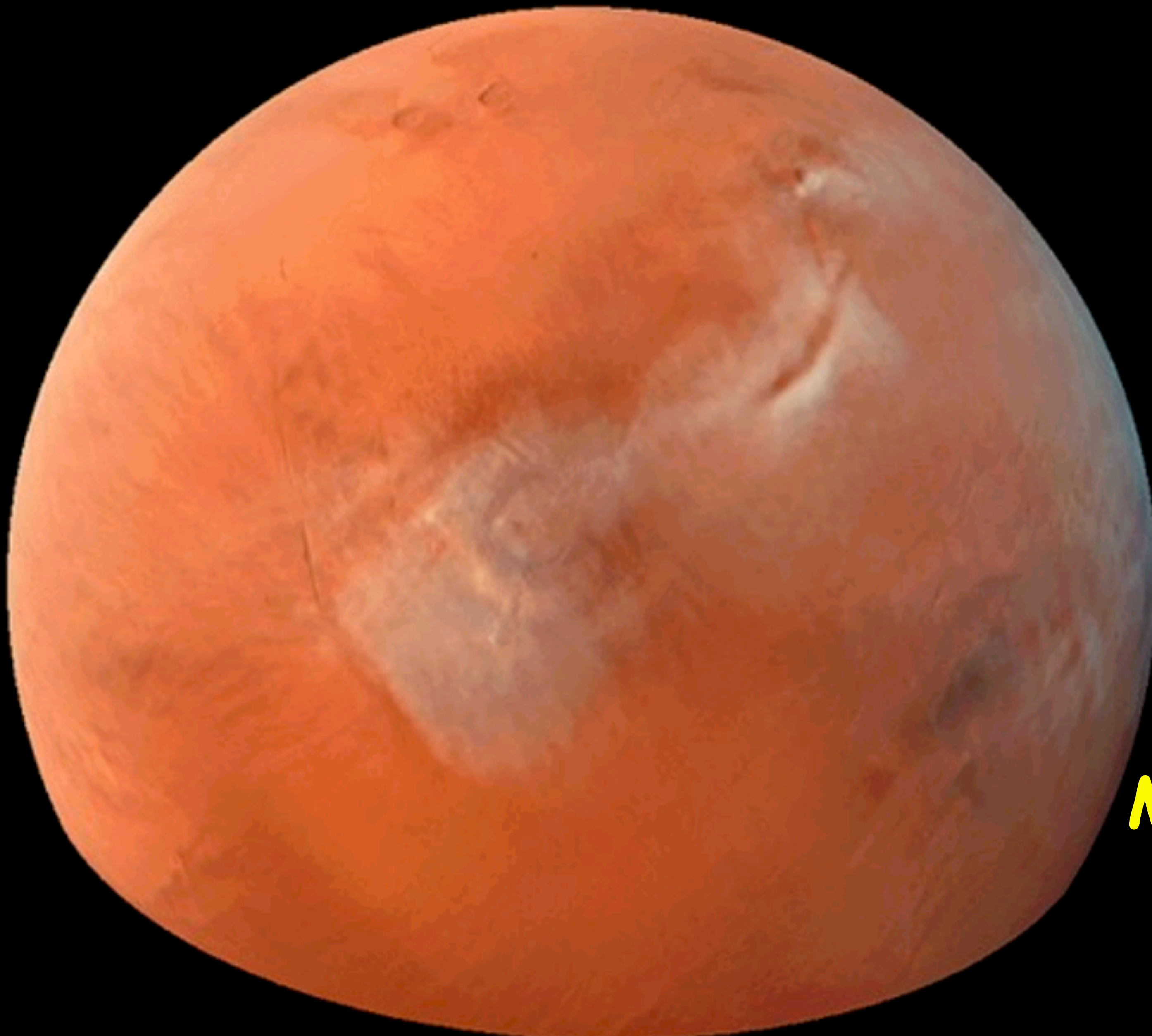
Habitable Zone (Ecosphere)

- Region around a star where planet temperature allows liquid water to exist
- $273 < T_p < 373$ K
- distance of the habitable zone from the star will vary depending on the type of star



480°C
too hot
no water

Venus atm.
90 bar
96% CO_2
3% N_2



-60°C
too
cold.....

Mars Atm.
95% CO_2
2.7% N_2
1.6% Ar

no atmosphere

…..and no life

Atm of
78% N_2

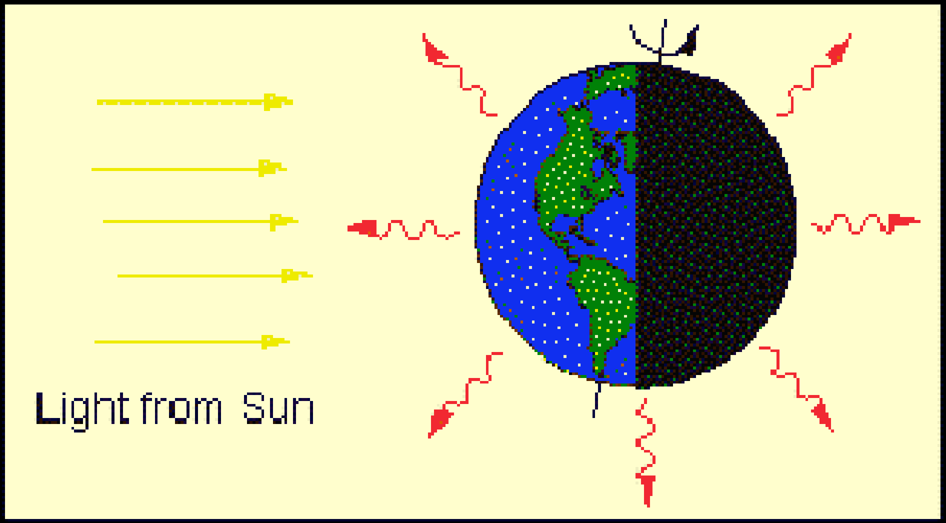
1% Ar, CO_2
and H_2O

15°C
+ oceans

just
right……

What Keeps the Earth Warm?

The power received from the Sun is balanced by heat emission from the Earth.



Simple Planetary Energy Balance

$E_{emitted} = E_{absorbed}$

① $E_{emitted}$

- Blackbody w/ effective radiating temperature, T_e
- Stefan-Boltzmann law

$$E = \sigma T_{eff}^4 \qquad \left(\sigma = 5.67 \times 10^{-8} \frac{W}{m^2 \cdot K^4}\right)$$

→ Energy emitted per unit area

- For entire surface of Earth

$$E_{emitted} = 4\pi R_{Earth}^2 \times \sigma T_{eff}^4$$

(Multiply by area of sphere)

• Blackbody:
emits radiation w/
100% efficiency
@ all λ

Adapted from Kump et al. (1999)

Energy Balance (cont'd.)

② Energy Absorbed

$$E_{absorbed} = E_{intercepted} - E_{reflected}$$

$$= \pi R_E^2 \times S \quad - \pi R_E^2 \times S \times A$$

$$= \pi R_E^2 S (1 - A)$$

Cross section of Earth = area of circle with Earth radius

$$E_{emitted} = E_{absorbed}$$

$$4\pi R_E^2 \times \sigma T_{eff}^4 = \pi R_E^2 S (1 - A)$$

$$\sigma T_{eff}^4 = \frac{S}{4} (1 - A)$$

$$T_{eff} = \sqrt[4]{\frac{S}{4\sigma} (1 - A)}$$

S = solar radiation received at the radius of the planet's orbit around star (so S is a function of the luminosity of the star and the distance the planet is from the star)

A = albedo; the fraction of solar radiation reflected back to space from clouds, ice, deserts, etc.

Adapted from Kump et al. (1999)

The Surface T (T_s) of a Planet can Differ from the Radiating T (T_{eff}) if its Atmosphere Contains Heat-Trapping ("Greenhouse") Gases

$$T_{eff} = \sqrt[4]{\frac{S}{4\sigma}(1-A)}$$

Today: $= 255\,K = -18\,°C$

Earth surface Temp $= 15\,°C$ ← T_s

$T_s - T_{eff} = \Delta T_g$ Greenhouse Effect

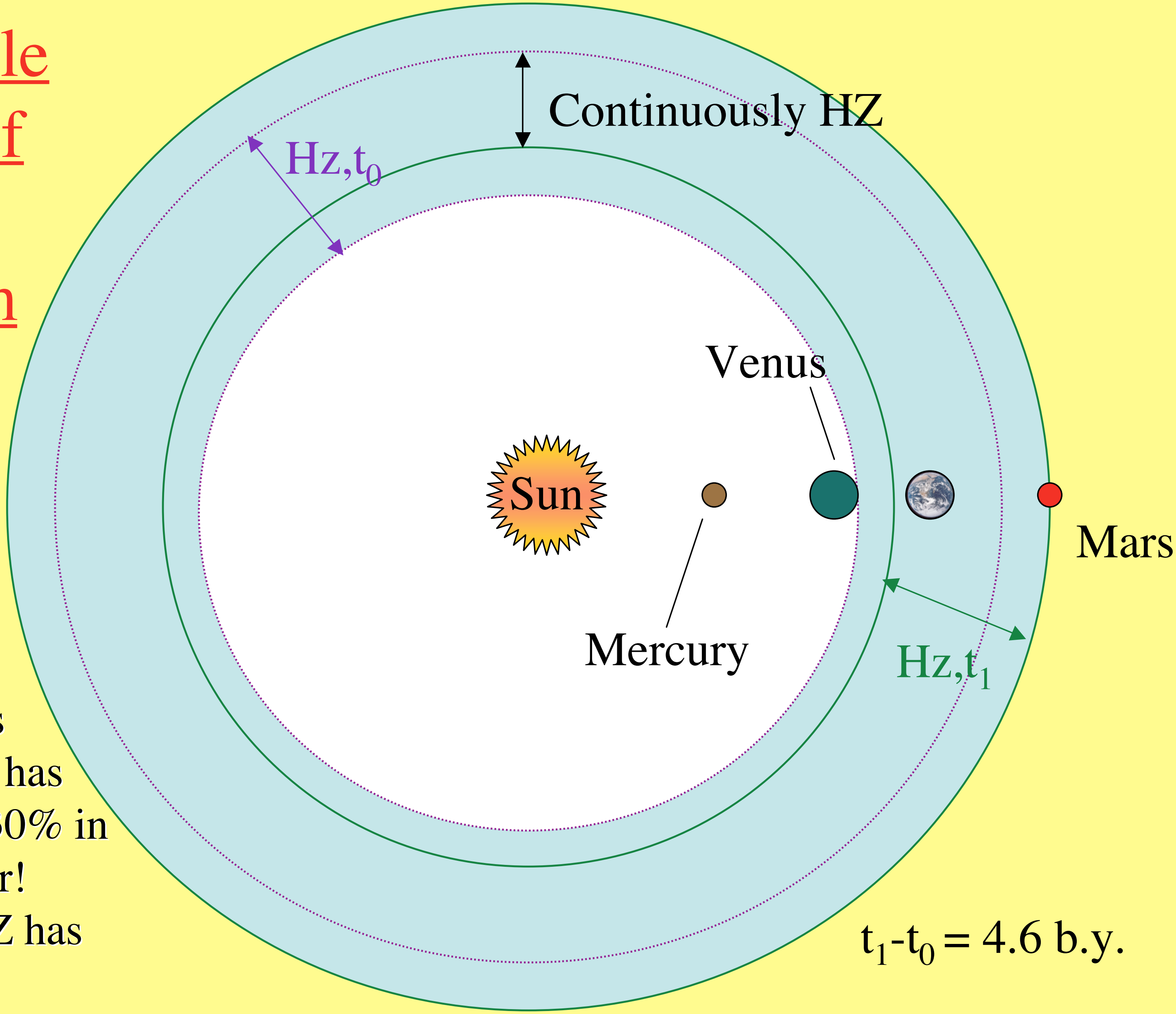
$15° - (-18°) = 33\,°C$

Habitability requires:

$0°C < Ts < 100°C$

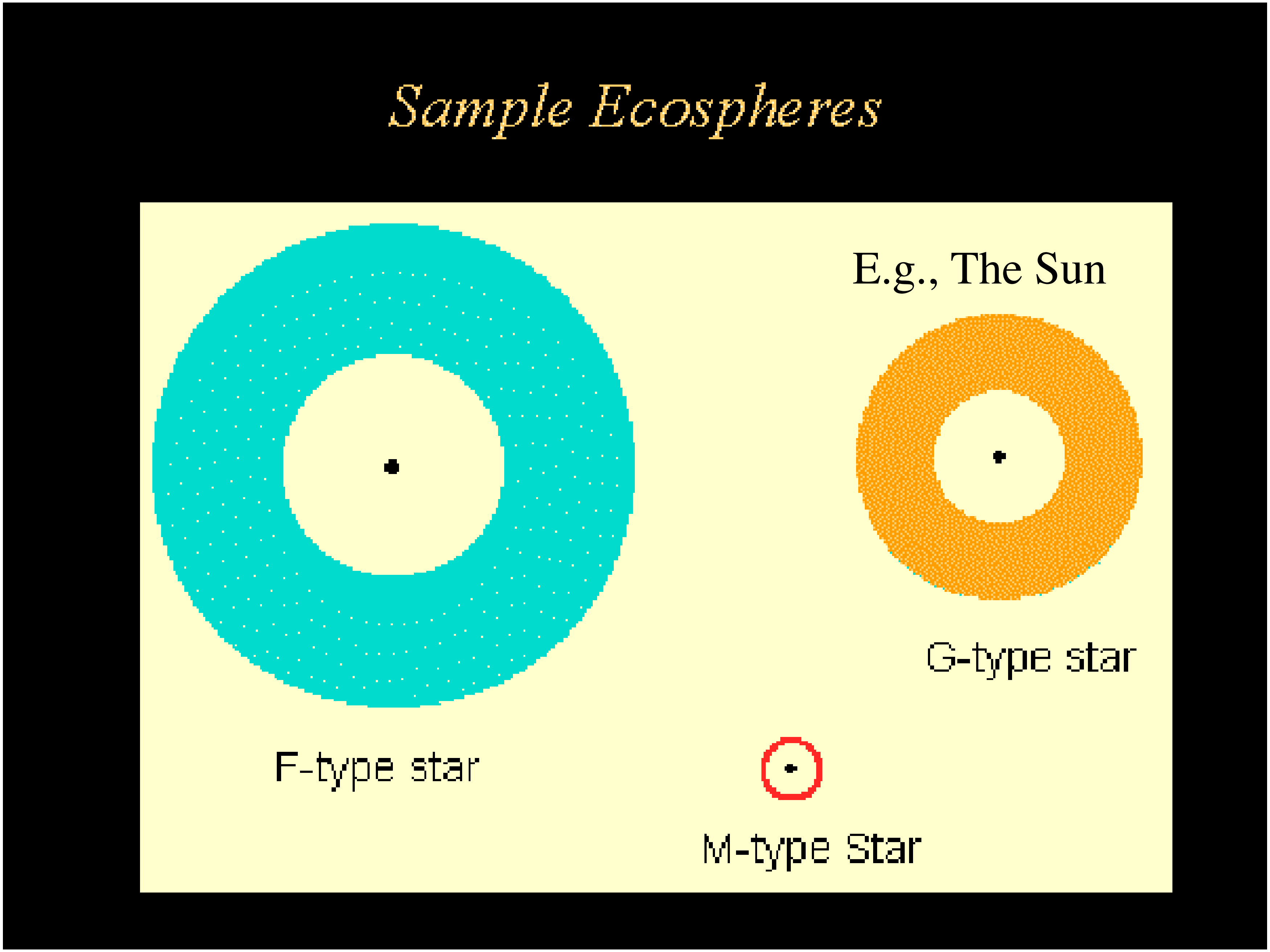
Adapted from Kump et al. (1999)

Habitable Zone of Solar System

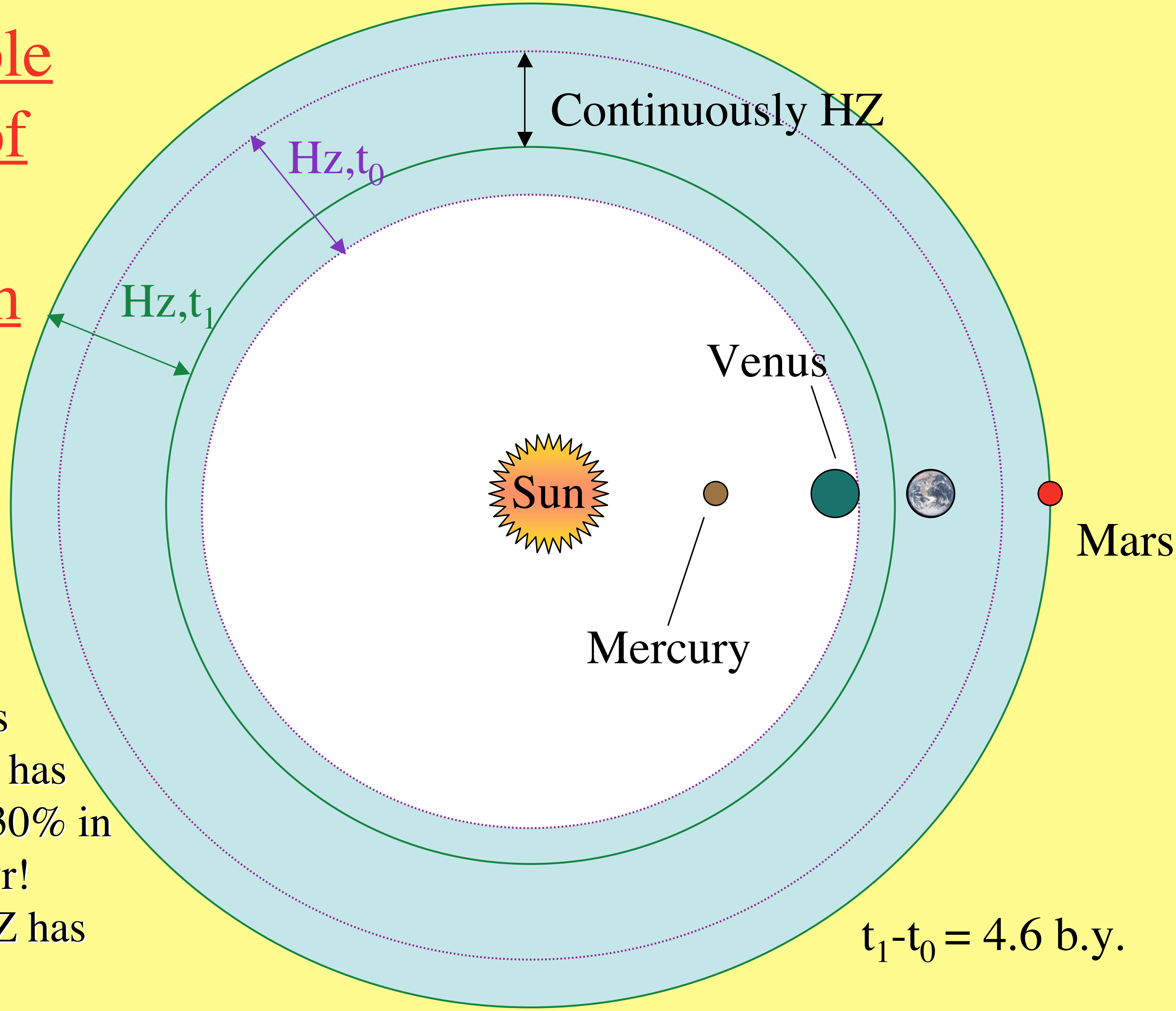


- The sun's luminosity has increased 30% in last 4.6 Gyr!
- So the HZ has moved out

$t_1 - t_0 = 4.6$ b.y.


Sample Ecospheres
E.g., The Sun
G-type star
F-type star
M-type Star

Habitable Zone of Solar System



- The sun's luminosity has increased 30% in last 4.6 Gyr!
- So the HZ has moved out

$t_1 - t_0 = 4.6$ b.y.

Other Considerations Influencing HZ

Caveat: We are relegated to only considering life as we know it & to considering physical conditions similar to Earth

- Greenhouse effect: Increases surface T
  (e.g., Venus, at 0.72 AU, is within HZ, but T_s~745 K!)

- Lifetime of star: larger mass = shorter lifetime
  (must be long enough for evolution)

- UV radiation emission: larger mass = more UV
  (deleterious to life… as we know it)

- Habitable zone moves outward with time
  (star luminosity increases with age)

Further Characteristics of the Habitable Zone

- Liquid water
- Sources of carbon and energy
  - CO_2, organic matter
  - energy from chemistry of rocks + water
  - energy from the sun
- Mechanisms of renewal and recycling
  - Nutrients limited
  - Space = habitat limited (continents…)
  - **Mechanism = Tectonism. Is it that simple?**

The Galactic Habitable Zone

"The Galactic habitable zone (GHZ), analogous to the concept of the circumstellar habitable zone, is an annular region lying in the plane of the Galactic disk possessing the heavy elements necessary to form terrestrial planets and a sufficiently clement environment over several billion years to allow the biological evolution of complex multicellular life."

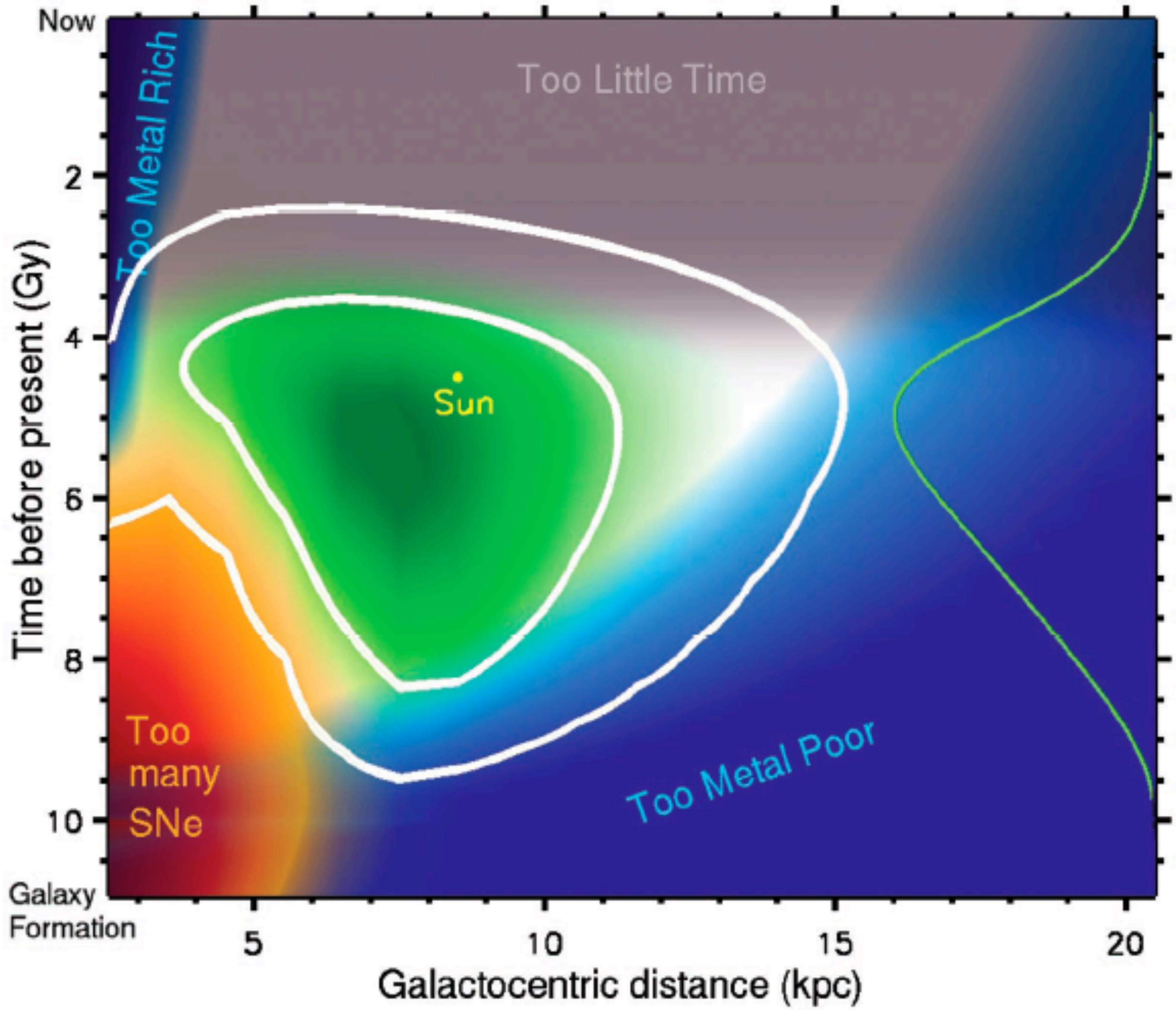
Lineweaver et al. (2004) ***Science*** **Vol. 303: 59-62.**

Habitable Zone of the Milky Way Galaxy

Requirements

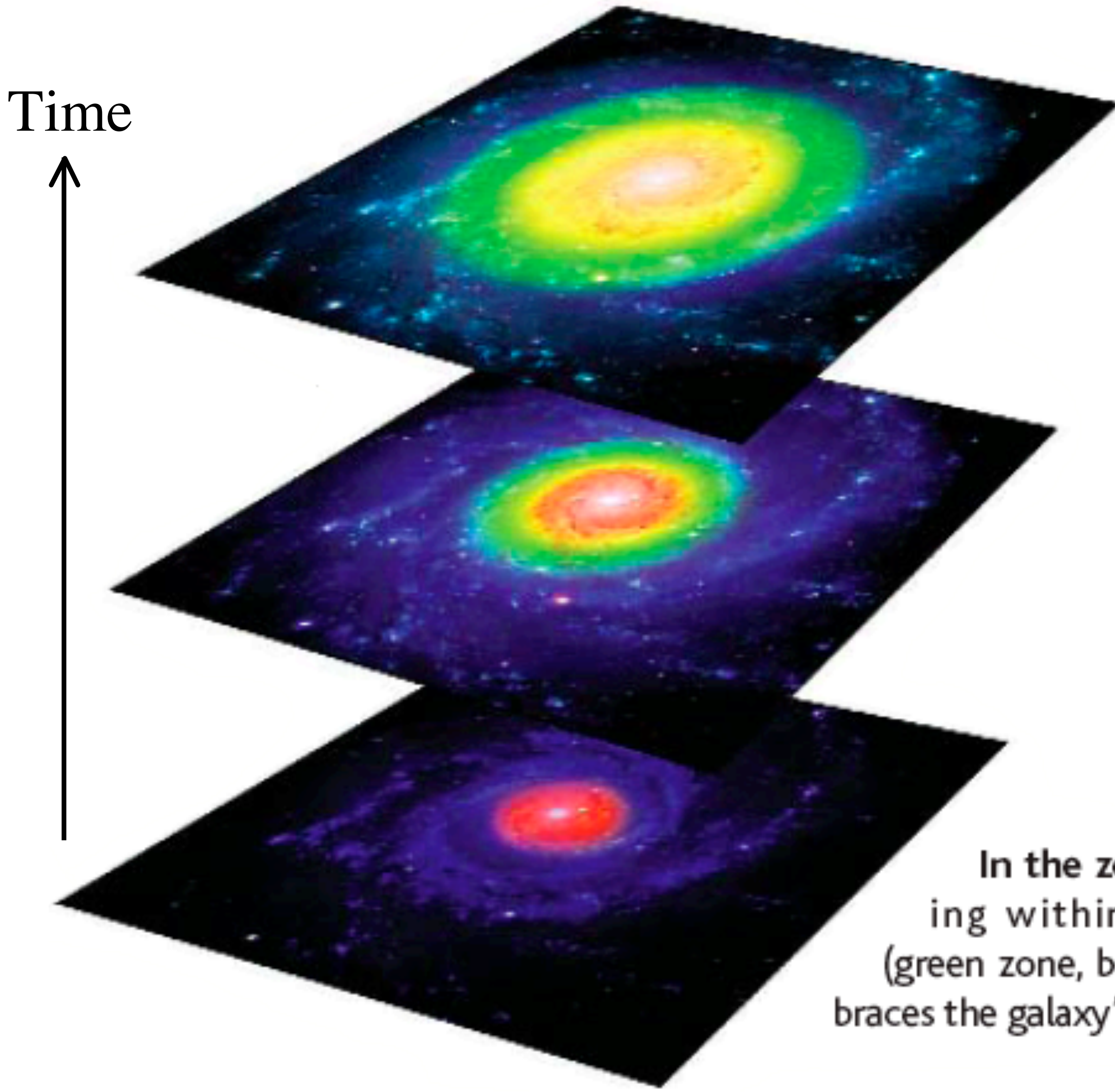
- Metals (Fe)
- Protection from supernovae
- Time for evolution

1 Parsec = 3.26



Lineweaver et al. (2004) *Science* Vol. 303: 59-62.

Milky Way HZ



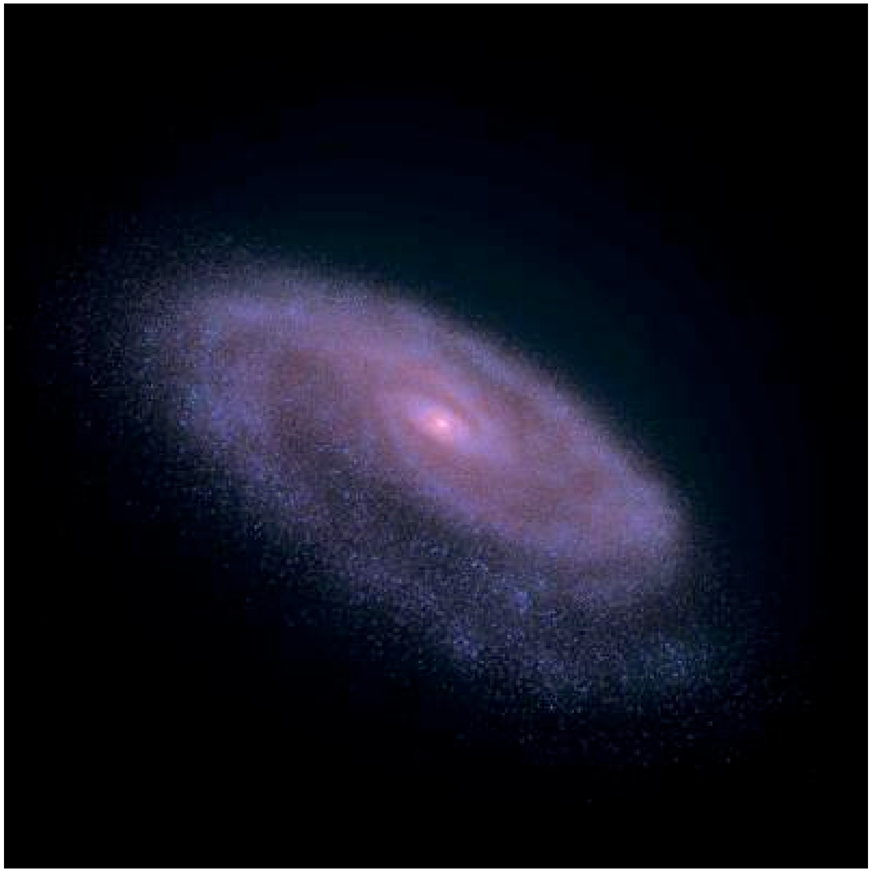
In the zone. A ring spreading within the Milky Way (green zone, bottom to top) embraces the galaxy's life-friendly stars.

Irion (2004) *Science* Vol. 303: 26-27.


Caveat: If Stars (& therefore solar systems) move around the galaxy then the concept of a Galactic Habitable Zone is less tenable….
Roskar et al. (2008) Riding the Spiral Waves: Implications of Stellar Migration for the Properties of Galactic Disks. *The Astrophysical Journal Letters*, 684(2), L79-L82

Stars Appear to Migrate Long Distances In Spiral Galaxies like the Milky Way

- The sun might have traveled far from where it formed, contradicting a belief that stars generally remain static
- According to UW astronomers, 9/16/08, using "*N*-body + smooth particle hydrodynamics simulations of disk formation" (100,000 hrs of computer time!)
- May challenge idea of "habitable zones" in galaxies -- where metal abundances, radiation, water, etc. are amenable to life



Roskar et al. (2008) Riding the Spiral Waves: Implications of Stellar Migration for the Properties of Galactic Disks. *The Astrophysical Journal Letters*, 684(2), L79-L82.

Immigrant Sun: Our Star Could Be Far From Where It Started In Milky Way, *Science Daily*, 9/16/08

Sun might be a long-distance traveler, *UPI.com*, 9/16/08

Simulation of Spiral Galaxy formation: http://www.astro.washington.edu/roskar/astronomy/12M_hr_rerun_angle.mpg

Views of The Early Earth

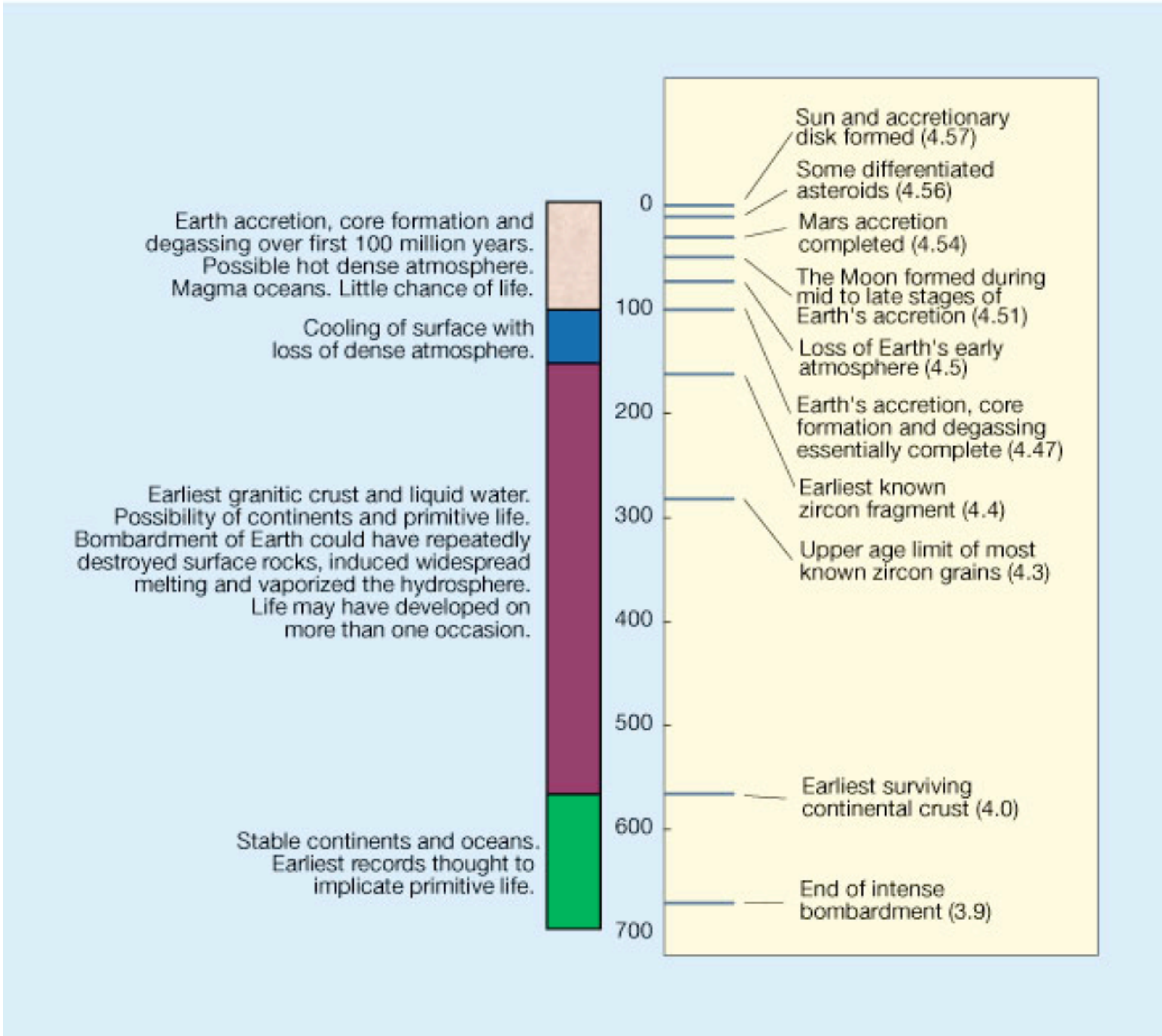
- Artist's view of Hadean Earth, 1952 cover of *Life*
- While such energetic conditions prevailed for ~50 Myr after Earth formed, recent studies indicate a more clement & cool climate within 200–300 Myr after formation

From *Elements Magazine*, August 2006, p. 201.

- Artist's view of cool early Earth ca. 4.2 Ga
- Meteorite impacts still common on Earth & Moon
- Moon in closer orbit & dark maria not yet formed
- Volcanism & other magmatic processes on Earth start to form granitic rocks & proto-continental crust
- Liquid water covered much of Earth

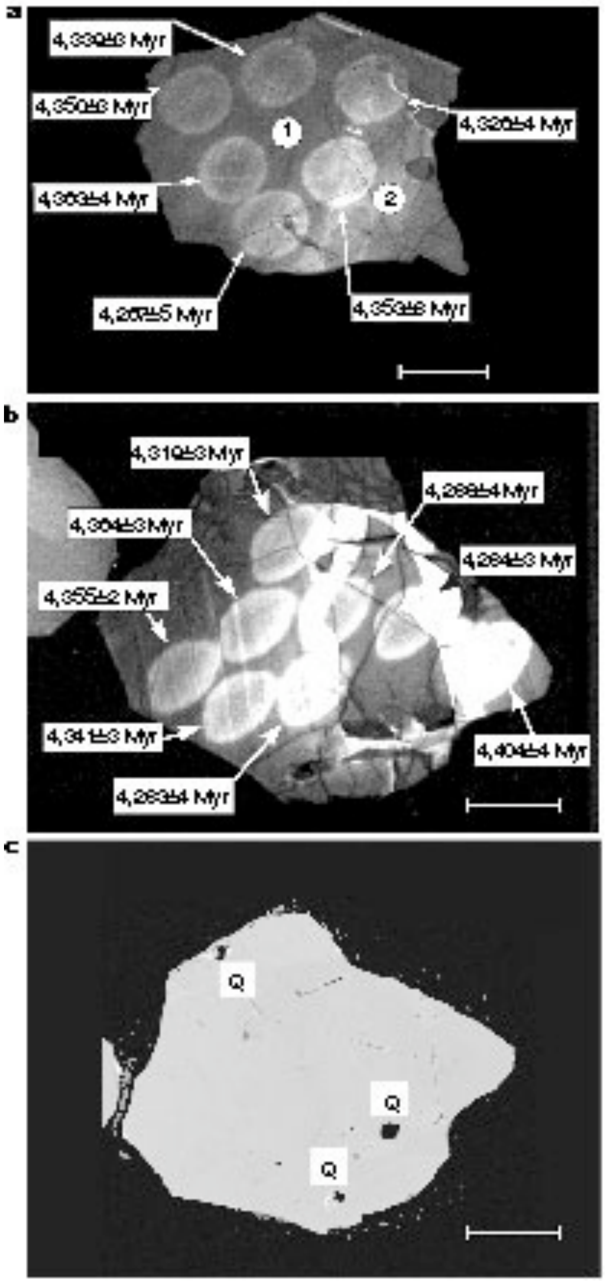
From Valley JW (2005) A cool early Earth? *Scientific American* October 2005: 40-47, IMAGE COURTESY OF DON DIXON, in *Elements* (cover), Aug. 2006.

Early Earth History

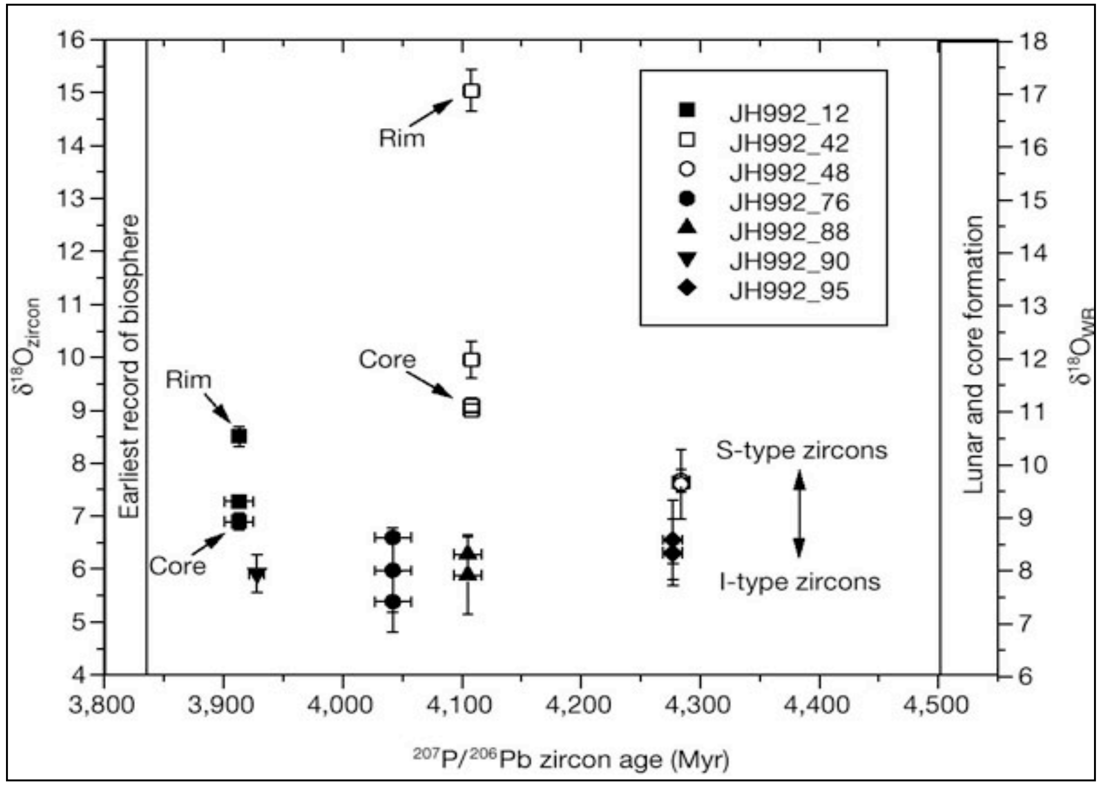



Formation of Earth's Atmosphere and Ocean

Evidence from Zircons for Liquid Water 4.3 Ga



Wilde et al., *Nature* (2001)



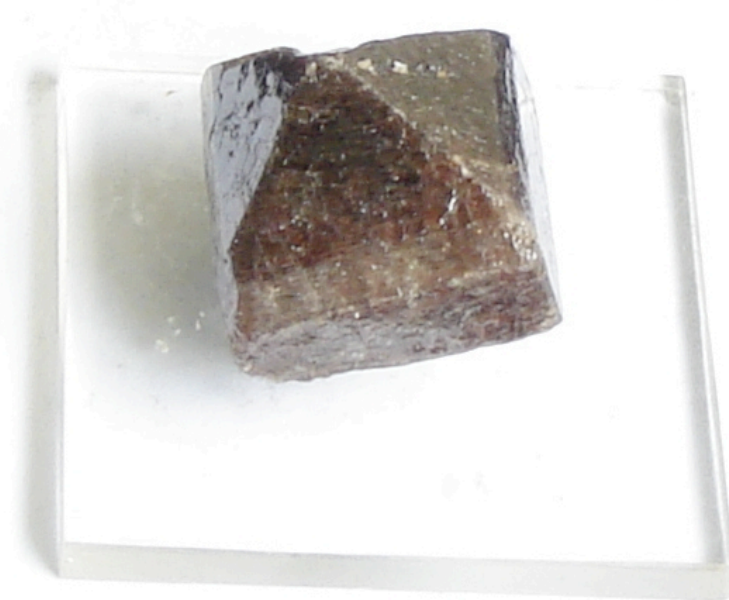
•Heavy oxygen isotope ratios ($^{18}O/^{16}O$) are produced by low-temperature interactions between rock & *liquid* water.

•4.3 Ga zircons have high $^{18}O/^{16}O$, implying the rocks that were melted to form the magma from which the zircons crystallized included material that had been at the surface in the presence of liquid water.

*** Ended here - 10/10/08 ***

-from Wikipedia, 10/10/08

What is Zircon?

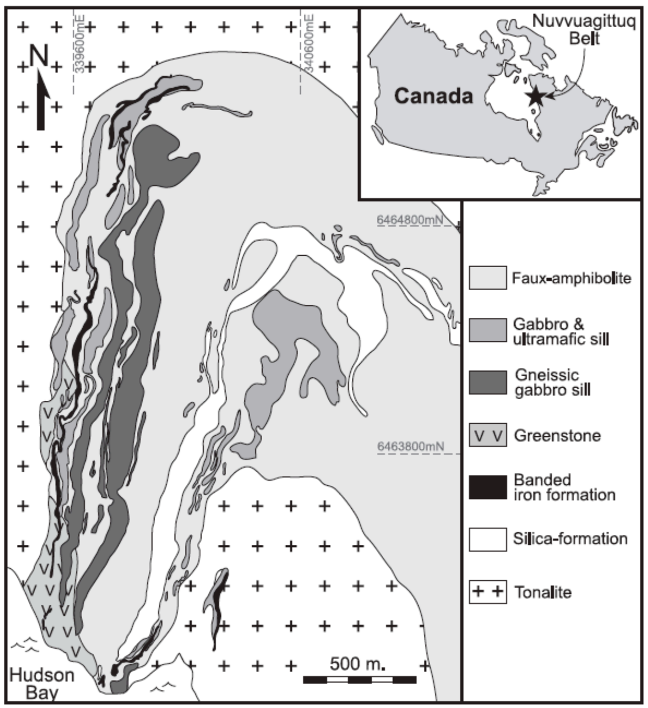


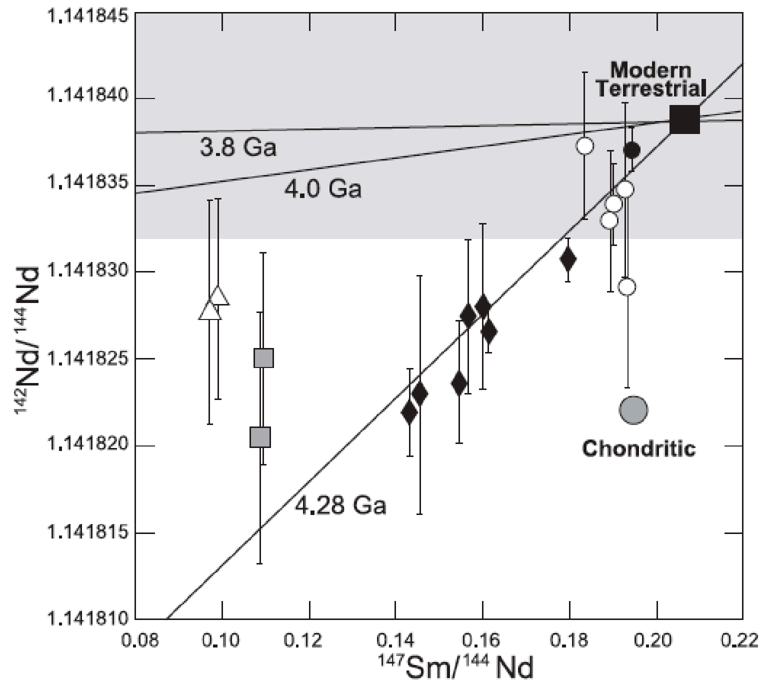
• Zircon is a mineral belonging to the group of nesosilicates. Its chemical name is zirconium silicate and its corresponding chemical formula is $ZrSiO_4$.

-from Wikipedia, 10/10/08

Oldest Continental Crust (yet found): **4.28 Ga**

- Based on Samarium-Neodymium dating
- Measure ^{142}Nd, produced by radioactive decay of ^{146}Sm: ^{146}Sm-->^{142}Nd, $t_{1/2}=1.08\times10^8$ yr
- ^{146}Sm has been extinct since ~4.2 Ga
- Rocks that formed while ^{146}Sm was still around have larger than usual quantities of ^{142}Nd





- Initial solar system ^{146}Sm/^{144}Nd = 0.008 at 4567 Ma
- ^{146}Sm/^{144}Nd = 0.00116 from isochron diagram
- Corresponds to an age of 4280^{+53}_{-81} Ma

^{144}Nd ~ invariant, $t_{1/2}$ = 2.29 x 10^{15} yr (α, ^{140}Ce)
^{147}Sm --> ^{143}Nd, $t_{1/2}$ = 1.06 x 10^{11} yr (α)
^{146}Sm (extinct nuclide) --> ^{142}Nd, $t_{1/2}$ = 1.08 x 10^8 yr (α)

O'Neil et al. (2008) *Science*, Vol. 321: 1828-1831 (Sept. 26, 2008)

Theories for Origin of Earth's Volatile Components: Atmosphere & Oceans

- Arrived with the planetesimals, partly survived the accretion process and outgassed during volcanic activity (Hogbom 1894, Rubey 1951-5). Volcanic gases vary in composition; not primordial and may have been recycled many times. No record of the time and conclusive answers about this scenario (Turekian, 1972; Delsemme, 1997).
- Arrived with comets during the late bombardment - late veneer hypothesis (Delsemme, 1997)
- Arrived with one or more hydrated planetesimals from the outer asteroid belt (Morbidelli, 2001)
- **Arrived with meteorites (seems likely….)**

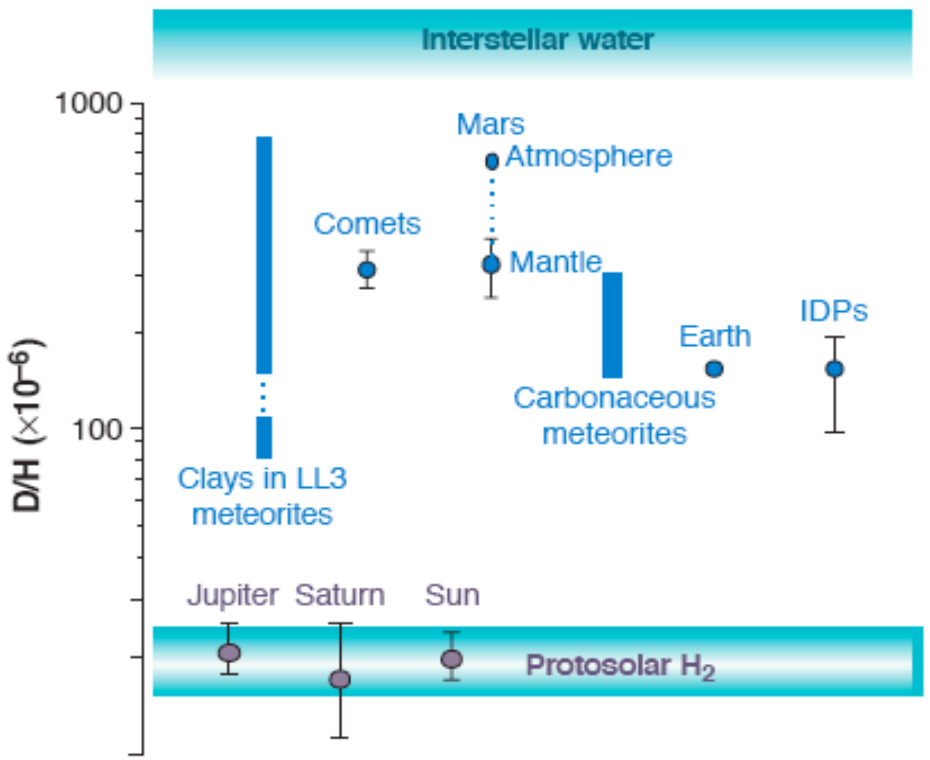
Modeled Composition of Volatiles in Comet Halley

• Comets seem like a plausible source for Earth's Volatile Veneer (Atmosphere/Ocean)

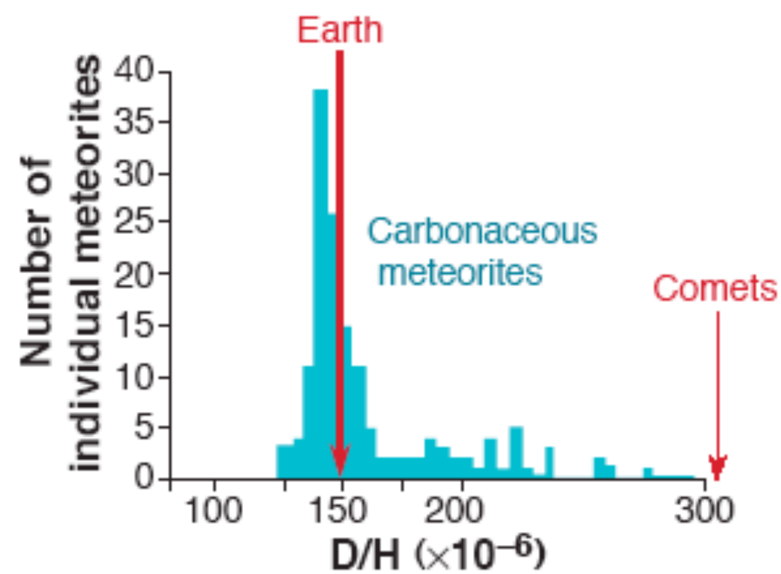
78.5 % H_2O	**2.6% N_2**	**1.5% C_2H_4**	**0.1% H_2S**
4.0% H_2CO	**0.8% NH_3**	**0.5% CH_4**	**0.05% S2**
4.5% HCO-OH	**1.0% HCN**	**0.2% C_3H_2**	**0.05% CS_2**
1.5% CO	**0.8% N_2H_4**		
	0.4% $C_4H_4N_2$		
92% with O	**5.6% with N**	**2.6% H/C**	**0.2% S**

… however, other evidence points to meteorites

D/H Evidence for Origin of Earth's Water from Meteorites



An isotopic enigma. Distribution of the hydrogen isotopic composition in solar system bodies. Blue, water; purple, molecular hydrogen.

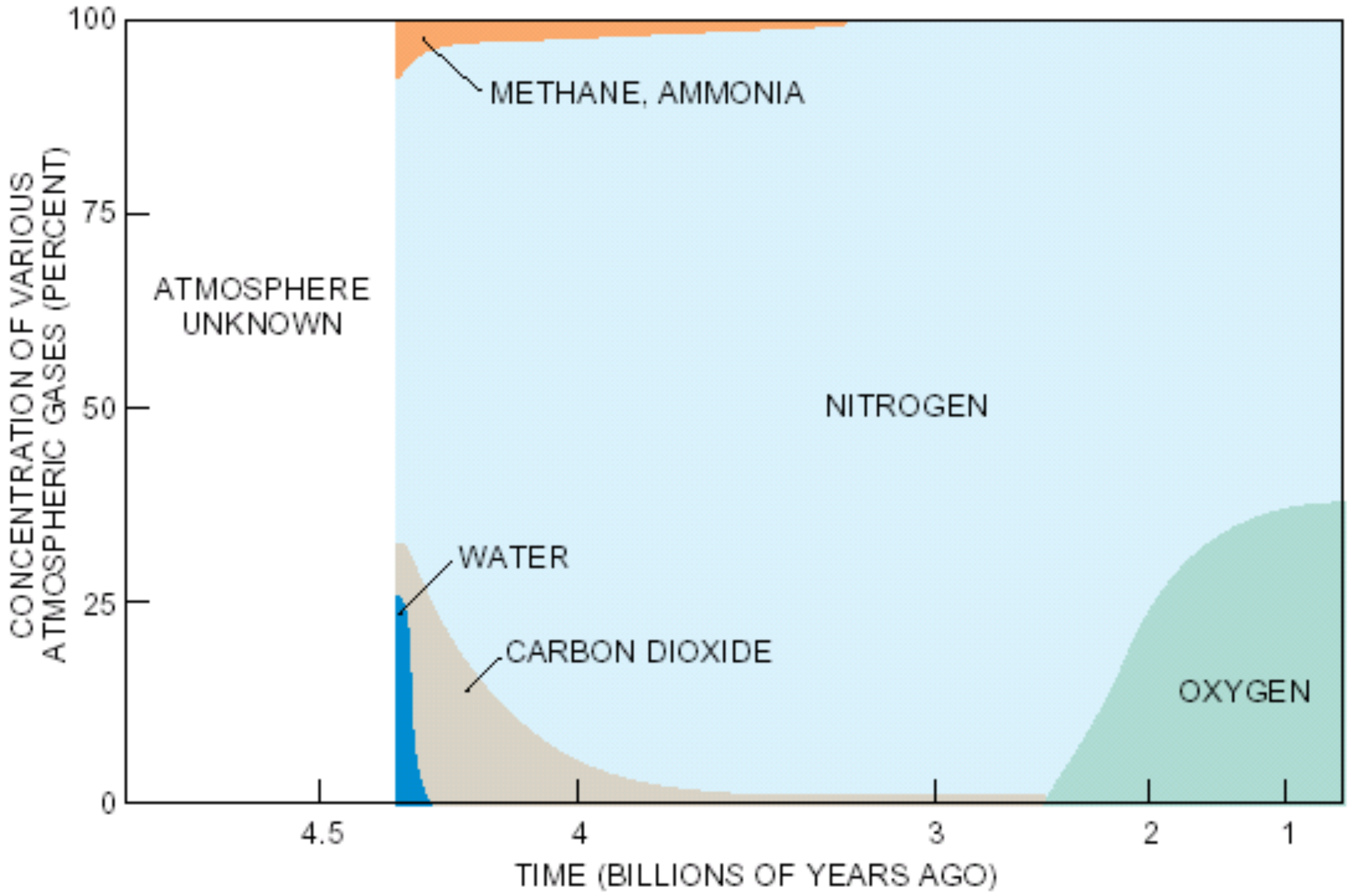


Water from meteors. Distribution of the hydrogen isotopic ratio in carbonaceous meteorites compared with Earth and comets. According to this distribution, water on Earth seems mostly derived from a meteoritic source.

- **Planets formed from collisional accretion of many primitive planets (10-1000 km diam) w/ unstable orbits around Sun**
- **Addition of water-rich bodies during accretion contributed small fraction of water but most added by a few late giant impactors.**
- **Late impactors had D/H ratios similar to carbonaceous meteorites because they originated from the same cold region of the asteroid belt.**

Robert (2001) *Science* Vol. 293: 1056-1058

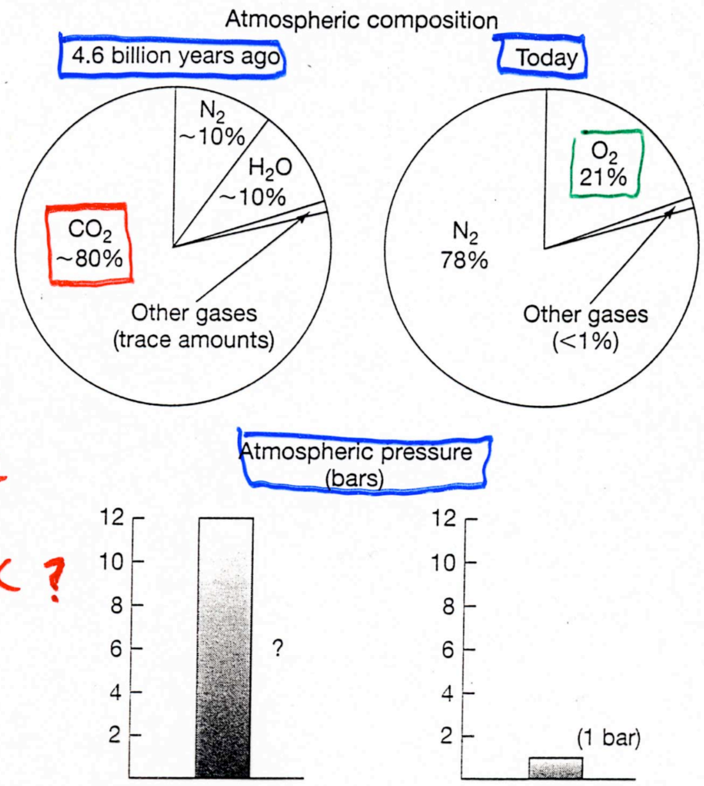
Composition of Earth's Early Atmosphere



ATMOSPHERIC COMPOSITION, shown by the relative concentration of various gases, has been greatly influenced by life on the earth. The early atmosphere had fairly high concentrations of water and carbon dioxide and, some experts believe, methane, ammonia and nitrogen. After the emergence of living organisms, the oxygen that is so vital to our survival became more plentiful. Today carbon dioxide, methane and water exist only in trace amounts in the atmosphere.

Allegre & Schneider (1994)

Earth's Early Atmosphere



Extreme Greenhouse Effect

T_s ~80°-90°C ?

FIGURE 8-7

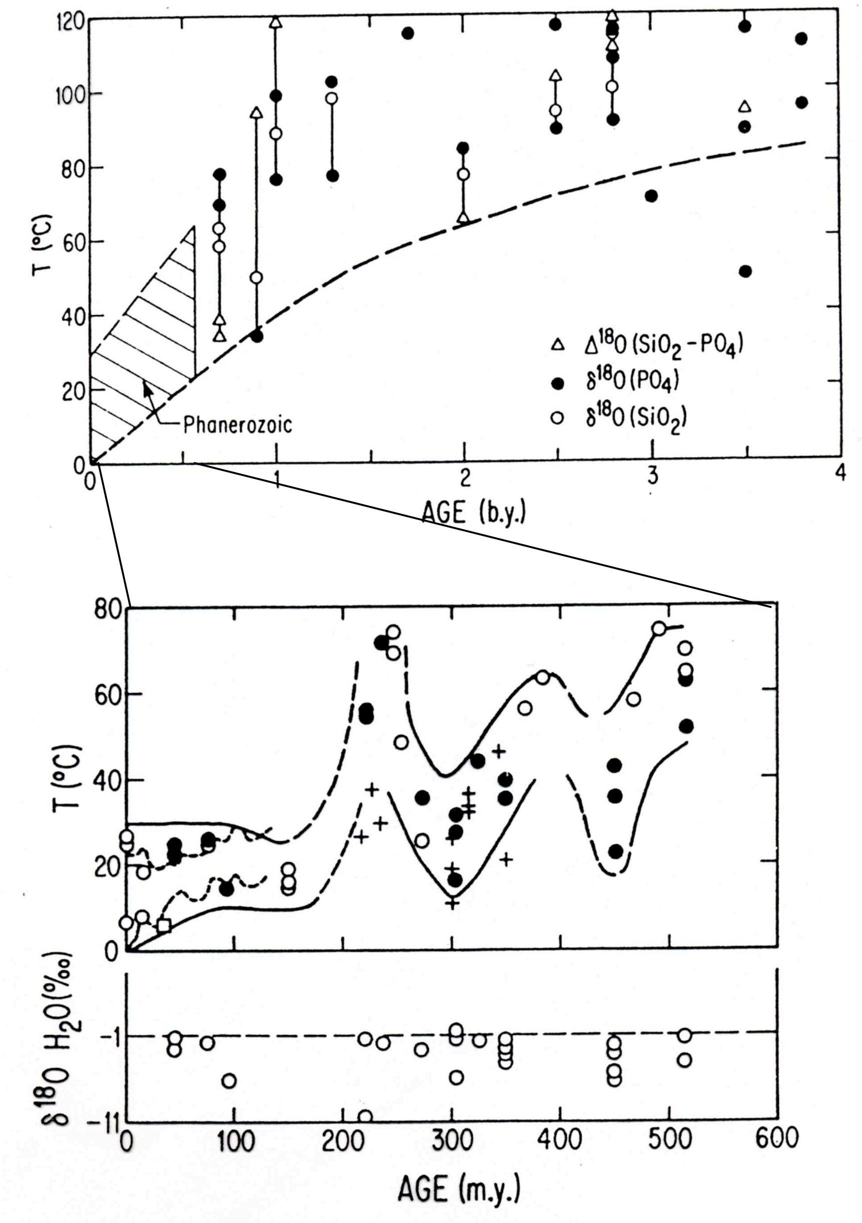
Possible atmospheric composition and pressure during the heavy bombardment period, compared with today.

Changes in Atmospheric Composition over Time

	Prebiotic Atmosphere	*Archean Atmosphere*	*Modern Atmosphere*
Surface pressure	1–10 bars	1–2 bars	1 bar
Constant → N_2	10–80%	50–80%	78%
O_2	about 0	about 0	21%
CO_2	30–90%	10–20%	0.036%
CH_4	10–100 ppm	1000–10,000 ppm	1.6 ppm
CO	100–1000 ppm		0.1–0.2 ppm
H_2	100–1000 ppm		0.5 ppm

Kump et al. (1999) The Earth System, Prentice Hall, New Jersey, 351 p.

Evidence for High Surface T on early Earth



- Oxygen isotope ratios in cherts & phosphates indicate surface temperatures may have exceeded 60°C for most of the period 4-1 Gyr ago

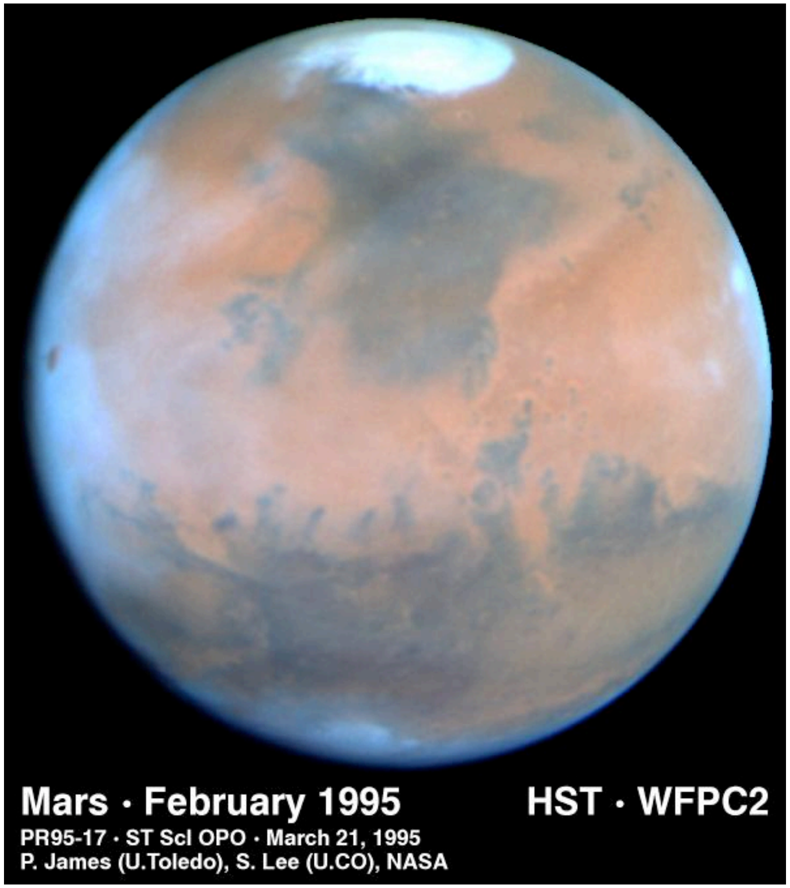
USGS image

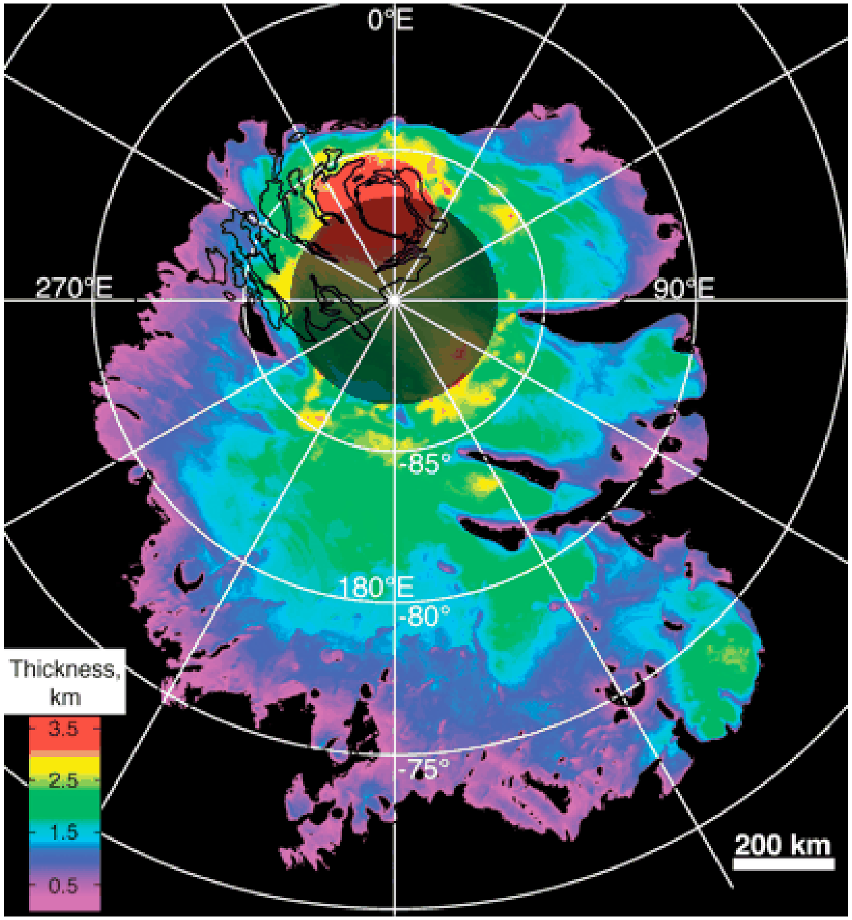
Karhu & Epstein (1986) *Geochimica et Cosmochimica Acta* Vol. 50: 1745-1756

Water Elsewhere in Solar System:

Water Ice on Mars

http://photojournal.jpl.nasa.gov/

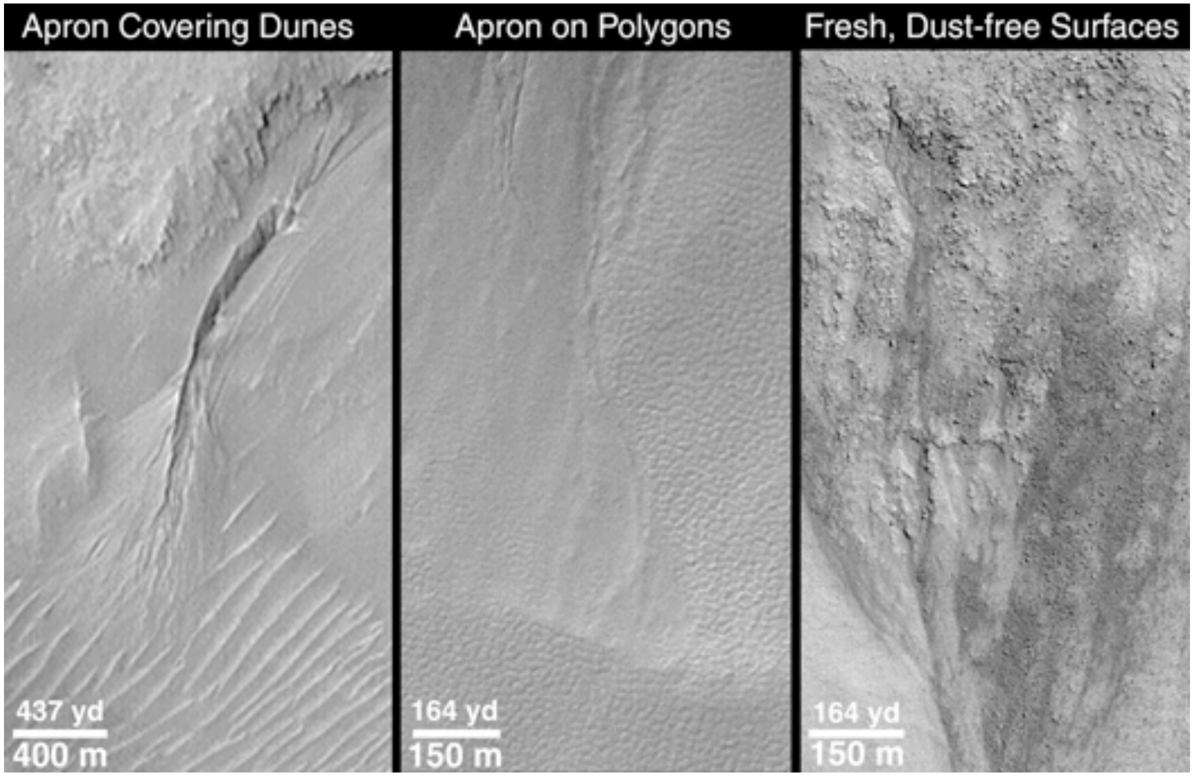




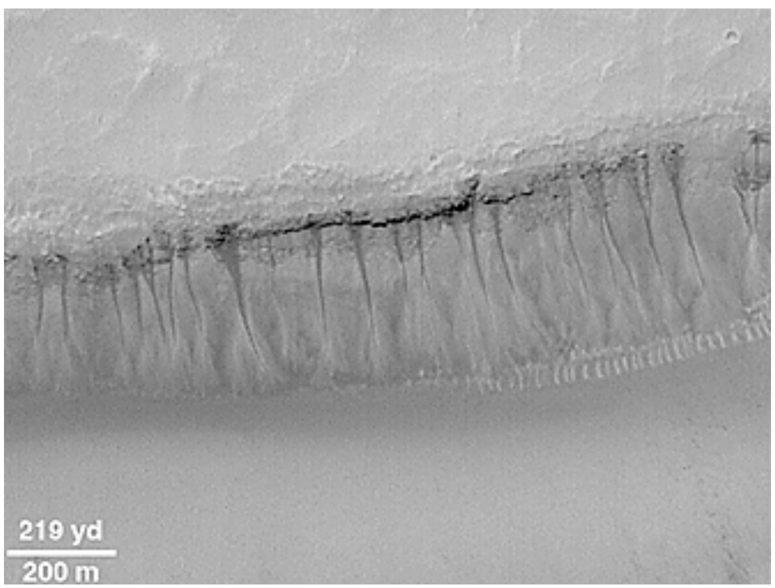
Plaut et al. (2007) *Science*
Vol. 316: 92-95.

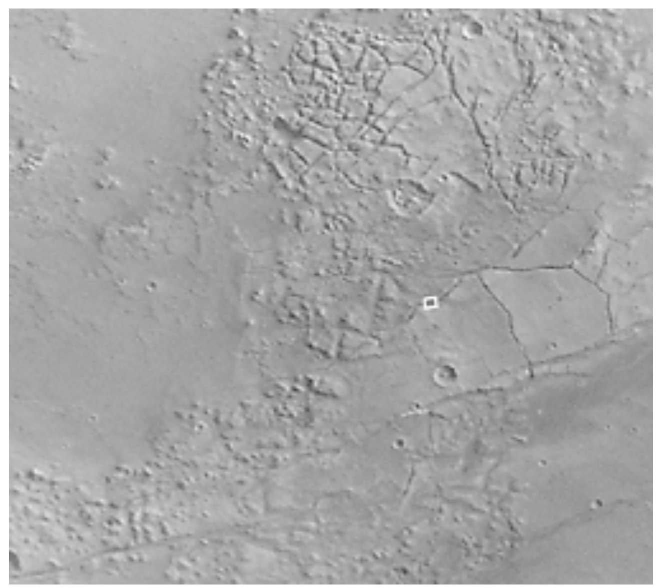
• South Pole water ice thickness: The total volume is estimated to be 1.6 x 10^6 cubic kilometers, which is equivalent to a global water layer approximately 11 meters thick.

Evidence of Recent Water flow on Mars

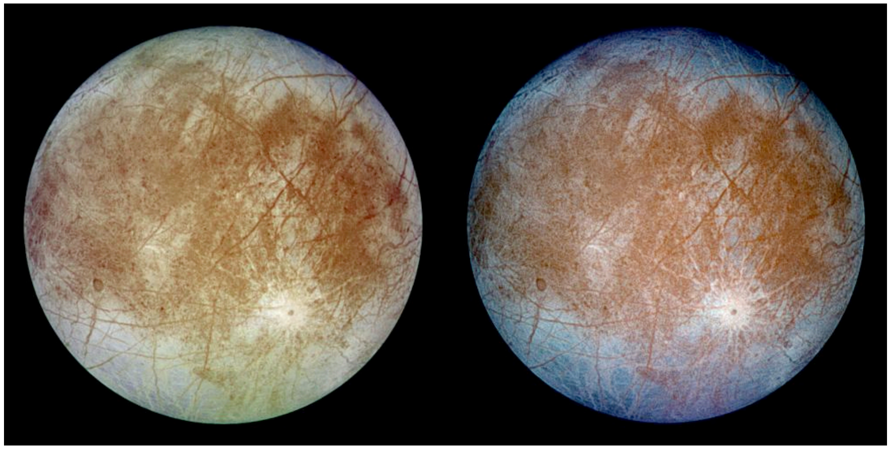


• Martian gullies proposed to have formed by seepage & runoff of liquid water in recent martian times





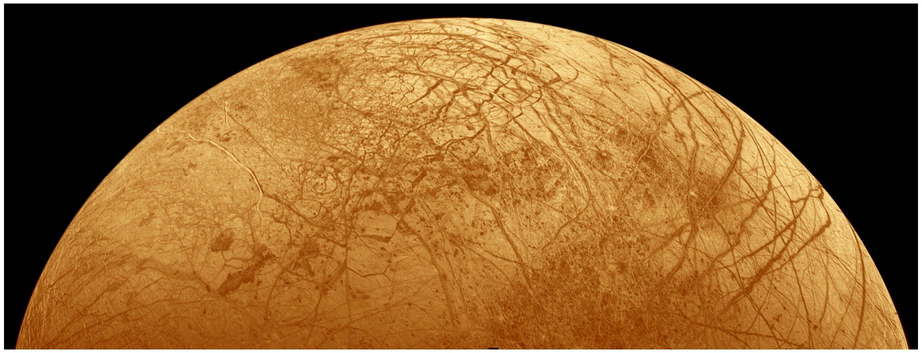
http://www.msss.com/mars_images/moc/june2000/age/index.html



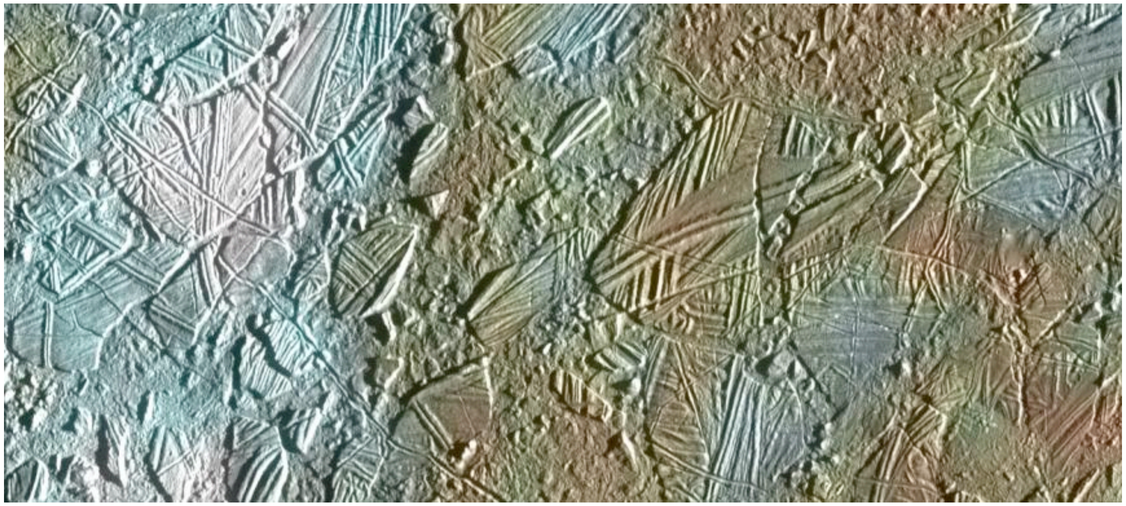
Water on Europa

- One of Jupiter's 4 large (Galilean) satellites
- 25% of Earth's radius

- Crust composed of water & ice



- Fragmented chunks of water ice on Europa's surface



http://science.nasa.gov/newhome/headlines/ast09sep99_1.htm; http://www.solarviews.com/eng/europa.htm