

When the wind has been blowing for a long time and a steady state has been set up, the surface rises at a small uniform elevation rate  $\alpha$  and there is a pressure gradient that induces a recirculation cell in the lake. This ensures that the mass flow rate through any cross section is zero  $\Rightarrow$  mass is conserved and cannot accumulate anywhere in steady-state.

CONSERVATION OF MOMENTUM (N-S EQ.)

X-direction

$$\cancel{\rho \frac{\partial v_x}{\partial t}} + \cancel{v_x \frac{\partial v_x}{\partial x}} + \cancel{v_y \frac{\partial v_x}{\partial y}} + \cancel{v_z \frac{\partial v_x}{\partial z}} = -\frac{\partial P}{\partial x} + \mu \left( \cancel{\frac{\partial^2 v_x}{\partial x^2}} + \cancel{\frac{\partial^2 v_x}{\partial y^2}} + \cancel{\frac{\partial^2 v_x}{\partial z^2}} \right)$$

Labels:   
 $\cancel{\rho \frac{\partial v_x}{\partial t}}$ : STEADY   
 $\cancel{v_x \frac{\partial v_x}{\partial x}}$ : FULLY DEVELOPED   
 $\cancel{v_y \frac{\partial v_x}{\partial y}}$ : negligible   
 $\cancel{v_z \frac{\partial v_x}{\partial z}}$ : negligible   
 $\mu \left( \dots \right)$ : FULLY DEVELOPED

Y-direction

$$\cancel{\rho \frac{\partial v_y}{\partial t}} + \cancel{v_x \frac{\partial v_y}{\partial x}} + \cancel{v_y \frac{\partial v_y}{\partial y}} + \cancel{v_z \frac{\partial v_y}{\partial z}} = -\frac{\partial P}{\partial y} + \mu \left( \cancel{\frac{\partial^2 v_y}{\partial x^2}} + \cancel{\frac{\partial^2 v_y}{\partial y^2}} + \cancel{\frac{\partial^2 v_y}{\partial z^2}} \right) - \rho g$$

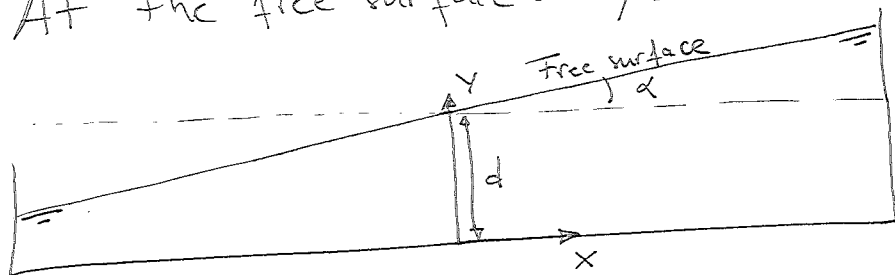
Labels:   
 $\cancel{\rho \frac{\partial v_y}{\partial t}}$ : negligible   
 $\cancel{v_x \frac{\partial v_y}{\partial x}}$ : negligible   
 $\cancel{v_y \frac{\partial v_y}{\partial y}}$ : negligible   
 $\cancel{v_z \frac{\partial v_y}{\partial z}}$ : negligible   
 $\mu \left( \dots \right)$ : negligible

$$\frac{\partial P}{\partial y} = -\rho g \Rightarrow P(x,y) = P_0(x) - \rho g y$$

At the free surface:  $y = d + \alpha x \Rightarrow P_{atm} = P_0(x) - \rho g(d + \alpha x)$

$$P_0(x) = P_{atm} + \rho g(d + \alpha x)$$

$$\frac{\partial P}{\partial x} = \frac{\partial P_0}{\partial x} = \rho g \alpha$$



From the x-direction equation:  $0 = -\rho g \alpha + \mu \frac{d^2 v_x}{dy^2}$

$$\frac{dv_x}{dy} = \frac{\rho \alpha}{\mu} y + C_1 \Rightarrow \mu \left. \frac{dv_x}{dy} \right|_{y=d} = S' = \rho g \alpha d + C_1 \mu$$

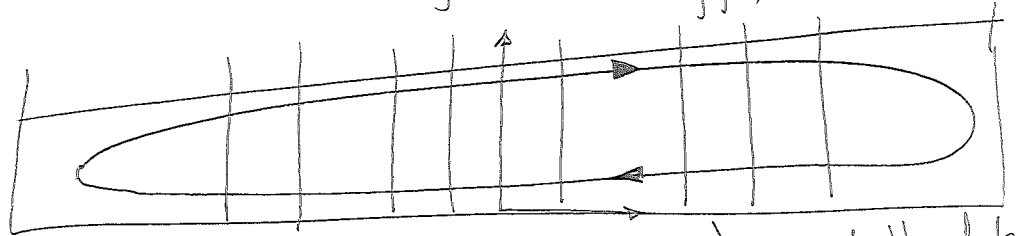
$$C_1 = \frac{S'}{\mu} - \frac{\rho \alpha}{\mu} d$$

$$v_x = \frac{g\alpha}{\nu} \frac{y^2}{2} + \left( \frac{S'}{\mu} - \frac{g\alpha}{\nu} d \right) y + C_2$$

$$v_x(y=0) = 0 = C_2$$

$$v_x(y) = \frac{g\alpha}{\nu} \frac{y^2}{2} + \left( \frac{S'}{\mu} - \frac{g\alpha}{\nu} d \right) y$$

But  $\alpha$  and  $S'$  are not independent, the elevation angle of the free surface is a function of the shear rate imposed by the wind. To link the together we apply conservation of mass:



Along all these cross sections of the lake, away from the ends, the flow is quasi-unidirectional but mass has to be conserved and that means the net flux across a section has to be zero:

$$\int_{\text{cross-section}} g v_x(y) dA = \int_0^d v_x(y) \cdot \underbrace{b}_{\text{width of the lake perpendicular to the paper}} dy = 0$$

$$\int_0^d \left[ \frac{g\alpha}{\nu} \frac{y^2}{2} + \left( \frac{S'}{\mu} - \frac{g\alpha}{\nu} d \right) y \right] dy = \frac{g\alpha}{\nu} \frac{y^3}{6} + \left( \frac{S'}{\mu} - \frac{g\alpha}{\nu} d \right) \frac{y^2}{2} \Bigg|_0^d =$$

$$\frac{g\alpha}{\nu} \frac{d^3}{6} + \frac{S' d^2}{\mu} - \frac{g\alpha}{\nu} \frac{d^3}{2} = 0$$

$$\frac{S' d^2}{2\mu} = \frac{g\alpha}{\nu} \frac{d^3}{3} \Rightarrow$$

$$\alpha = \frac{3}{2} \frac{S'}{g\alpha d}$$