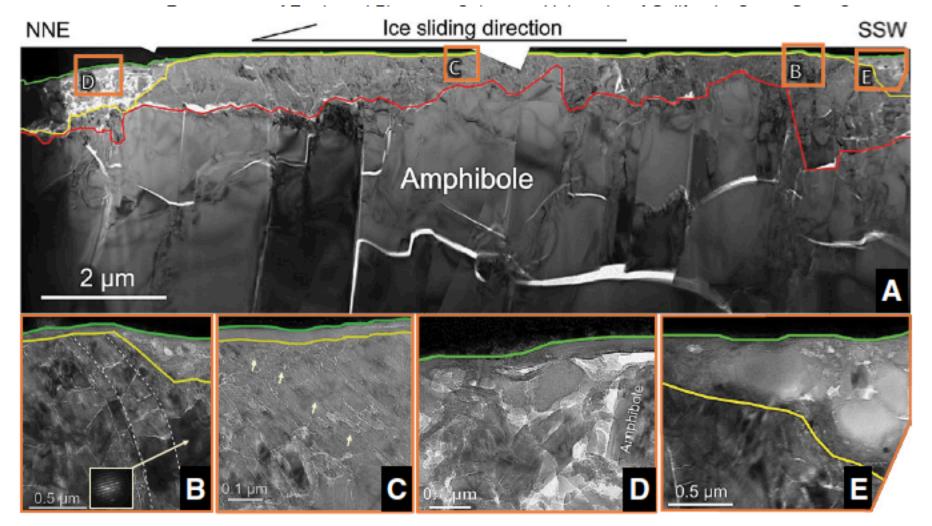


Physico-Chemical case-hardening: nano-scale particles filling micro-cracks and sealing the surface (2017, Geology)

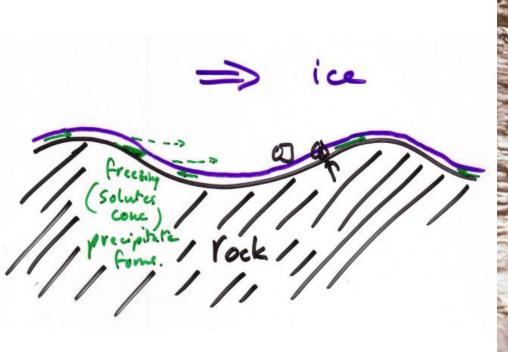
The coating layer of glacial polish

Shalev Siman-Tov¹, Greg M. Stock², Emily E. Brodsky¹, and Joseph C. White³



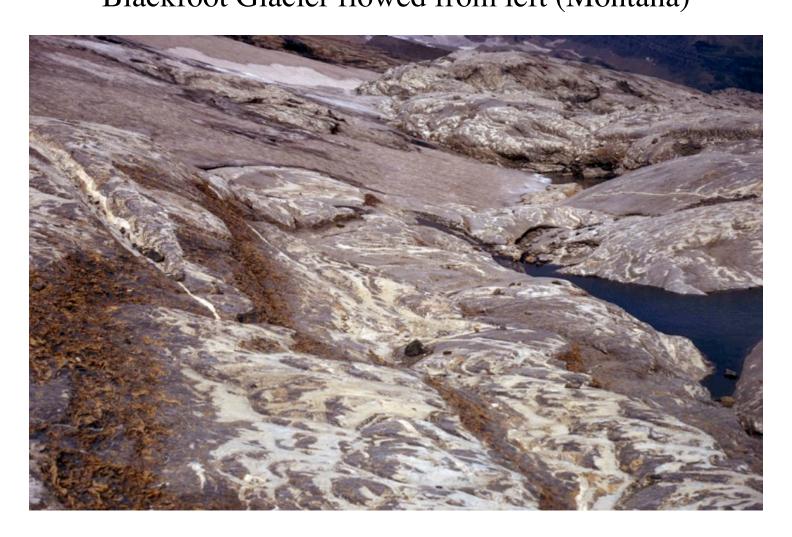
Sliding physics (regelation) & suglacial chemical processes

Sliding over small bumps is dominated by regelation, which involves melting/freezing, and water flow in a thin basal film. Solutes in the water film that are rejected during the freezing process can exceed saturation, causing chemical precipitation.



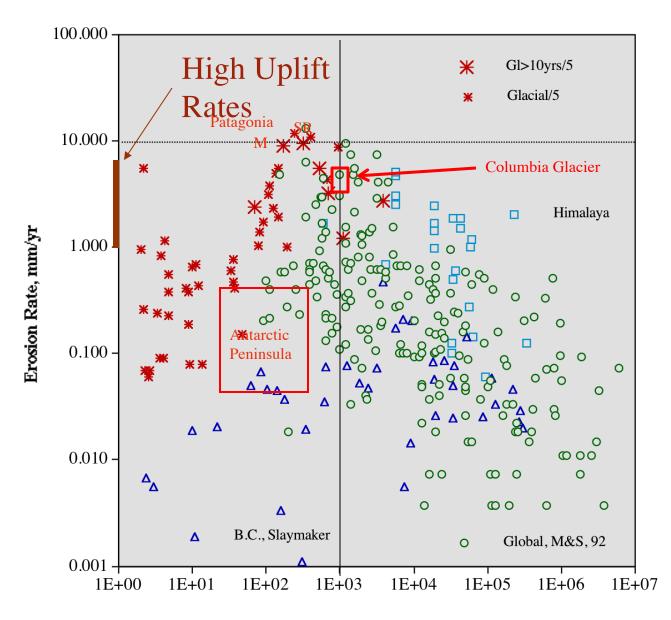


Contrasting hydraulic systems: microscopic, saturated films, where the white carbonate coatings cover the bedrock, traversed by interconnected cavities full of under-saturated water as evident from the dissolution features (buff-colored argillite seams), Blackfoot Glacier flowed from left (Montana)



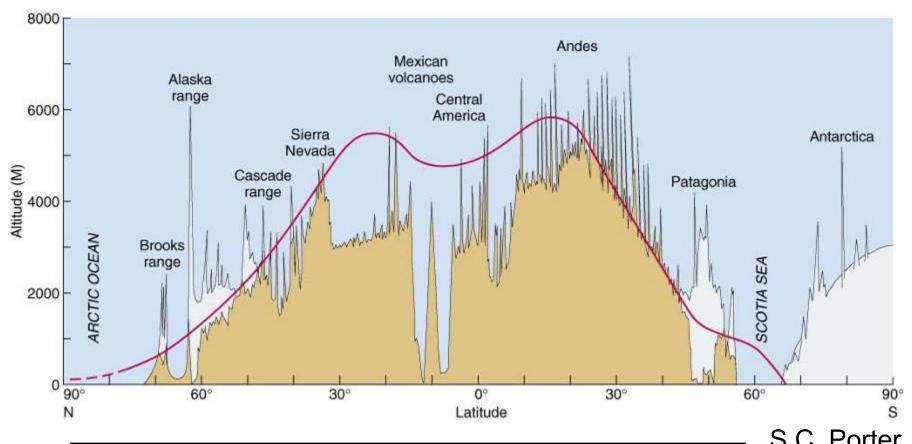
Glacial & Non-Glacial Denudation RatesRates from Alaskan Tide-water glacier reduced 5-fold

Global erosion rates



Basin Area, km²

Glacial Buzz saw



Equilibrium Line Altitude "ELA"

S.C. Porter

Near Polar Regions: Alpine topography at sea level, Lofoten Islands, Norway



Near the Equator, the World's Highest Mountains, Mt Everest in center





The ridge lower with the valleys. 2015 landslide, Taan Fjord below Mt. St. Elias, Alaska. The mountain fell, 150 million tons of rock shooting down...displacing water that surged uphill as a tsunami, reaching ~600 feet above"

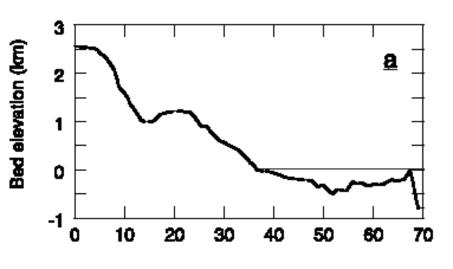
Sediments, the product of glacial erosion, and associated landforms: braided rivers, Gilbert deltas and turquoise waters, Peyto Lake, Banff NP, Alberta

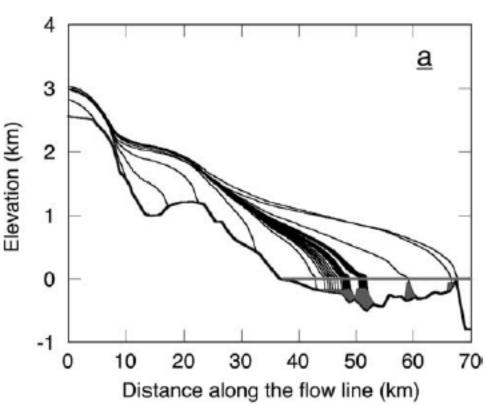


Controls on advance of tidewater glaciers: Results from numerical modeling applied to Columbia Glacier

F. M. Nick, C. J. van der Veen and J. Oerlemans

(JGR, 112, F03S24, doi:10.1029/2006JF000551, 2007)



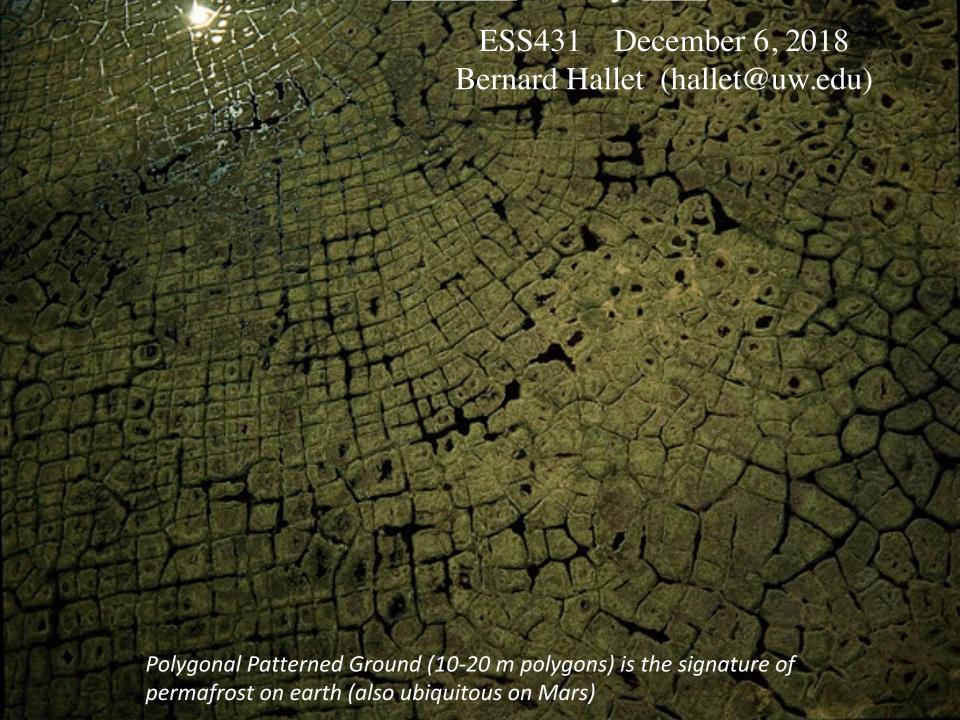


"The results suggest that irrespective of the calving criterion and accumulation rate in the catchment area, it is impossible for the glacier terminus to advance into deeper water (>300 m water depth) unless sedimentation at the glacier front is included."

Summary

- Insights into glacial erosion mechanisms: quarrying, abrasion, chemical processes
- Introduction to ideas, models, and data: model validation is sparse
- Products of erosion: from scratches to spectacular valleys & fjords; massive amounts of sediment
- Rates of erosion: unsurpassed but highly variable
- Diverse roles of glacial erosion: from modulating glacial motion & fluctuations, to shaping much of the terrain we live on

Thank you!



Permafrost: brief intro

Thermal regime & the warming Arctic

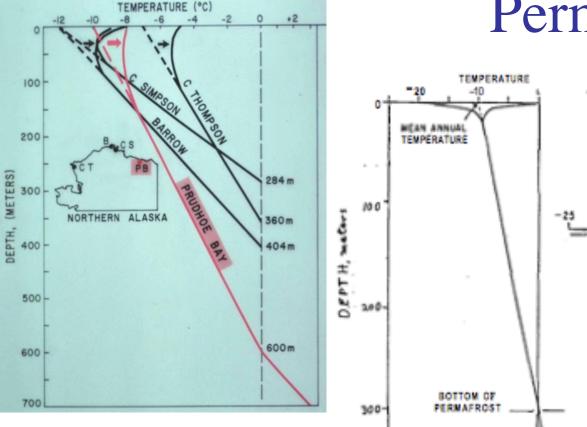
Soil is well known to expand, heave, and crack as it freezes but why and why it matters?

Patterned ground

Pingos, rock glaciers, etc

Mars, the permafrost planet

Permafrost thermal regime Permanently frozen



Note that the frozen ground records the ongoing warming over the last century, long before instruments were available in the North

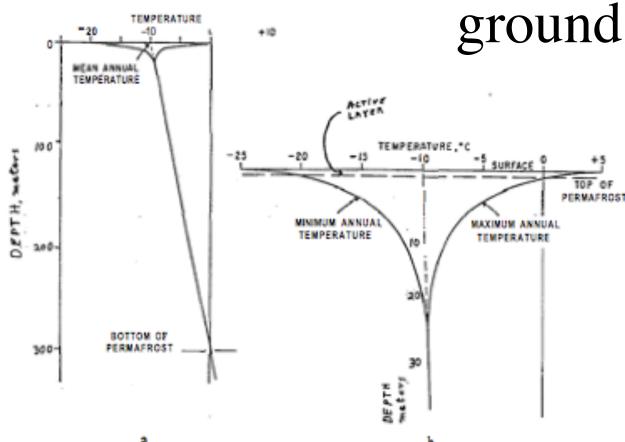
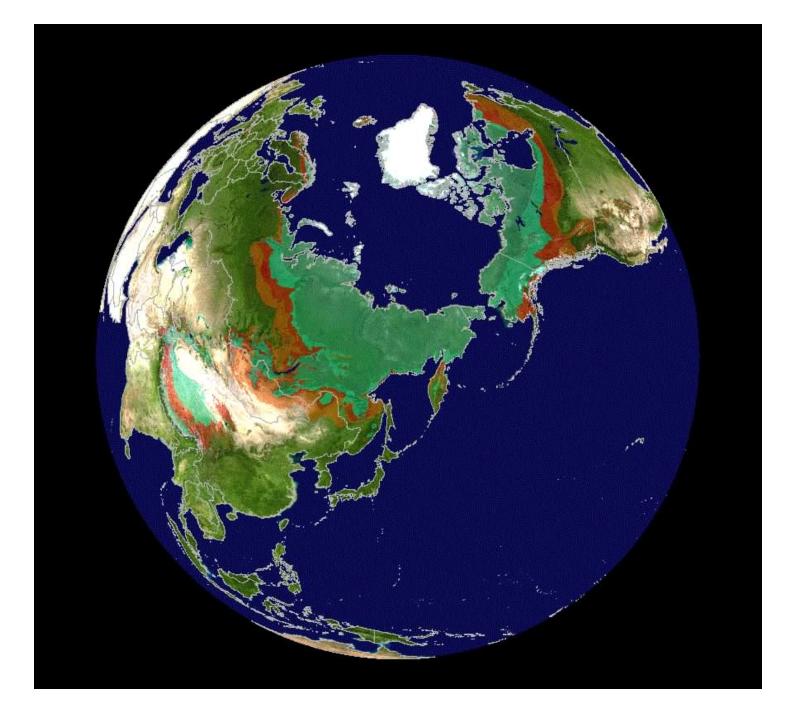
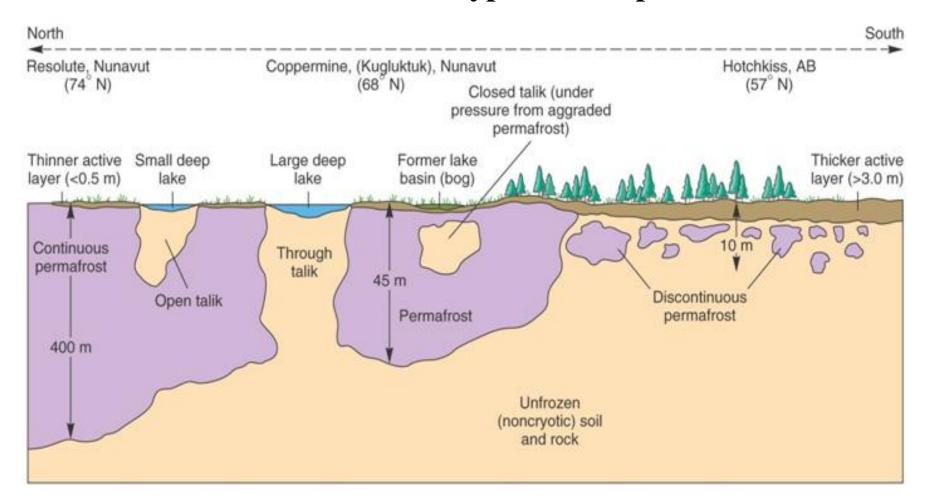


Fig. 2. Relationships between temperature and depth showing how the lower limit (a) and upper limit (b) of permafrost are established. Shown for ground surface temperature with seasonal amplitude of 15°C, mean, T_p, of -10°C, and geothermal gradient of 1°C per 30 m.



Permafrost Types and Depths





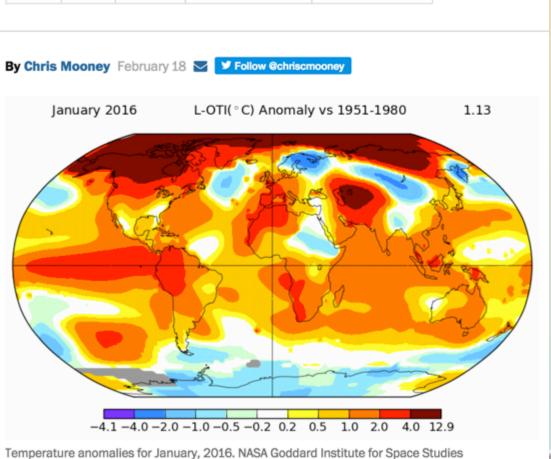
http://wpo.st/nHOH1

Energy and Environment

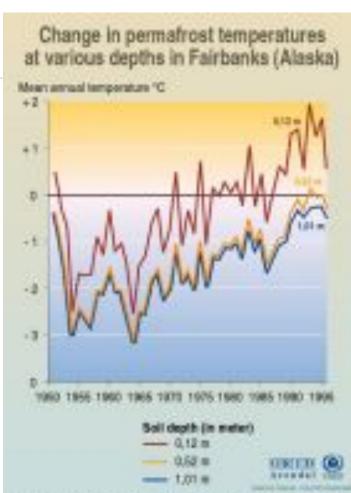
3131

Scientists are floored by what's happening in the Arctic right now

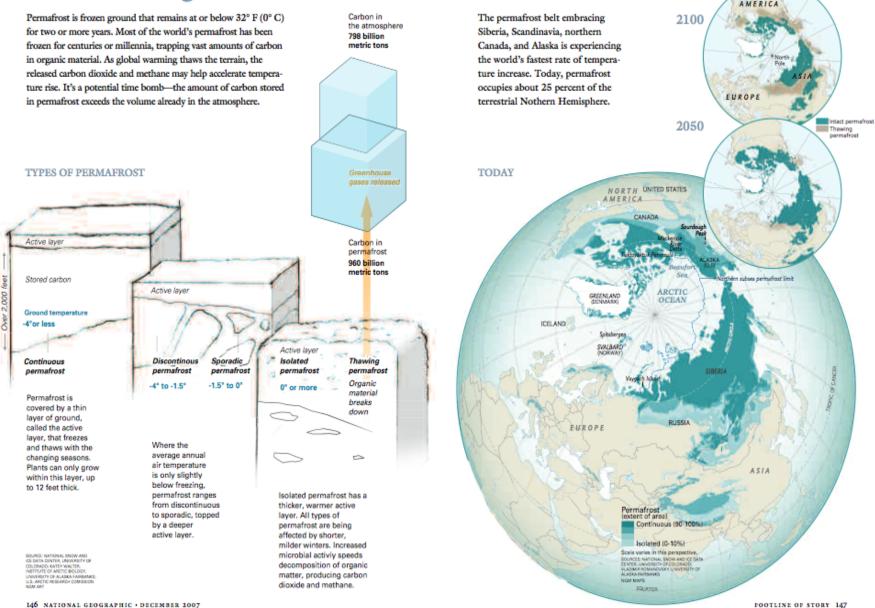
□ Reading List



□ Save for Later



A Threatening Thaw



. .

A Threatening Thaw

Permafrost is frozen ground that remains at or below 32° F (0° C) for two or more years. Most of the world's permafrost has been frozen for centuries or millennia, trapping vast amounts of carbon in organic material. As global warming thaws the terrain, the released carbon dioxide and methane may help accelerate temperature rise. It's a potential time bomb-the amount of carbon stored in permafrost exceeds the volume already in the atmosphere.

in billions of tons Atmosphere: 798

Global Carbon Mass

Permafrost: 960

TYPES OF PERMAFROST

Active layer Stored carbon Active layer Ground temperature -4°or less Active layer Continuous Discontinous Sporadic Isolated permafrost permafrost permafrost permafrost -4° to -1.5° -1.5° to 0° 0° or more Permafrost is covered by a thin layer of ground, called the active layer, that freezes Where the and thaws with the average annual changing seasons. air temperature Plants can only grow

is only slightly below freezing. permafrost ranges from discontinuous to sporadic, topped layer. All types of by a deeper active layer. affected by shorter,

Isolated permafrost has a thicker, warmer active permafrost are being milder winters. Increased microbial activiy speeds decomposition of organic matter, producing carbon dioxide and methane.

Carbon in the atmosphere

798 billion

metric tons

Greenhouse

Carbon in permafrost 960 billion metric tons

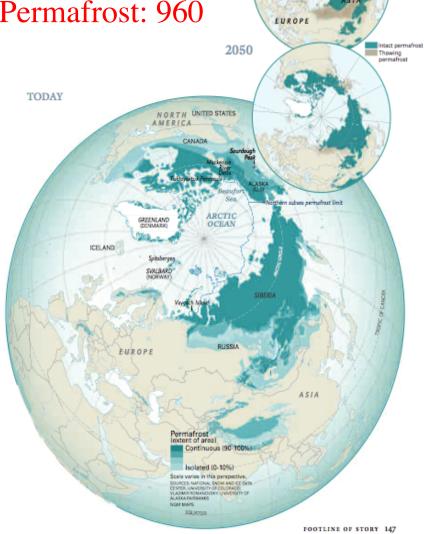
Thawing

material

breaks

down

permafrost Organic



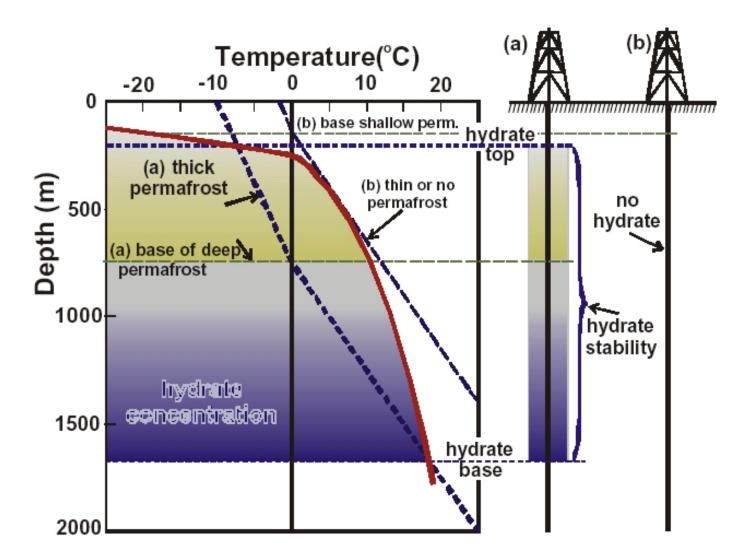
146 NATIONAL GEOGRAPHIC - DECEMBER 2007

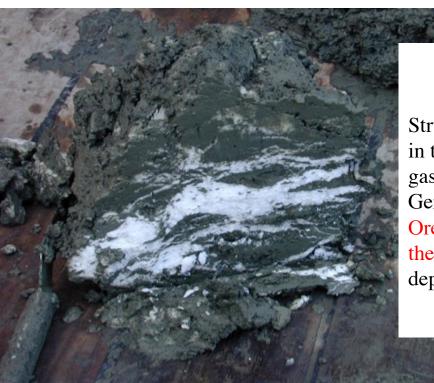
within this layer, up

to 12 feet thick.

COLORADO: KATEY WALTER, INSTITUTE OF ARCTIC BIOLOGY

Perhaps, a bigger concern about warming permafrost: gas hydrates





Gas hydrate-bearing sediment, from the subduction zone off Oregon

Structure of a gas hydrate (methane clathrate) block embedded in the sediment of hydrate ridge, off Oregon, USA. The gashydrates have been found during a reaserch cruise with the German reseach ship FS SONNE in the subduction zone off Oregon in a depth of about 1200 meter in the upper meter of the sediment. The shown gas hydrate (white) has been deposited in thin layers into the sediment



Note: at a depth of 1200 m, ocean water slightly above 0° C is well within the methane hydrate stability field

"Estimates on how much energy is stored in methane hydrates range from 350 years' supply to 3500 years' supply based on current energy consumption. That reflects both the potential as a resource and how little we really know about the resource" http://web.ornl.gov/info/reporter/no16/methane.htm

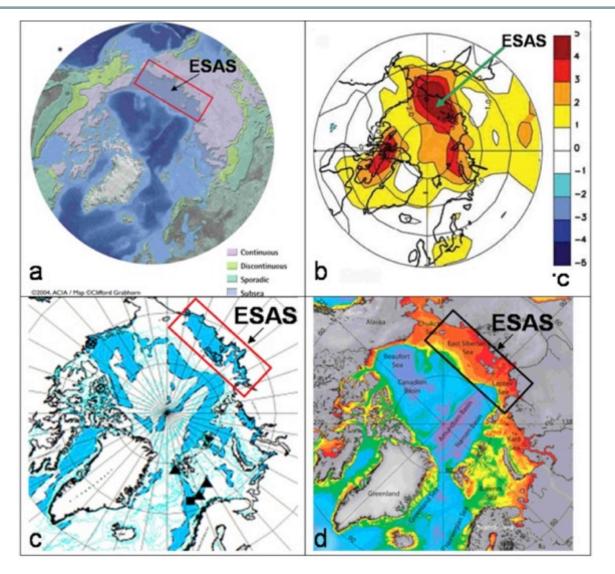
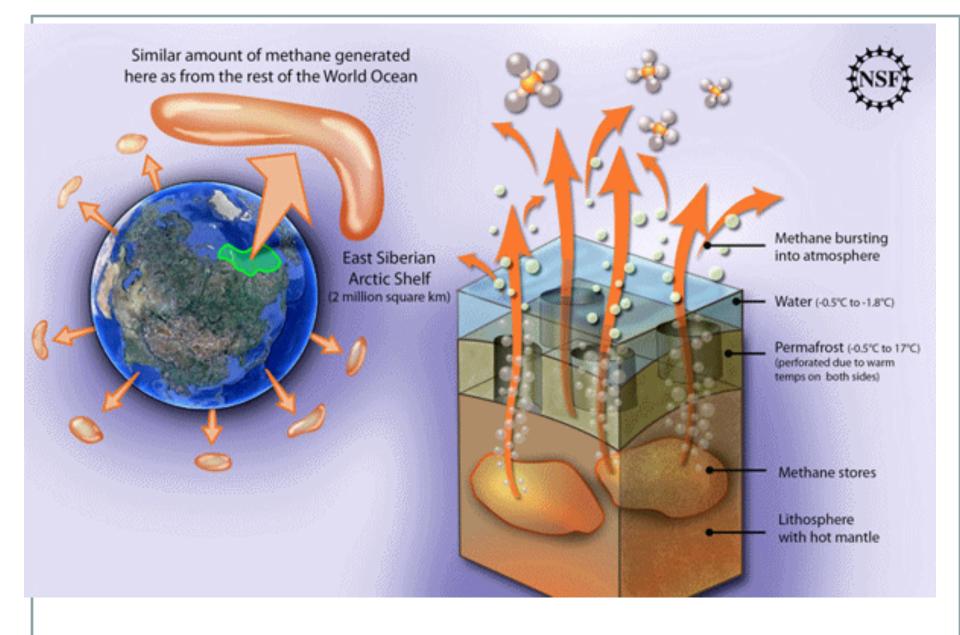


Figure 1. (a) Distribution of Arctic Ocean sub-sea permafrost (shown in purple); (b) springtime surface air temperature anomalies for 2000–5 relative to a 1968–96 base period; (c) predicted deposits of Arctic Ocean CH_4 hydrates (shown in blue); (d) bathymetric map of the Arctic Ocean (depth ≤ 50 m shown in shades of red) http://iopscience.iop.org/1748-9326/7/1/015201/article

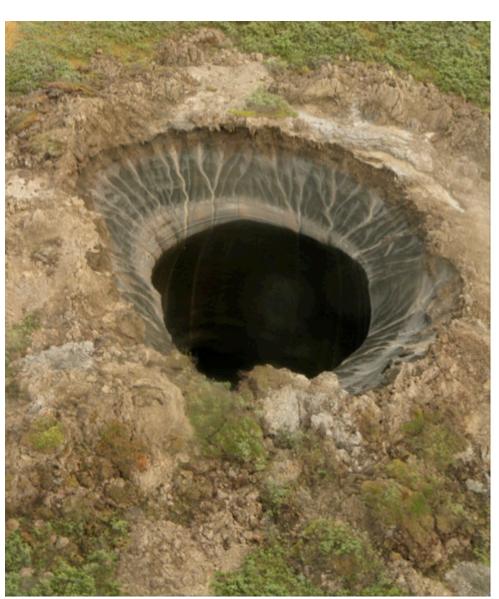


http://www.zeeburgnieuws.nl/nieuws/images/methane_hydrate_nsf.gif









Climate change headlines...

Are Siberia's methane blow-holes the first warning sign of unstoppable climate change?

Need to educate yourself about headlines like this and the underlying system.

Buildings affected by melting permafrost







Environmental Problems of Permafrost

Human activities degrade the permafrost environment

- Thermal erosion: surface layer removed, so the thaw extends deeper into the ground
- Thermokarst: natural surface removed over large areas of tundra leads to ground subsidence that results in depressions and lakes — ground may subside and buildings may collapse
- Buildings must be insulated
- Pipelines must be placed aboveground, for them not to thaw permafrost, or not to freeze

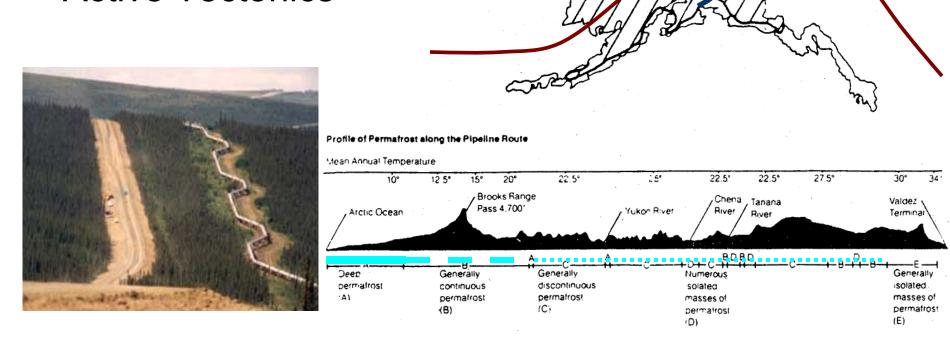
Alaska Pipeline: an engineering miracle instead of an environmental disaster

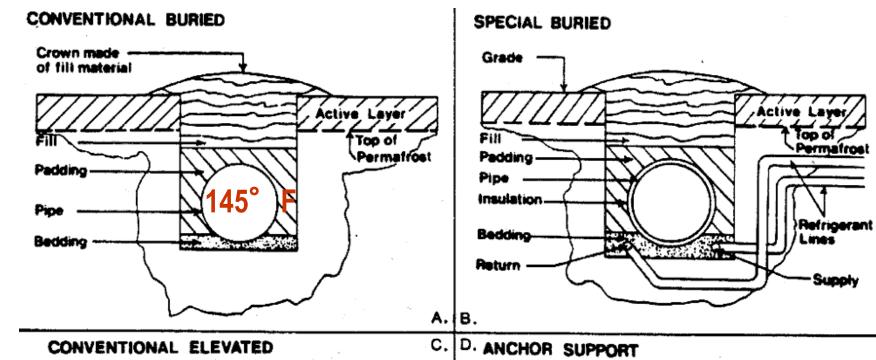
permafrost in Alaska

Discontinuous No Permatrost

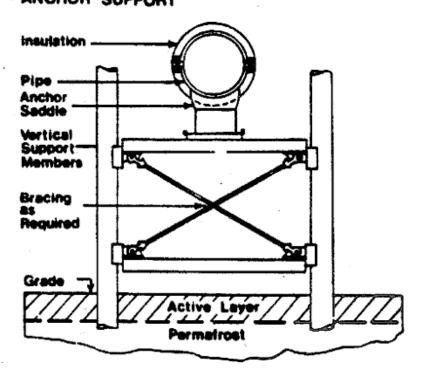
Continuous and discontinuous permafrost regimes

Active Tectonics









Trans-Alaskan Pipeline

Astonishing engineering success instead of environmental catastrophe



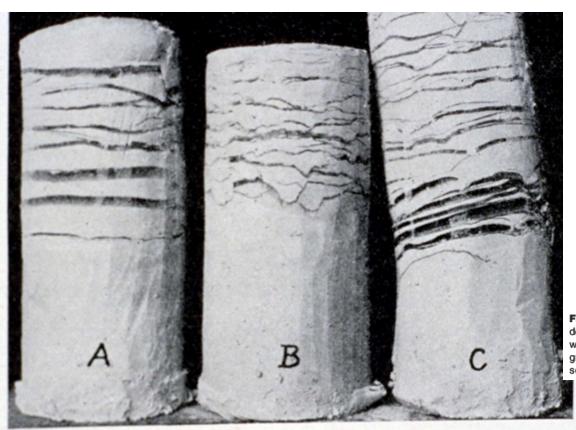
- Oil transport
 - emerges from the ground at 180° F (80° C)
 - flows in pipeline at temperatures above 120
- VSM's (vertical support members)
 - Equipped with heat exchangers to prevent warming of ground material
 - Allow the pipeline to efficiently transfer oil without affecting stability
 - Allows wildlife to maintain migratory patterns

Lots of ice below the surface in the Arctic





Freezing Clay: Ice lenses (Tabor)



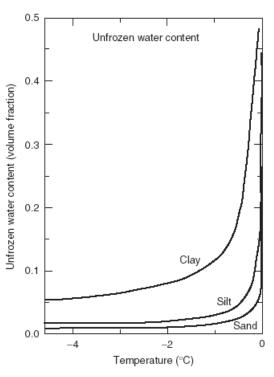
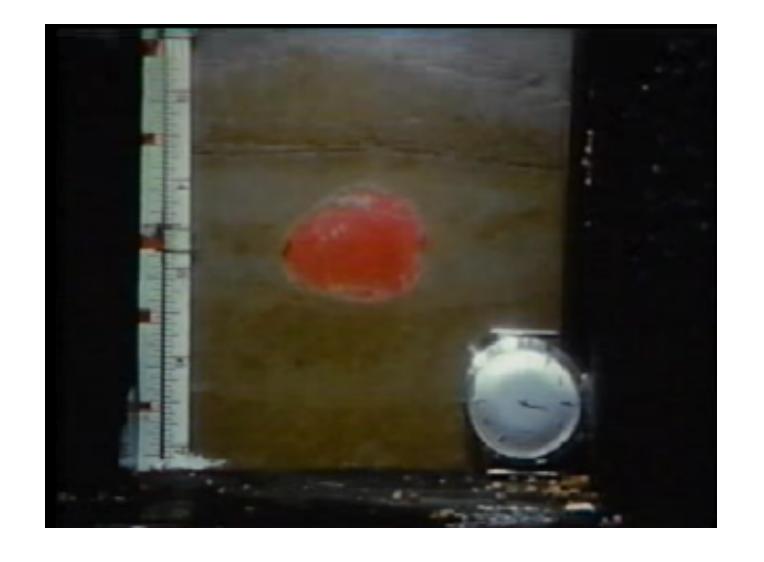
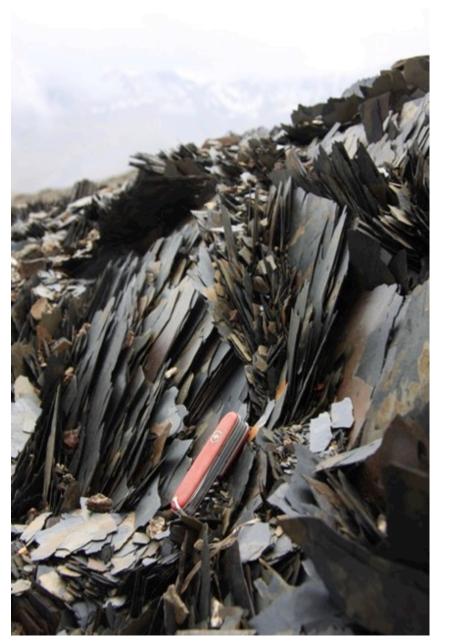


Figure 4 Representative values for the temperature dependence of unfrozen water contents in sand, silt, and clay. Unfrozen water contents typically increase with temperature and finergrained soil and are small in moss and peat. The presence of solutes increases unfrozen water contents.

Fig. 11.—Clay cylinders frozen (A) under no surface load; (B) under iron weight sulated from clay by wooden disc; and (C) under iron weight in contact with clay.



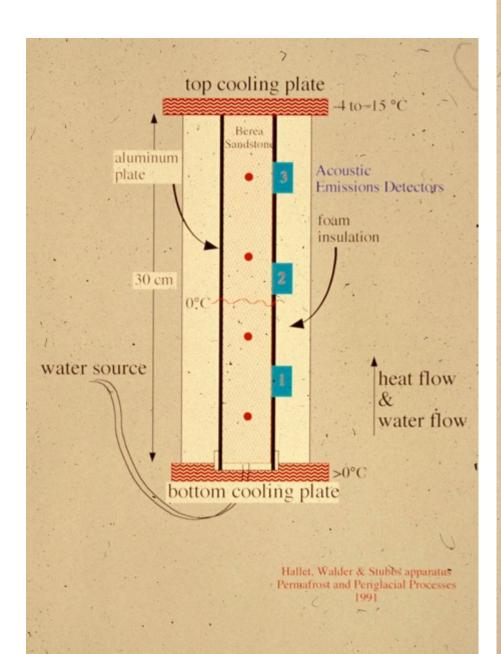
Cold Regions Res. & Eng. Lab (CRREL) movie of ~1-inch rock in freezing silt About 3 days of freezing from the surface down in ~1 minute Kaplar, C.W.,1976. Stone migration by freezing of soil [In:King, C.A.M. (Ed.), Periglacial processes . Stroudsburg, PA, Dowden, Hutchinson & Ross, Inc., pp. 44–45.]



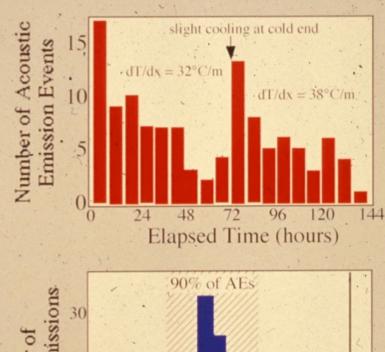
Freezing is also very effective in breaking down rock, but why?

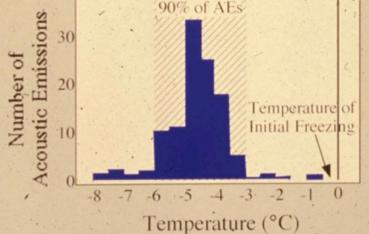


Lab test of theory...



6 Day Sustained Freezing Experiment using Berea Sandstone





from Hallet, Walder & Stubbs
Permafrost and Periglacial Processes 1991



"The strange beauty of ice extruded from the earth on a cold morning contains hidden secrets of nature for those who look beneath their feet."

by David Cavagnaro

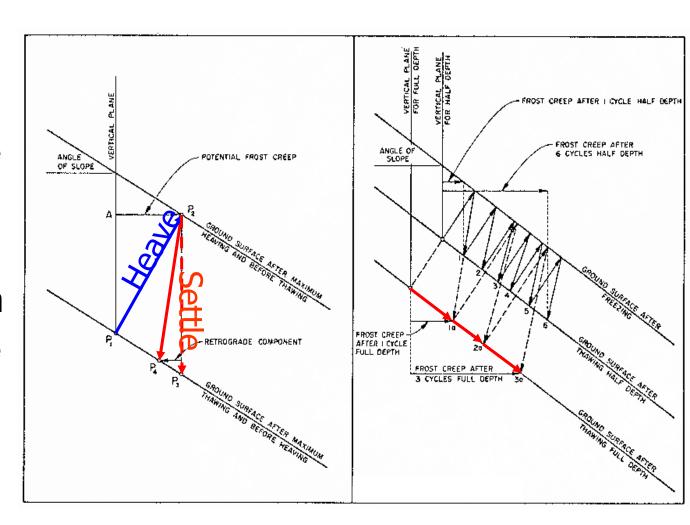
This Living Earth



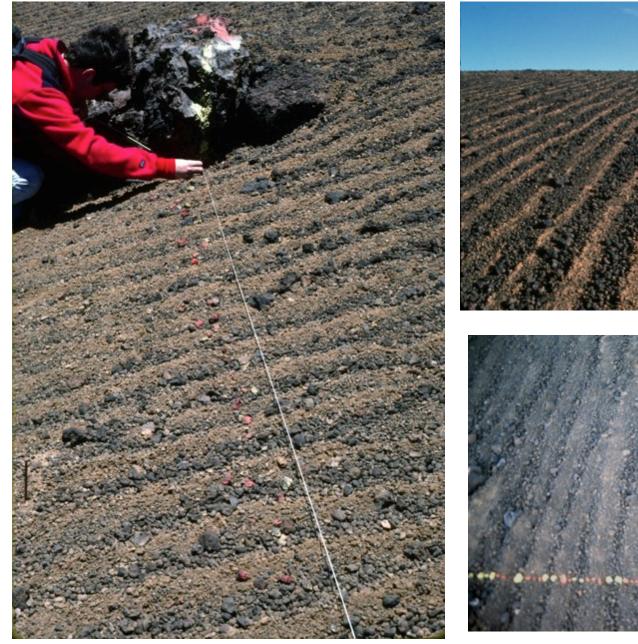
Needle Ice, Sierra Nevada

Frost Heave

- Heave
 - Moisture
 - Grain size
- Creep
 - Saturation
 - Grain size
 - Slope



Subsurface ice growth & melt creates patterns in Hawaii: Sorted stripes, Mauna Kea

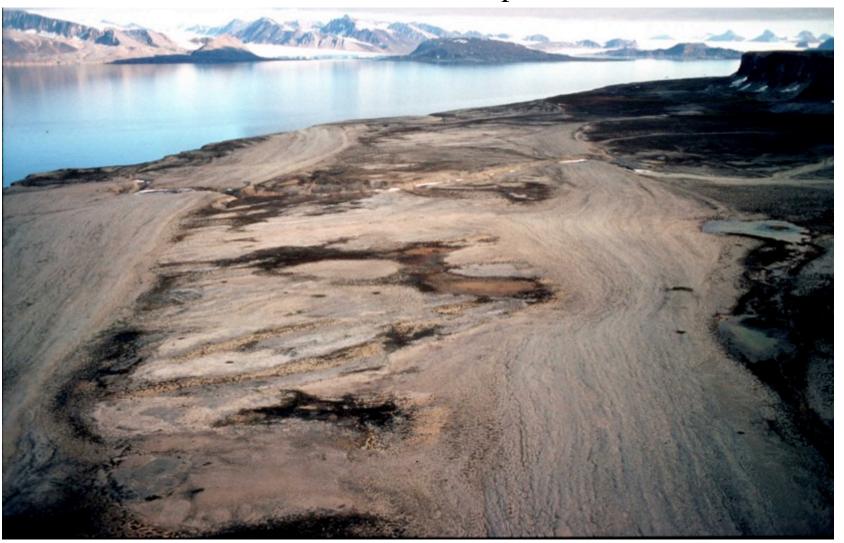






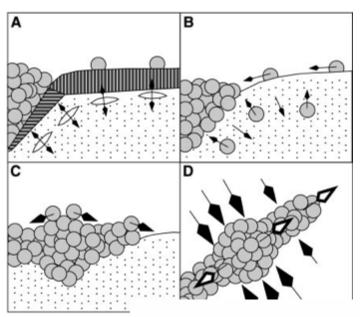
freeze/thaw on raised beaches,

Spitsbergenlift due to reduced ice load



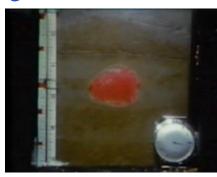
Stone Circles, Spitsbergen

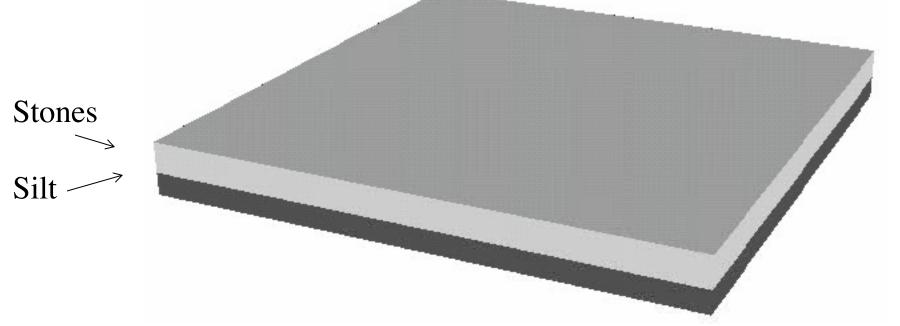




Mark Kessler's Model 4 simple elements

- A. Texture impacts freezing front
- B. Frost sorting of stones
- C. Downslope creep
- D. Lateral pressure

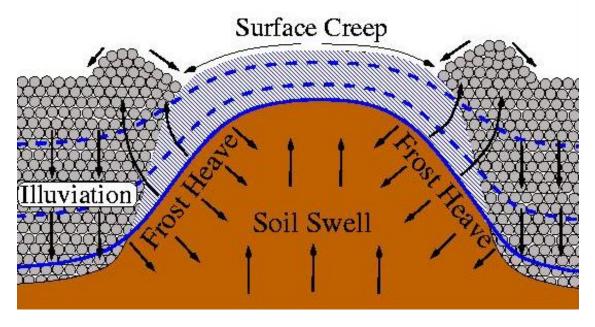




Patterns: hallmark of complexity

Soil diapirs form due to recurrent freeze-thaw cycles

Original state: layer of mixed material of nearly uniform thickness Initial Phase: larger stones/pebbles move to the surface (upfreezing) and then move laterally downslope



Courtesy of Mark Kessler

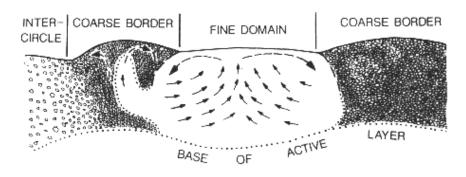
INTER- COARSE BORDER FINE DOMAIN COARSE BORDER

END DOMAIN COARSE BORDER

LAYER

BASE OF ACTIVE

Fig. 5. Movements in sorted circles proposed by Hallet et al. (1988).



Numerical model is consistent with conceptual model based on field observations



Fig. 6. Another possible explanation of the movements observed on the surface by Hallet et al. (1988).

Cryoturbation: broader importance

It impacts cycling of soil carbon and release of carbon from permafrost to the atmosphere

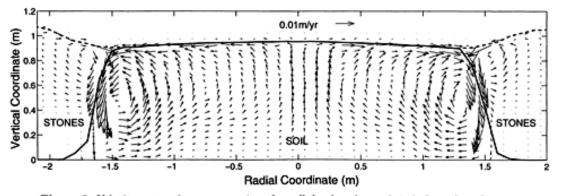
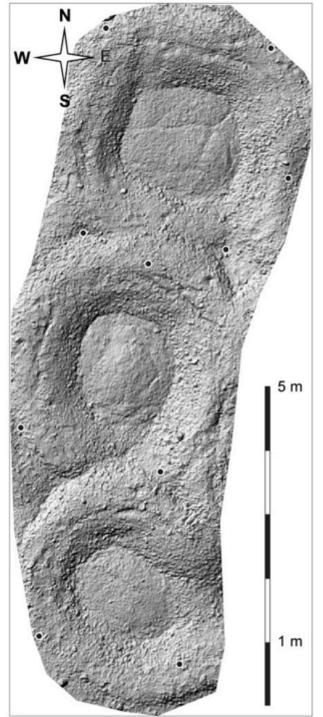
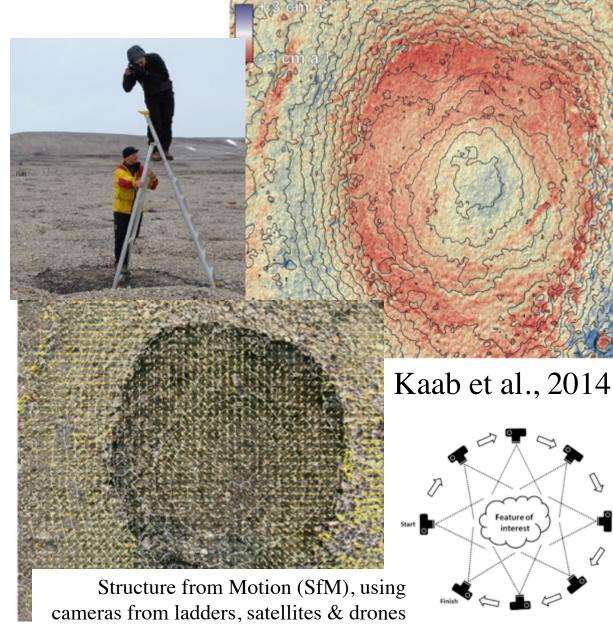


Figure 9. Velocity vectors in a cross section of a well-developed sorted circle from the reference model, as in Figure 8. Solid line is the soil surface; dashed line is the ground surface.

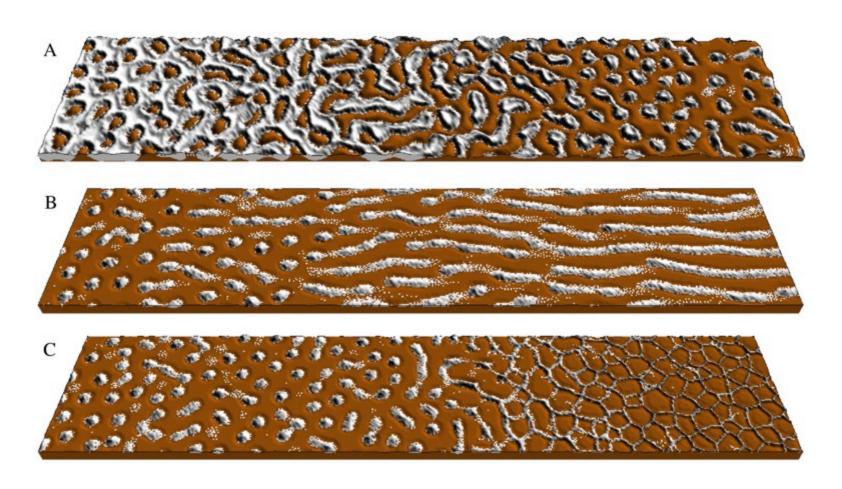


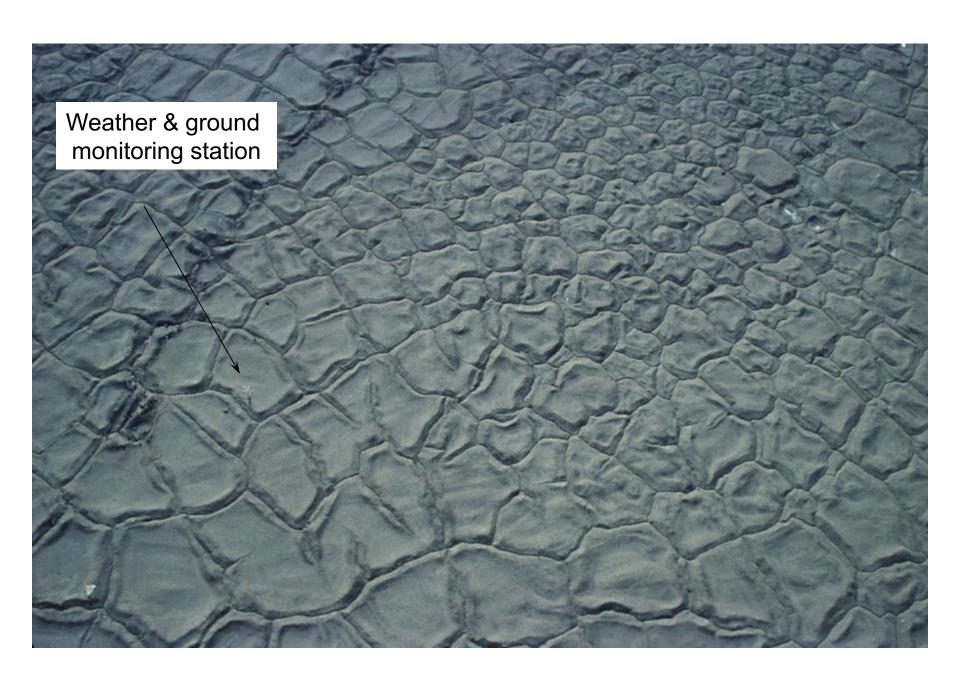
Exciting emerging technology (software &

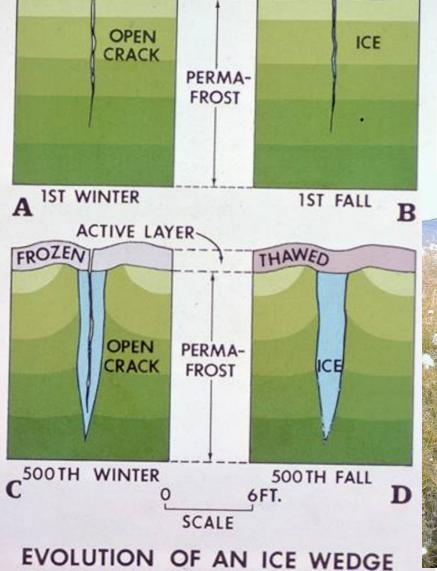
hardware) http://www.the-cryosphere.net/8/1041/2014/tc-8-1041-2014.pdf



Kessler and Werner's model produce rich and realistic diversity of patterns







THAWED

ACTIVE LAYER

FROZEN

Ice-Wedge Polygons Spacing is a small multiple of crack depth



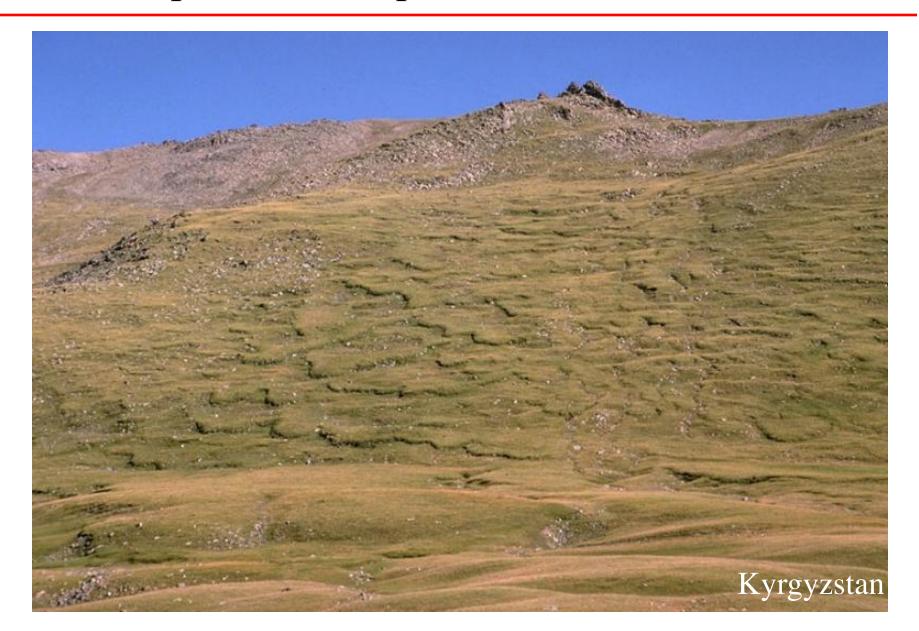
20 m Borehole (19 m into ice) Ground Temperature (C) 1.6 1.4 Crack Movement (cm) Air (2m) Temperature (C) 1/1/06 1/1/99 1/1/00 1/1/02 1/1/03 1/1/04 1/1/05

Summary for Central Beacon Valley; Underlain by massive ice

- •Temperature variations diminish with depth and lag the surface signal
- Seasonal "breathing": Greatest annual contraction/expansion
- No sustained divergence across crack



Rapid Soil Creep (Solifluction) Lobes





Rock Glacier S. Alaska

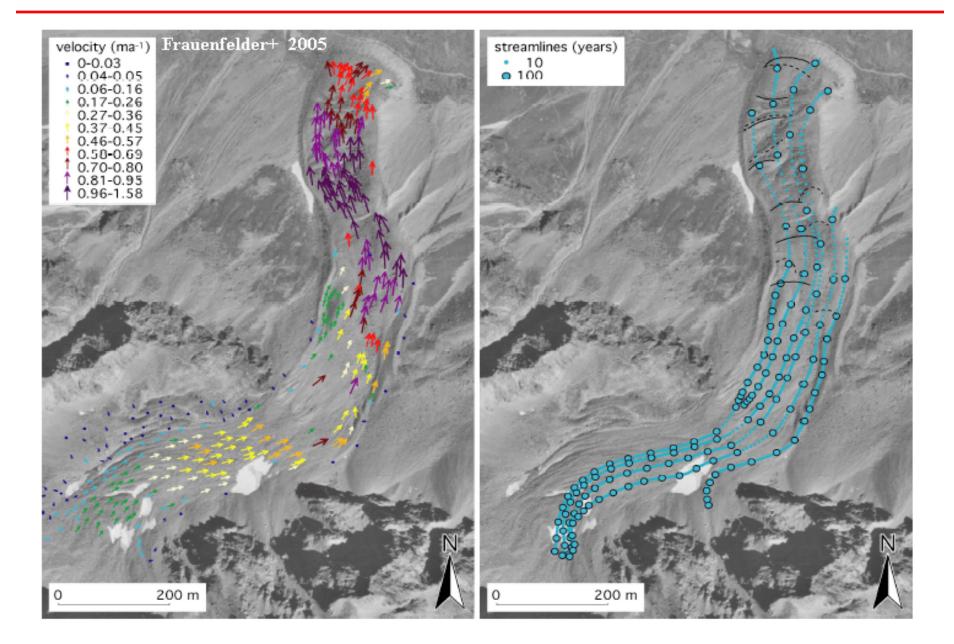
http://ngm.nationalgeographic.com/2007/12/permafrost/edmaler-photography.html

Mountain Permafrost: Glacial & periglacial processes

Ice-sculpted peaks and deep valleys filled with ice-debris mixtures (permafrost) in Peruvian Andes. Glacier perched on debris, and turquoise lakes. Why that color?



Rock Glacier

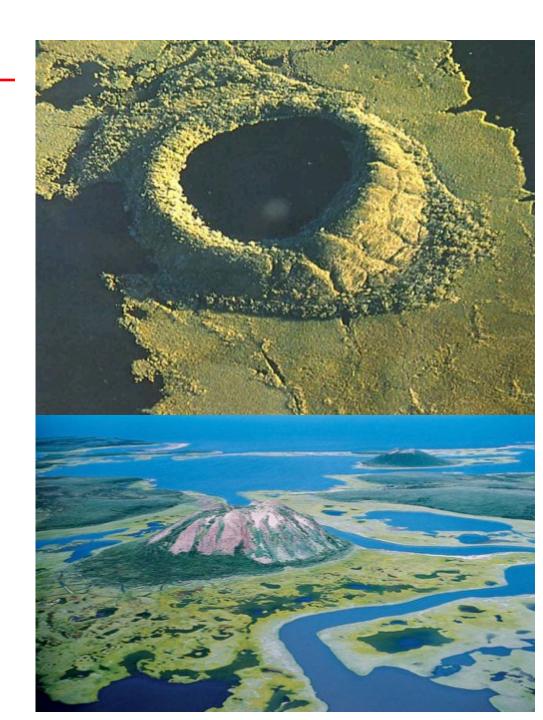




Pingos

Conspicuous conical mound or circular hill, with a core of ice, found on tundra where permafrost is present.

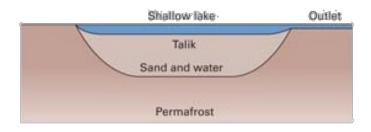
Form under initially unfrozen lakes (taliks).



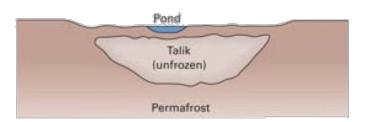
Pingos

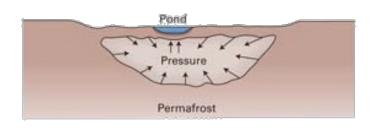
If the lake is infilled with sediment and vegetation, or drained by drainage capture the insulation of the ground surface will increase and permafrost will advance from the sides and bottom of the talik - forcing water up into a zone of refreezing.

Growth of this "pingo ice" pops the overlying sediment up and results in an ice-cored hill.









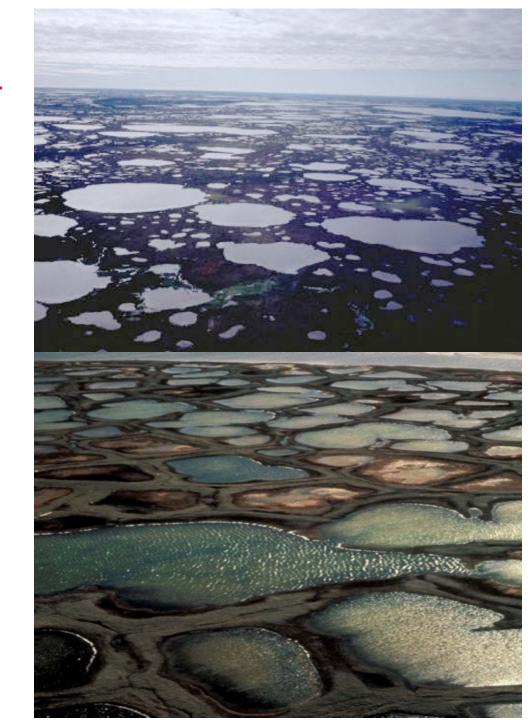


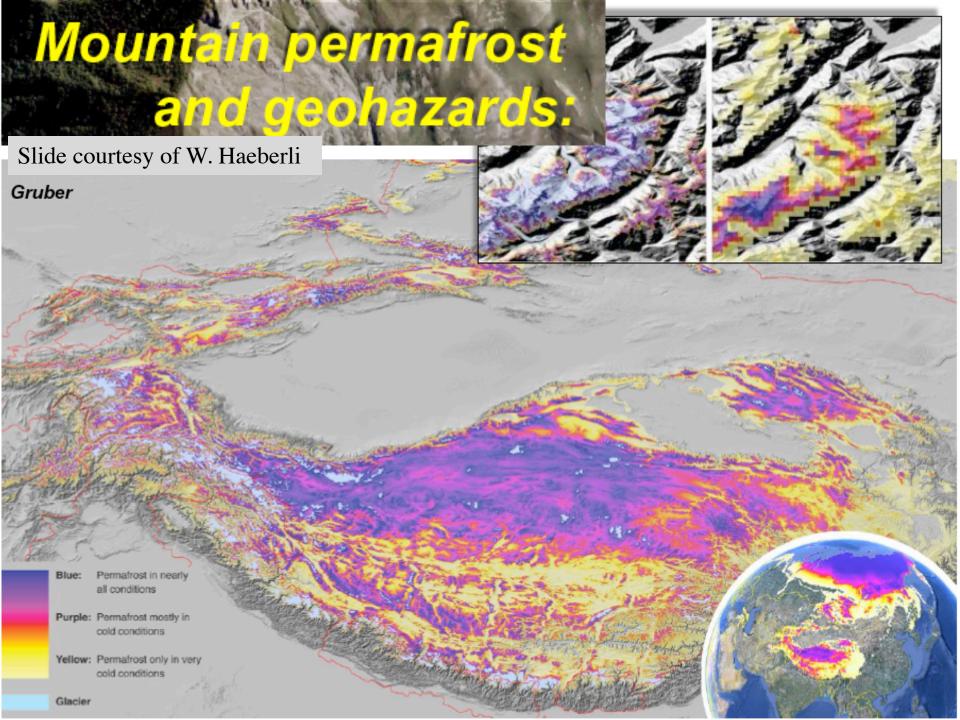
Thermokarst

Characteristic landforms from thawing of ice-rich permafrost or melting of massive ice blocks.



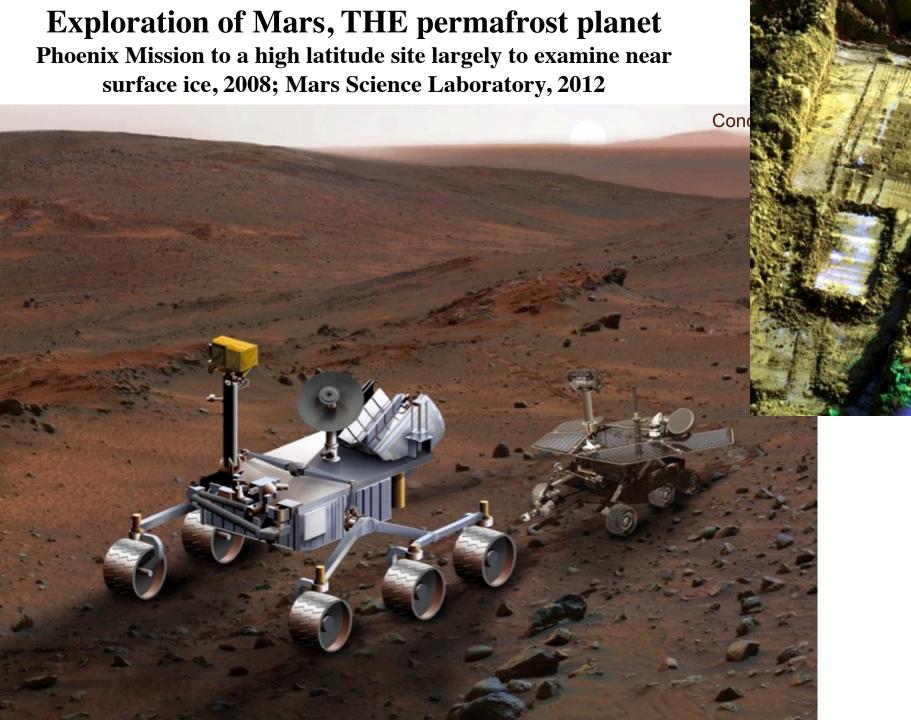
Characterized by irregular topography as irregular pits and depressions develop by thaw settling.





Changing water glaciers and permafrost in the Alps availability ice volume km^3 80 Boeckli glacial/periglacial dynamics 25 subsurface ice paraglacial dynamics time 2010 2050? Figures above courtesy of W. Haeberli Vulnerable villager &

Vulnerable villager & prayer stones, Khumbu region



Phoenix Mission Site

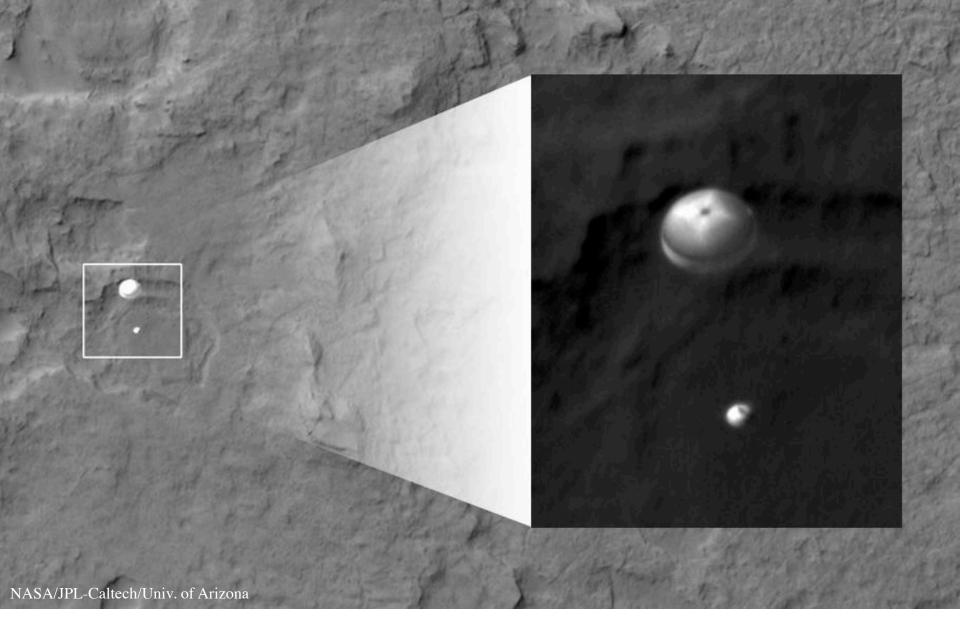




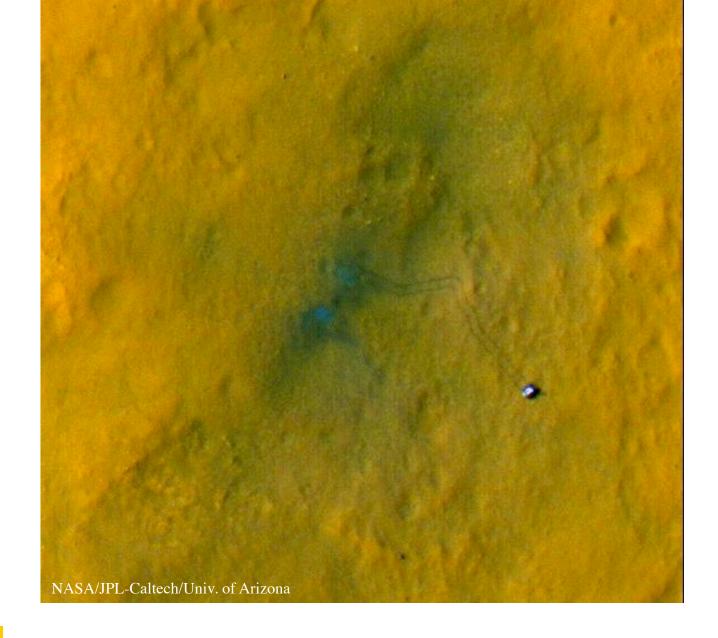






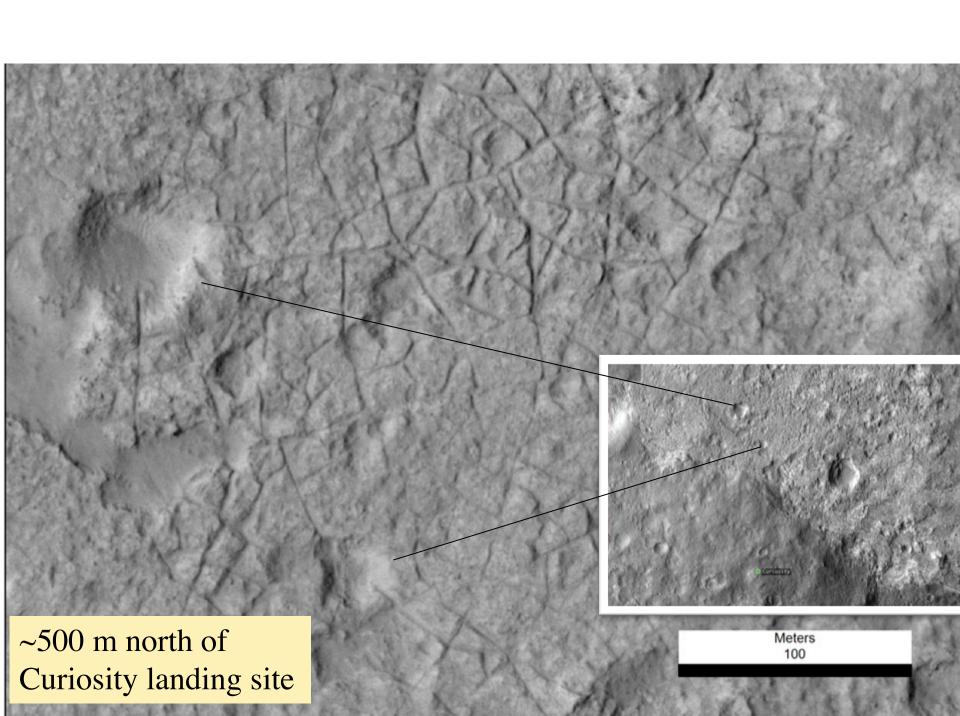




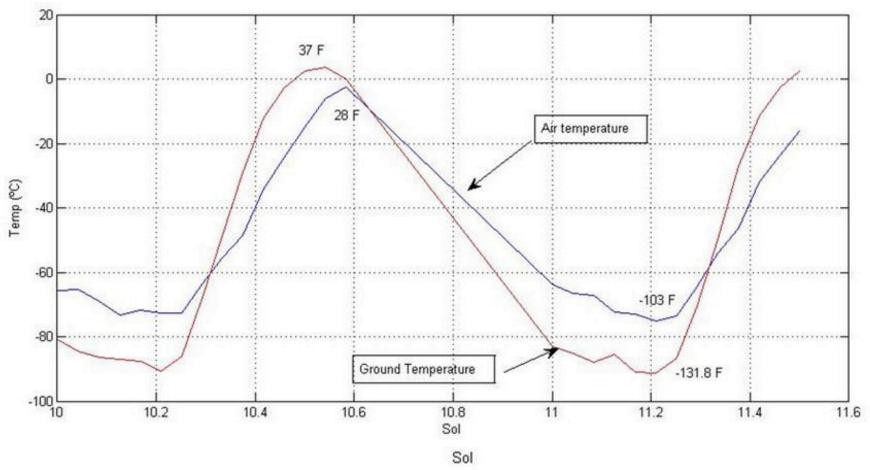




HIRISE images a few days after landing



GROUND AND AIR TEMPERATURE SENSOR



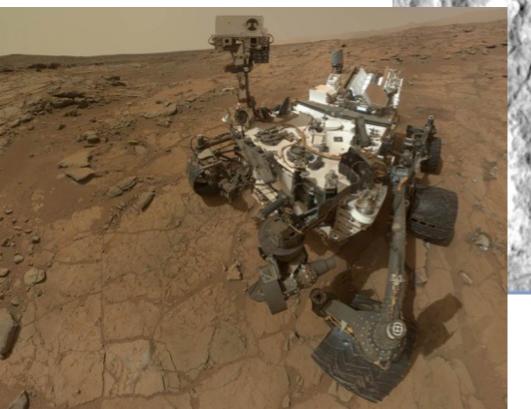


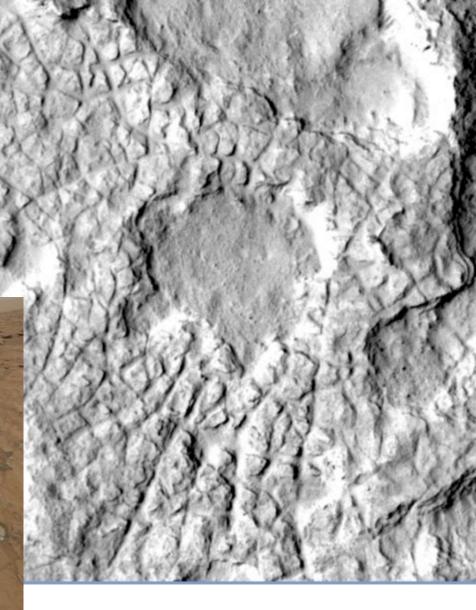
Permafrost prevails even at equator, mean T about - 50° C

Fracture patterns ubiquitous on Mars

But is contraction due to cooling or drying?







20 m polygons at Gale, Mars

Summary

Rich issues:

Thermal regime & the warming Arctic: complex system

Soil is well known to expand, heave, and crack as it freezes because of molecular scale properties of H₂O in confined spaces, and thermal stresses

Patterned ground: spectacular expressions of abiotic self-organization in Nature

Other characteristic periglacial landforms

Ice & Frozen ground on Mars: Earth provides instructive analogs

Useful references

Taber, S. 1929. Frost heaving, Journal of Geology, 37, 428-461.

Rosenberg, R. 2005. Why is ice slippery? Physics Today, 50-55.

Wettlaufer, J. S. and J.G. Dash, 2000, Melting below zero, Scientific American, 282(2), 50-53.

Stone Circles, Spitsbergen

